

Patrick Launeau

Quantitative Image Analysis off Minerals and Rocks

Fabric analysis

- Shape preferred orientation (SPO) vs. strain quantification.
- Intercepts in digital images : a tool to analyze interconnection of grains in rocks vs. inertia tensor of individualized grains
- SPO vs Spatial distribution (Fry)
- Ellipsoid of SPO and strain by combining 3 \perp images.



Ellipsoid2003



Intercepts2003



SPO2003

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Ellipsoid2003

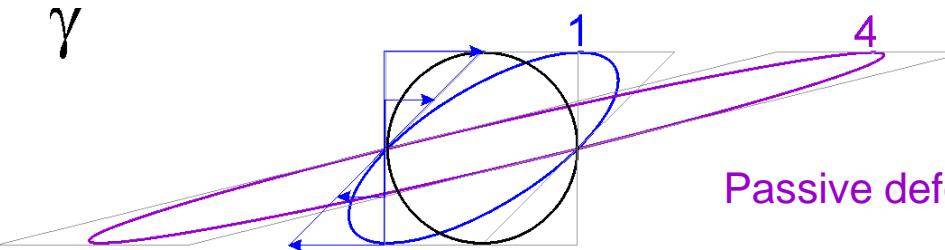


Intercepts2003

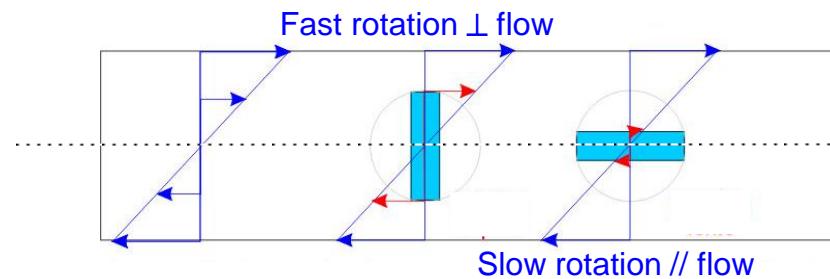


SPO2003

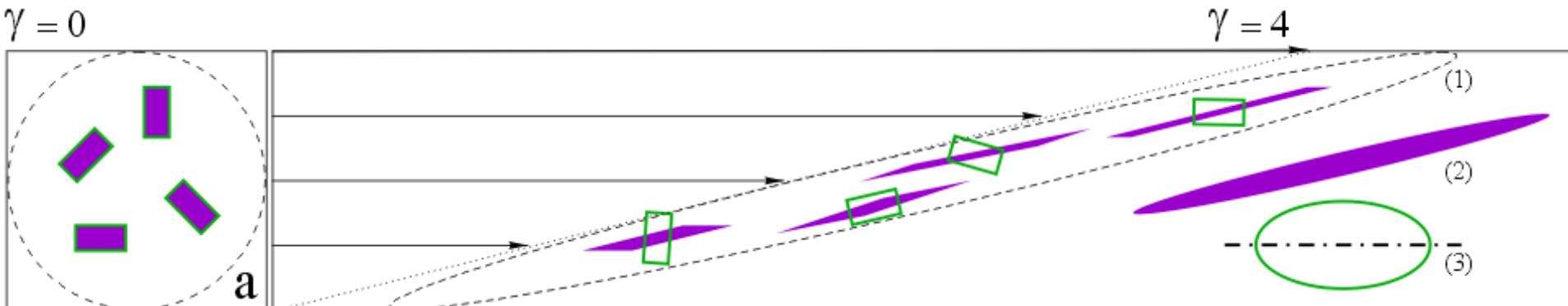
Passive / active deformation



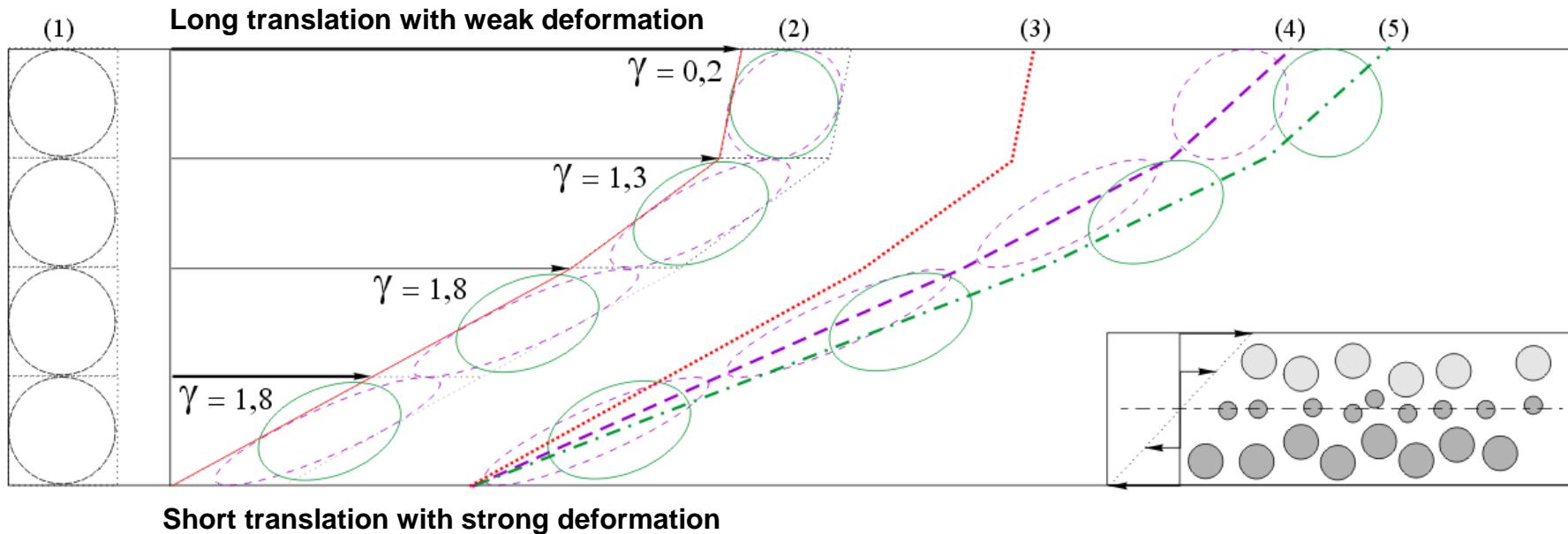
Passive deformation of an enclave
(with elongation of the shape)



Preferred Orientation of
rigid objects
(with out any change of the shape)



In a more realistic magma flow ...

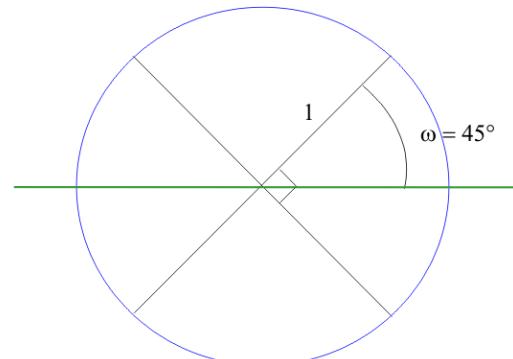


Short translation with strong deformation

- (1) Initial position of a magma bubble
- (2) Final position of its vertical section in red, its shape in purple, the preferred orientation of its microlithes in green
- (3) Alignment of the initial vertical sections of the magma bubbles
- (4) Alignment of the ellipses of magma deformed passively
- (5) Alignment of the ellipses of crystal preferred orientation

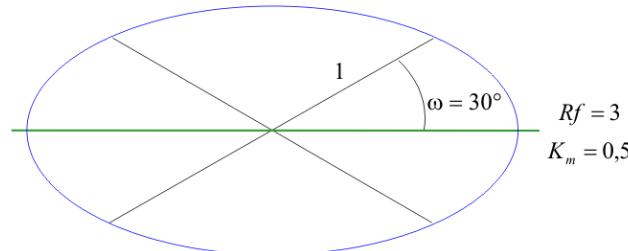
Preferred Orientation (PO)

Cosines directions (Harvey & Laxton, 1980)



$$\begin{aligned} Rf &= 1 \\ K_m &= 0 \end{aligned}$$

$$\mathbf{M} = \frac{1}{N} \begin{bmatrix} \sum \cos^2 \varphi_i & \sum \cos \varphi_i \sin \varphi_i \\ \sum \sin \varphi_i \cos \varphi_i & \sum \sin^2 \varphi_i \end{bmatrix}$$



$$\begin{aligned} Rf &= 3 \\ K_m &= 0,5 \end{aligned}$$

$$\mathbf{M} = \begin{bmatrix} \cos \varphi & \sin \varphi \\ -\sin \varphi & \cos \varphi \end{bmatrix} \cdot \begin{bmatrix} \sqrt{Rf} & 0 \\ 0 & 1/\sqrt{Rf} \end{bmatrix} \cdot \begin{bmatrix} \cos \varphi & -\sin \varphi \\ \sin \varphi & \cos \varphi \end{bmatrix}$$

$$Rf = \frac{\cos^2 \omega}{\sin^2 \omega} = \frac{1}{\tan^2 \omega}$$

$$K_m = \frac{Rf - 1}{Rf + 1}$$

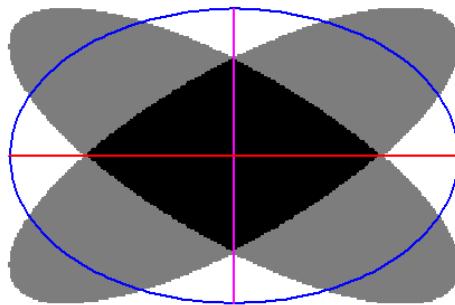
$$\begin{aligned} \omega &= 0 & Rf &= \infty \\ K_m &= 1 \end{aligned}$$

Shape Preferred Orientation (SPO)

SPO → of shapes with long and short axes

A=2, a=5,9926 cm b=3,9318 cm R=1,524 [1,729]b , 89,99°
K=0,398, Kn=0,498 (0,799), Kb=0,501 (0,795)

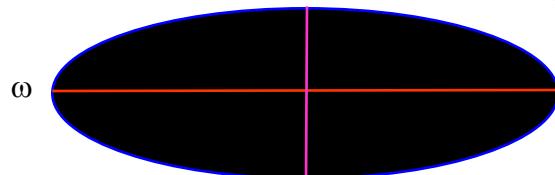
$$R=1,5 \quad \sqrt{Rf} = 1,7$$



$$R \leq \sqrt{Rf} \quad \text{when } r \geq 10$$

$$R < \sqrt{Rf} \quad \text{when } 1 < r < 10$$

$$\omega + \pi/2 \quad R=3 \quad \sqrt{Rf} = \infty$$



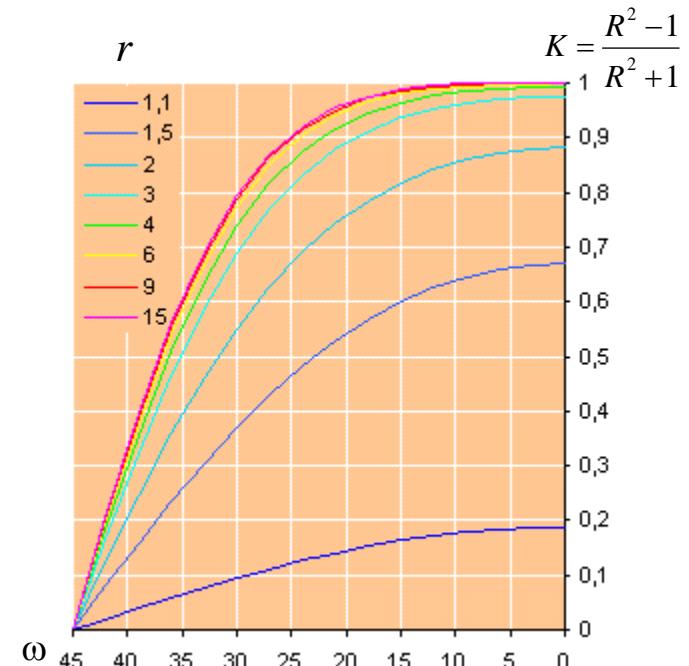
$$R = \frac{d(\omega)}{d(\omega + \pi/2)}$$

Normalization to k

$$k = \frac{r^2 - 1}{r^2 + 1}$$

$$K_n = K / k$$

$$R_n = \sqrt{\frac{1+K_n}{1-K_n}}$$



2D magma flow

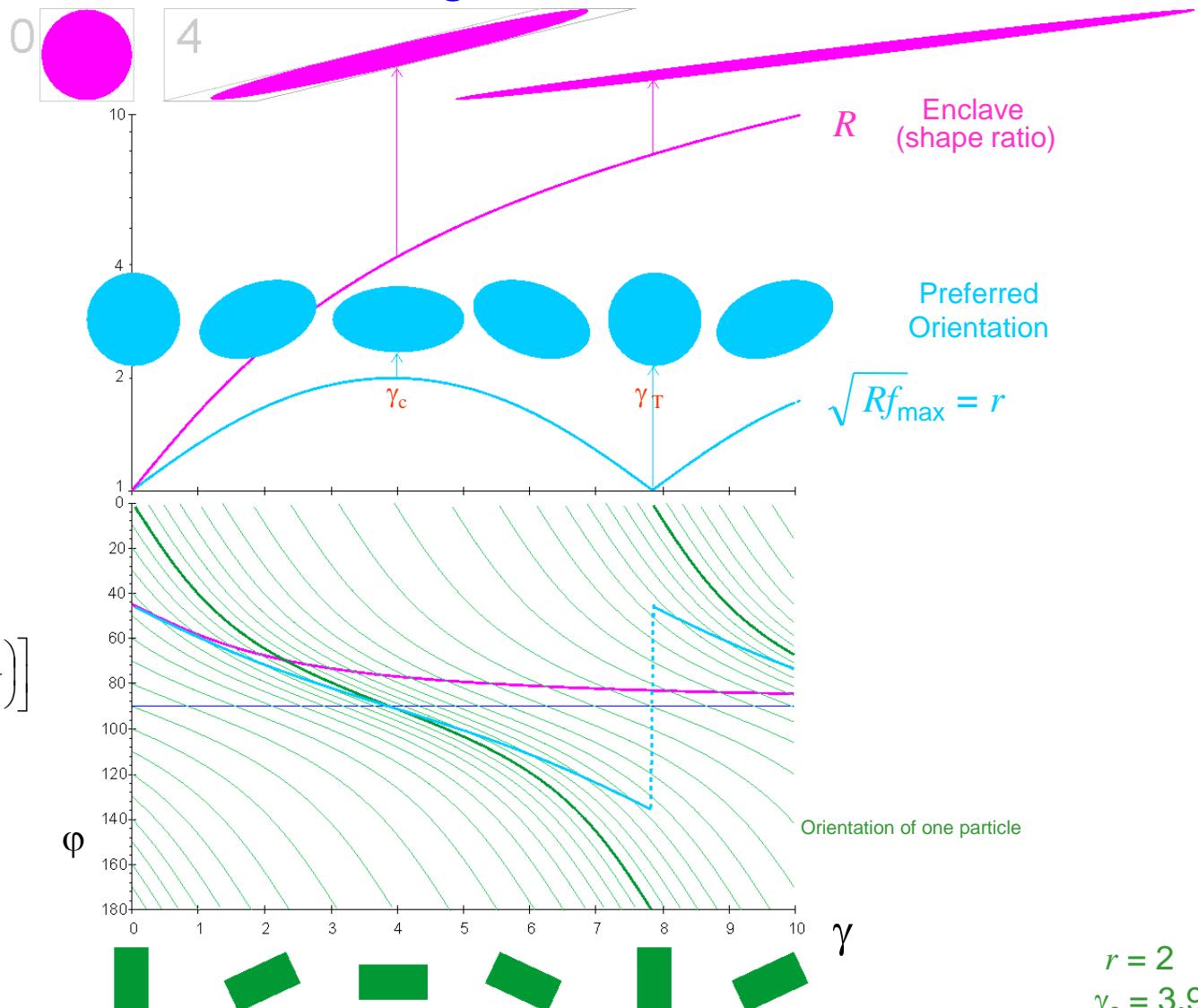
SPO with
strain record
(passive deformation)

SPO without
strain record
(active deformation)

Jeffery 1922

$$\tan \beta' = r \cdot \tan \left[\frac{r \cdot \gamma}{r^2 + 1} + \arctan \left(\frac{\tan \beta}{r} \right) \right]$$

$$\gamma_c = \frac{\pi}{\sqrt{1-k^2}} \quad k = \frac{r^2 - 1}{r^2 + 1}$$



$$r = 2$$

$$\gamma_c = 3.93$$

Simple shear in 3D

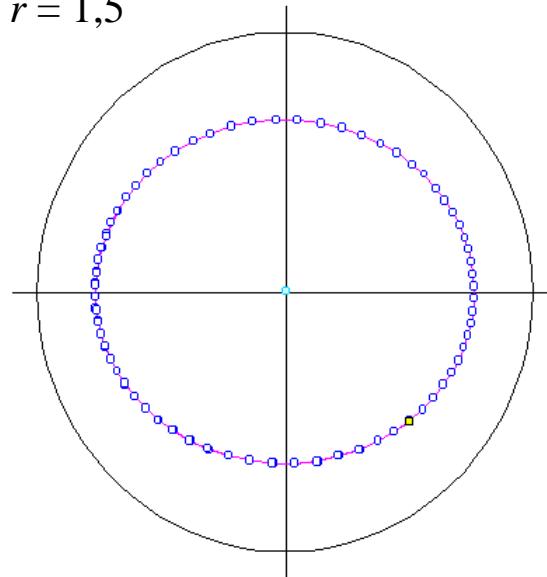
3D

Jeffery (1922), Reed and Tryggvason (1974) et Willis (1977)

$$\tan \beta' = r \cdot \tan \left[\frac{r \cdot \gamma}{r^2 + 1} + \arctan \left(\frac{\tan \beta}{r} \right) \right]$$

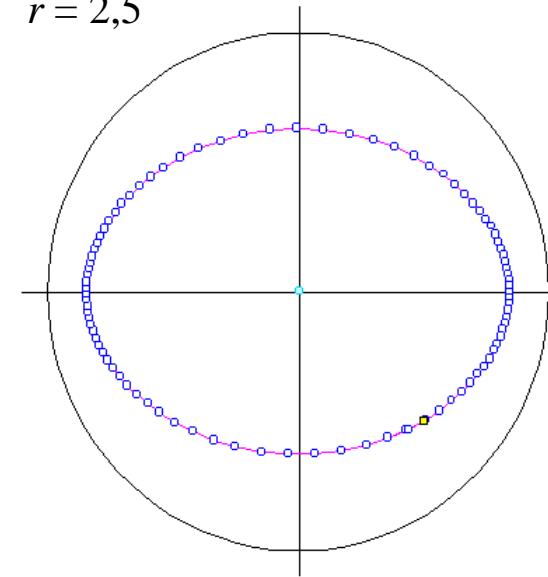
$$\tan^2 \psi' = \tan^2 \psi \cdot \left(\frac{r \cdot \cos^2 \beta + \sin^2 \beta}{r \cdot \cos^2 \beta' + \sin^2 \beta'} \right)$$

$r = 1,5$

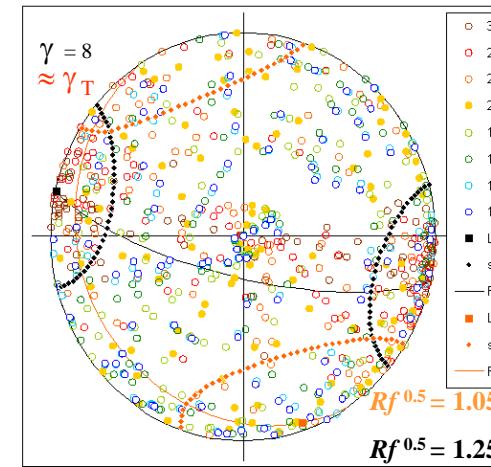
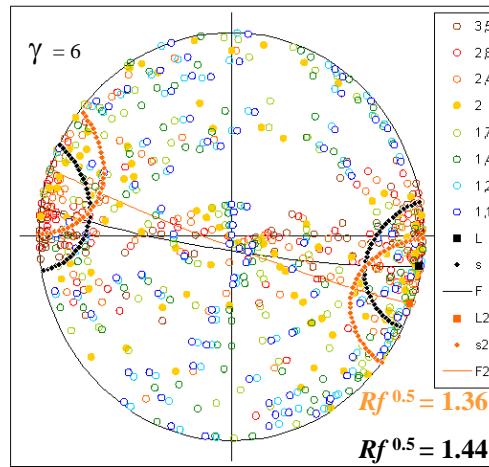
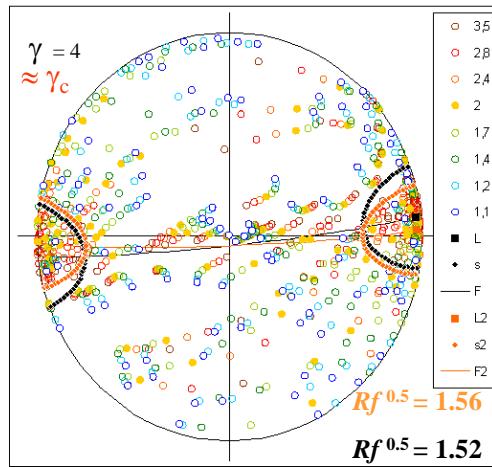
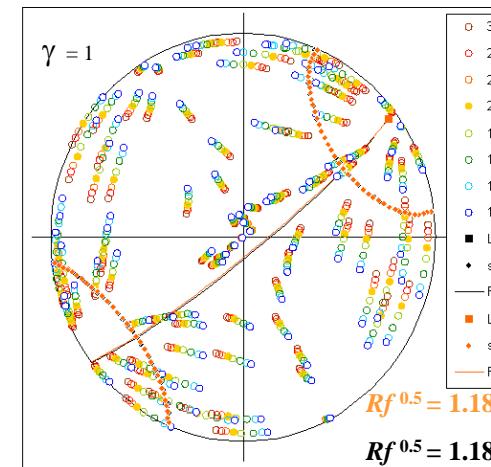
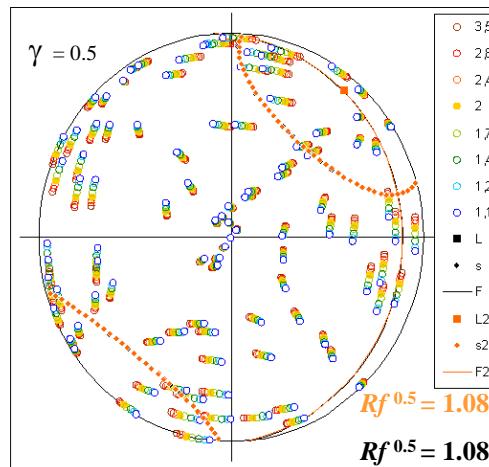
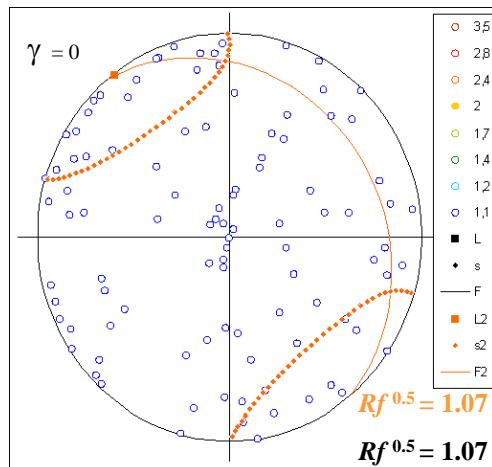


1,5 : 1 : 1

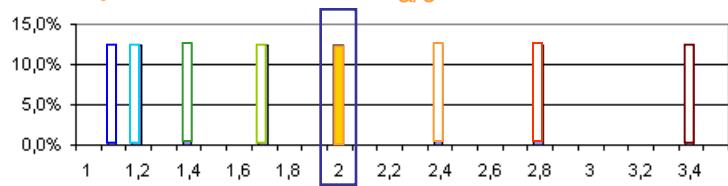
$r = 2,5$



2,5 : 1 : 1



Ellipsoid 2:1:1 $r_{a/c} = 2$



The PO of one class of aspect ratio is cyclic
The PO of standard CSD is not cyclic

Jeffery (1922), Reed and Tryggvason (1974) et Willis (1977)

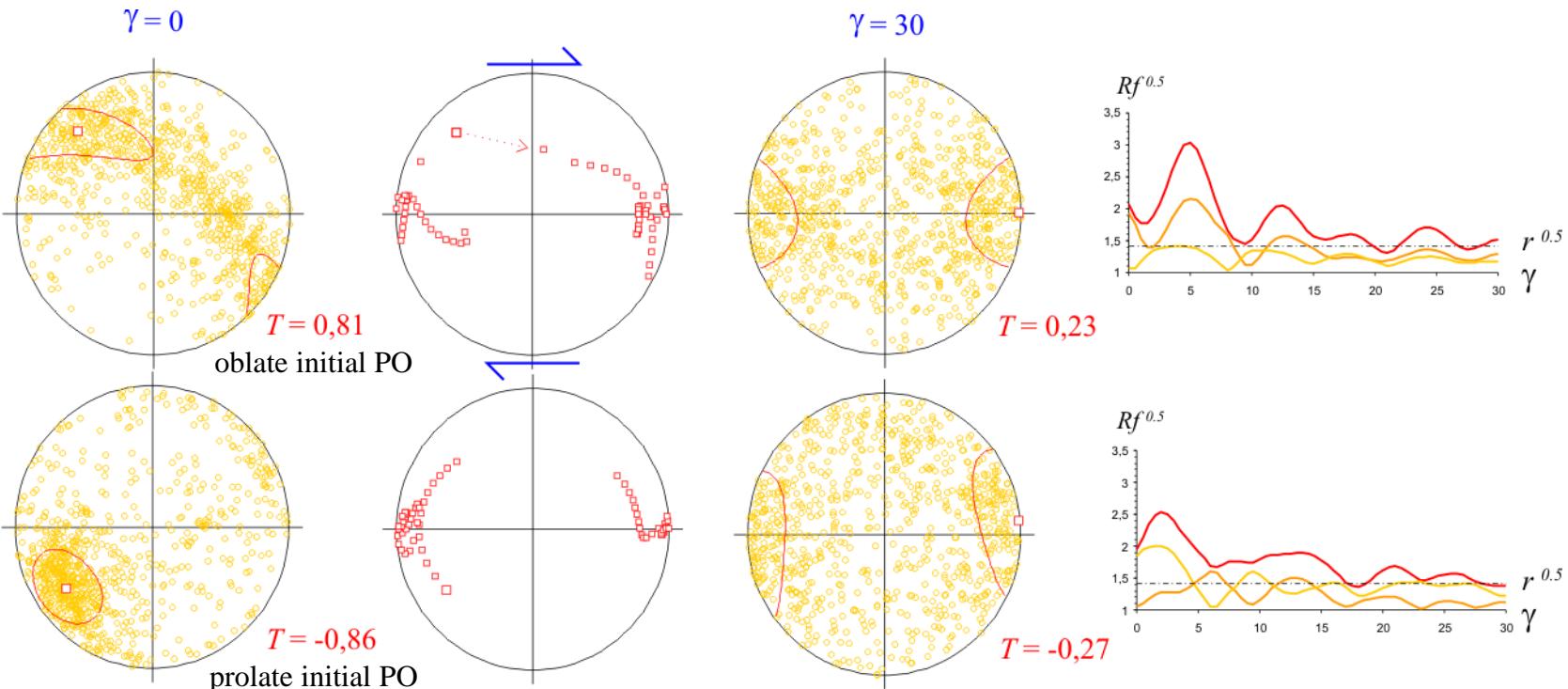
$$\tan \beta' = r \cdot \tan \left[\frac{r \cdot \gamma}{r^2 + 1} + \arctan \left(\frac{\tan \beta}{r} \right) \right]$$

$$\tan^2 \psi' = \tan^2 \Psi \cdot \left(\frac{r \cdot \cos^2 \beta + \sin^2 \beta}{r \cdot \cos^2 \beta' + \sin^2 \beta'} \right)$$

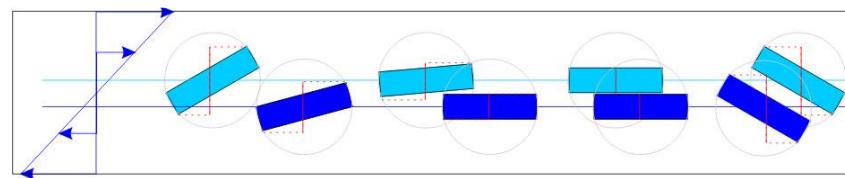
Ellipsoid 2:1:1

$r_{a/c} = 2$

SPO with strong initial SPO tend toward a weak SPO along a magma flow



Passive / active deformation



Particle interactions
may reduce the object
rotation

Patrick Launeau

Quantitative Image Analysis off Minerals and Rocks

Fabric analysis

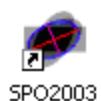
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Ellipsoid2003



Intercepts2003

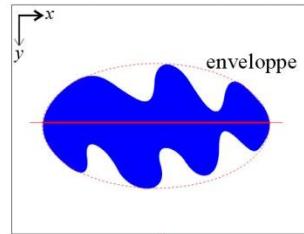


SPO2003

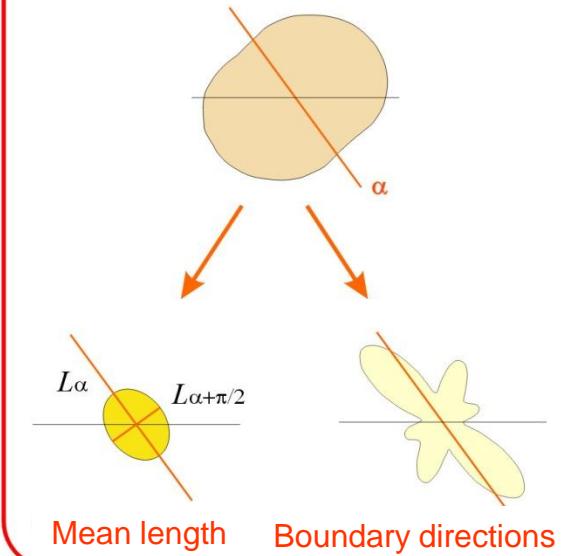
Image analysis

First hypothesis :
do we have
isolated crystals
or aggregates ?

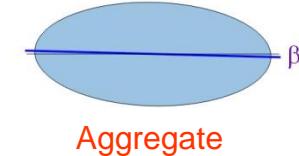
Aggregate



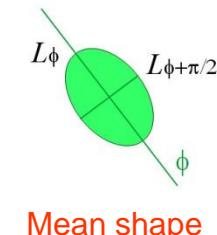
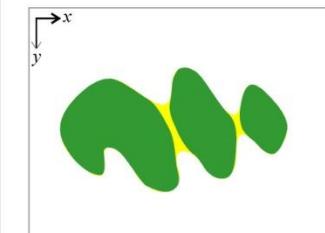
Intercepts method



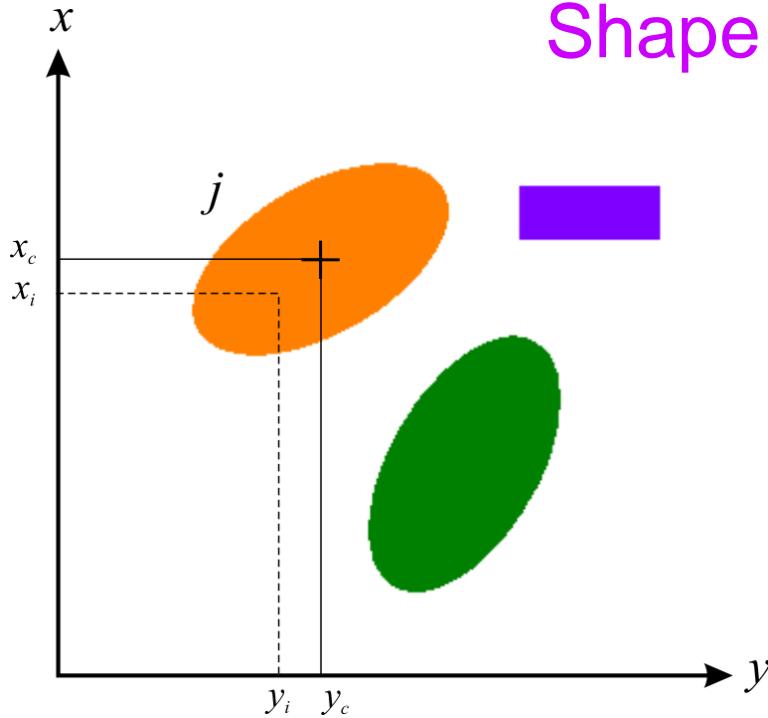
Inertia tensor method



Grain segmentation



Mean shape



Shape analysis with inertia tensor

Case of 1 object

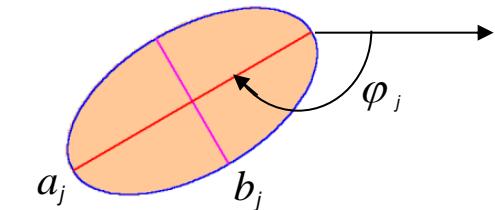
$$x_c = \frac{1}{A} \sum_i x_i$$

$$y_c = \frac{1}{A} \sum_i y_i$$

$$m_{xxj} = \frac{1}{A} \sum_i (x_i - x_c)^2$$

$$m_{xyj} = \frac{1}{A} \sum_i (x_i - x_c)(y_i - y_c)$$

$$m_{yyj} = \frac{1}{A} \sum_i (y_i - y_c)^2$$



$$r_j = \frac{a_j}{b_j}$$

$$\mathbf{M}_j = \begin{vmatrix} m_{xxj} & m_{xyj} \\ m_{xyj} & m_{yyj} \end{vmatrix}$$

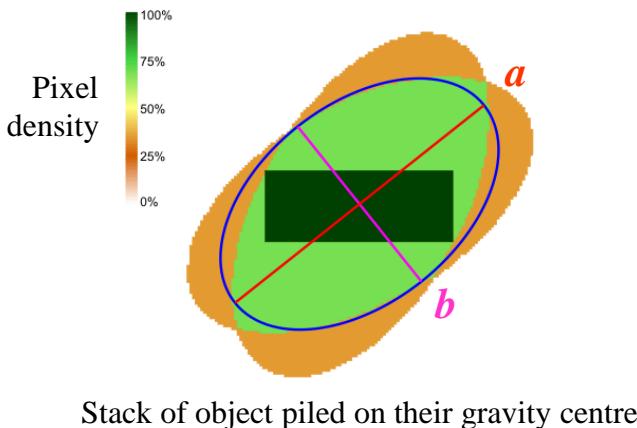
$$\mathbf{M}_j = \begin{bmatrix} \cos \varphi_j & \sin \varphi_j \\ -\sin \varphi_j & \cos \varphi_j \end{bmatrix} \cdot \begin{bmatrix} a_j^2/4 & 0 \\ 0 & b_j^2/4 \end{bmatrix} \cdot \begin{bmatrix} \cos \varphi_j & -\sin \varphi_j \\ \sin \varphi_j & \cos \varphi_j \end{bmatrix}$$

Rink, M. (1976). – A computerized quantitative image analysis procedure for investigating features and an adapted image process. *Journal of Microscopy*, 107: 267-286

Case of N objects

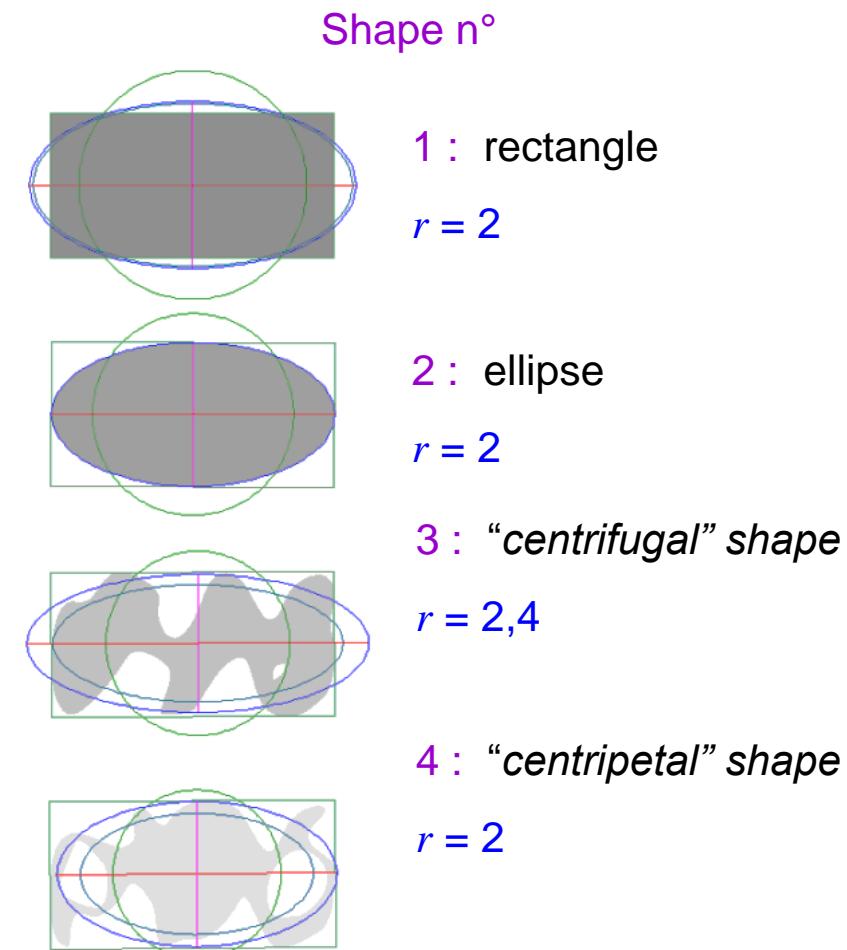
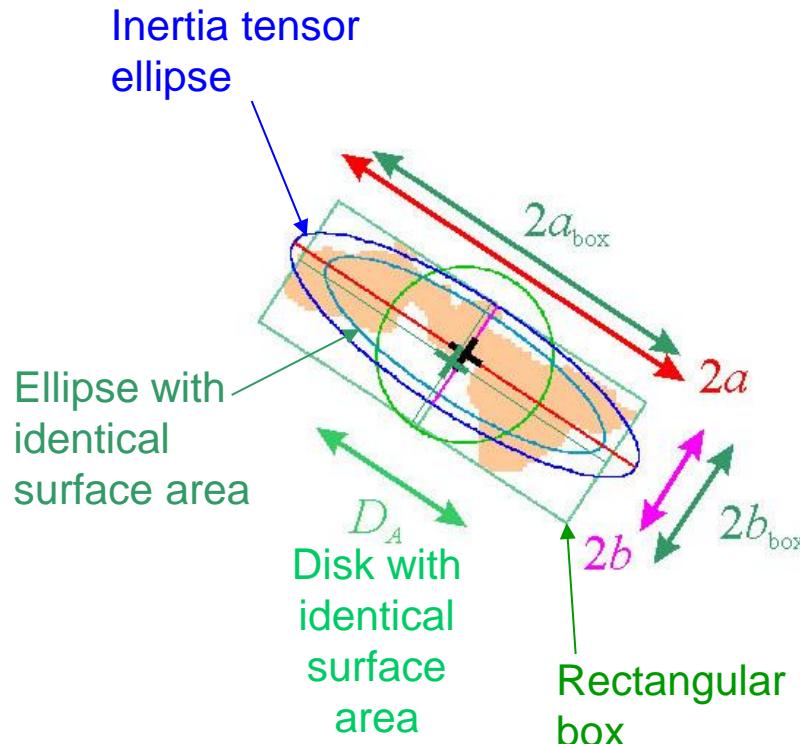
$$\mathbf{M} = \frac{1}{N} \sum_j \mathbf{M}_j = \frac{1}{N} \begin{vmatrix} \sum_j m_{xxj} & \sum_j m_{xyj} \\ \sum_j m_{xyj} & \sum_j m_{yyj} \end{vmatrix}$$

$$R = \frac{a}{b} = \sqrt{\frac{\rho_1}{\rho_2}}$$



P. Launeau
(2004) "Mise en évidence des écoulements magmatiques par analyse d'images 2-D des distributions 3-D d'orientations Préférentielles de Formes". *Bull. Soc. Géol. Fr.*, 175, 331-350

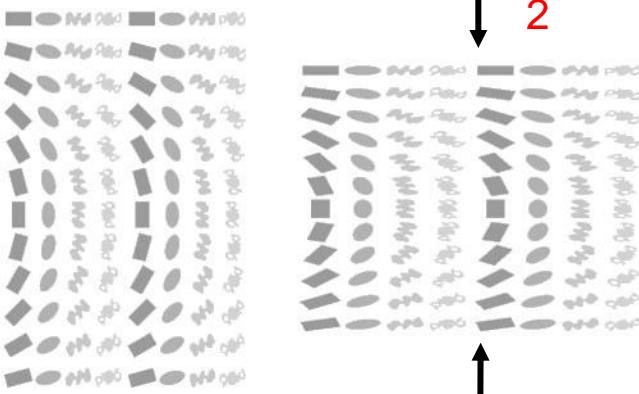
Other shape parameters ...



Quantitative Image Analysis of Minerals and Rocks

<http://www.sciences.univ-nantes.fr/lpgnantes/SPO>

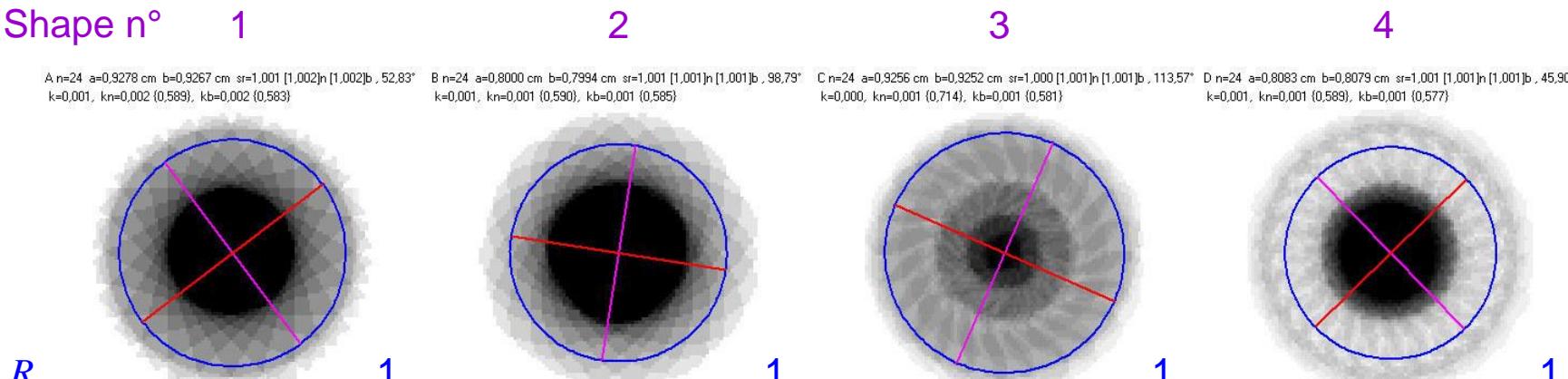
2



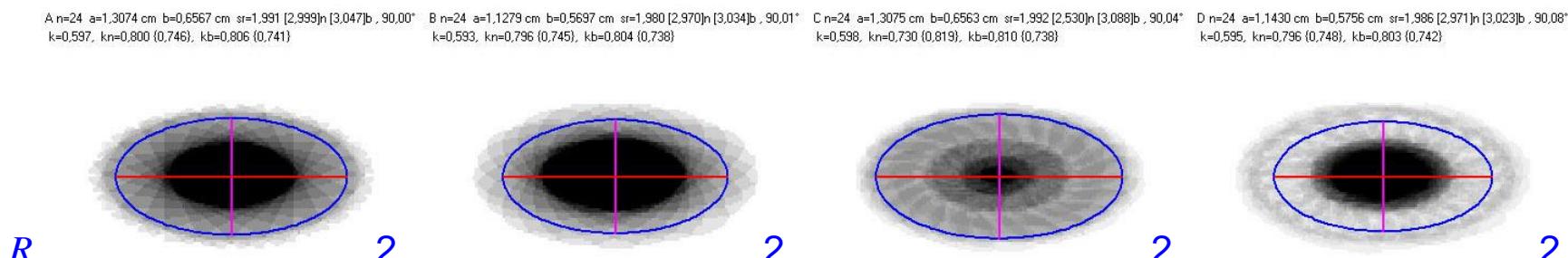
Passive Deformation no viscosity contrast

Shape n° 1

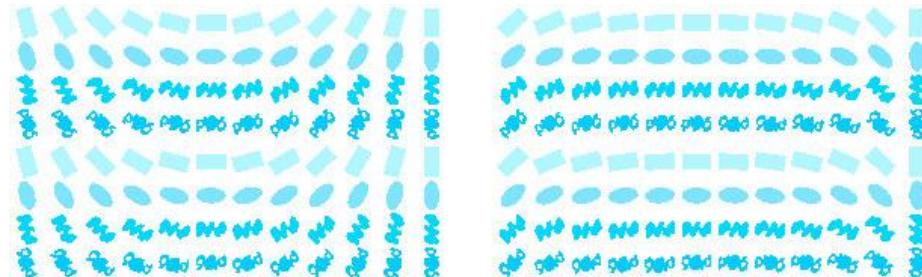
A n=24 a=0.9278 cm b=0.9267 cm sr=1,001 [1,002]n [1,002]b , 52,83° k=0,001, kn=0,002 (0,589), kb=0,002 (0,583)
B n=24 a=0,8000 cm b=0,7994 cm sr=1,001 [1,001]n [1,001]b , 98,79° k=0,001, kn=0,001 (0,590), kb=0,001 (0,585)
C n=24 a=0,9256 cm b=0,9252 cm sr=1,000 [1,001]n [1,001]b , 113,57° k=0,000, kn=0,001 (0,714), kb=0,001 (0,581)
D n=24 a=0,8083 cm b=0,8079 cm sr=1,001 [1,001]n [1,001]b , 45,90° k=0,001, kn=0,001 (0,589), kb=0,001 (0,577)



A n=24 a=1,3074 cm b=0,6567 cm sr=1,991 [2,999]n [3,047]b , 90,00° k=0,597, kn=0,800 (0,746), kb=0,806 (0,741)
B n=24 a=1,1279 cm b=0,5697 cm sr=1,980 [2,970]n [3,034]b , 90,01° k=0,593, kn=0,796 (0,745), kb=0,804 (0,738)
C n=24 a=1,3075 cm b=0,6563 cm sr=1,992 [2,530]n [3,088]b , 90,04° k=0,598, kn=0,730 (0,819), kb=0,810 (0,738)
D n=24 a=1,1430 cm b=0,5756 cm sr=1,986 [2,971]n [3,023]b , 90,08° k=0,595, kn=0,796 (0,748), kb=0,803 (0,742)



Active Deformation rigid body rotation



P.K. Harvey et C.C. Ferguson, 1978

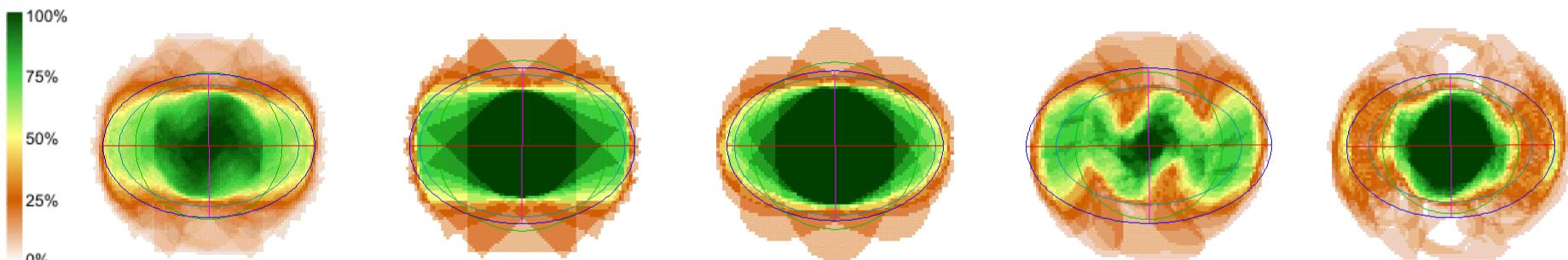
Shape n° 1234

1

2

3

4



$$R\gamma_c \quad 1,48$$

$$R_n \quad 1,99$$

$$r \quad 2$$

$$1,44$$

$$1,95$$

$$2$$

$$1,44$$

$$1,98$$

$$2$$

$$1,57$$

$$1,99$$

$$2,4$$

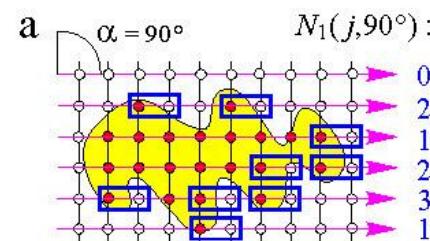
$$1,44$$

$$1,99$$

$$2$$

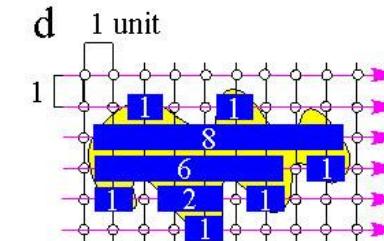
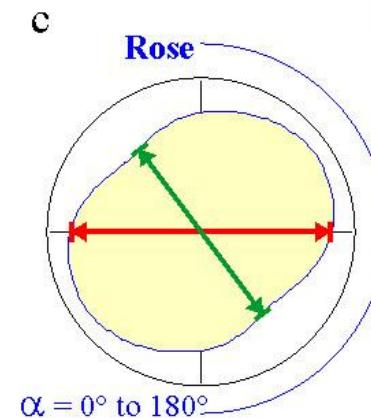
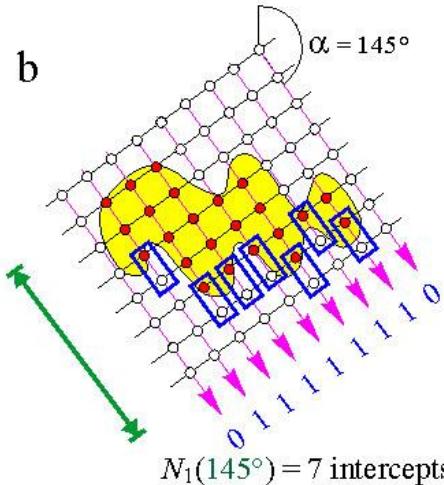
Case of 1 object

Shape analysis with the intercepts



scan line →
 ● in ○ out yellow phase ■ intercept

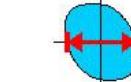
$N_1(90^\circ) = 9$ intercepts



mean length
 $22 \text{ units} / 9 \text{ segments} = 2.4 \text{ units per segment}$

$L(90^\circ) = 2.4$ units

e



mean length **Rose**

Launeau, P. and Robin, P.-Y.F. (1996). –
 Fabric analysis using the intercept method.
Tectonophysics 267, 91-119

Case of 1 object

Shape analysis with the intercepts

Theoretical shapes of roses

Shapes



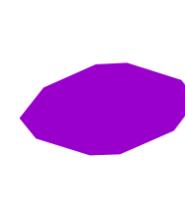
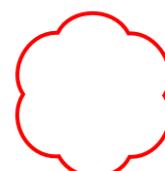
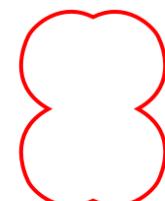
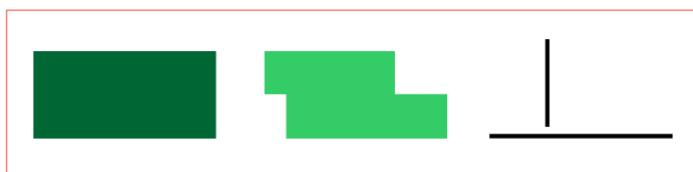
Intercept counts



Mean intercept lengths



shapes
without scale

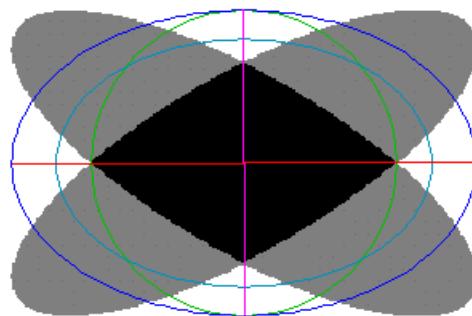


Launeau, P. and Robin, P.-Y.F. (1996). –
Fabric analysis using the intercept method.
Tectonophysics 267, 91-119

Case of N objects

$A = 2$ $a = 5,9926$ cm $b = 3,9318$ cm $R = 1,524$ [1,729]n [1,734]b , 89,99°
 $K = 0,398$, $K_n = 0,498$ (0,799), $K_b = 0,501$ (0,795)

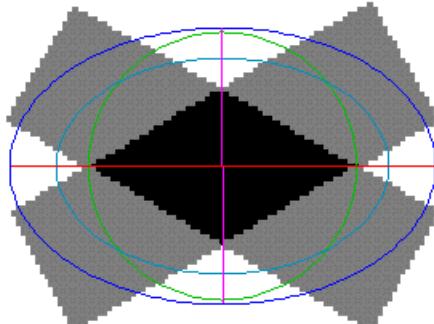
$$r = 3 \quad R = 1,52$$



Inertia tensor

$A = 2$ $a = 0,4883$ cm $b = 0,3171$ cm $R = 1,540$ [1,732]n [1,746]b , 89,82°
 $K = 0,407$, $K_n = 0,500$ (0,814), $K_b = 0,506$ (0,804)

$$r = 3 \quad R = 1,54$$

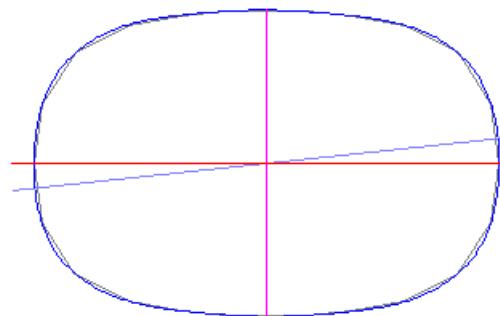


Shape analysis with the intercepts

$A = 3,9357$ cm $b = 2,5886$ cm $R = 1,520$, 90,12°

(1) 84°

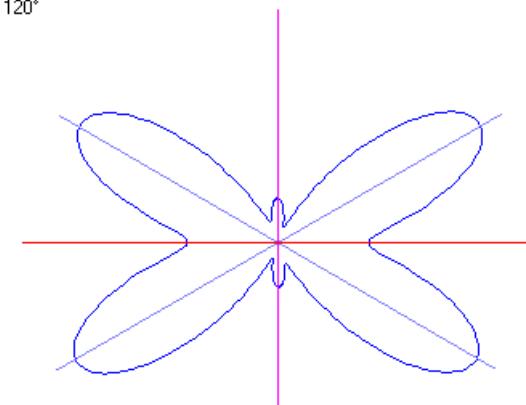
$$R = 1,52$$



Lengths of intercepts

$A = 11,9422$ cm $b = 7,8547$ cm $R = 1,520$, 90,12°

(1) 60°
(2) 120°

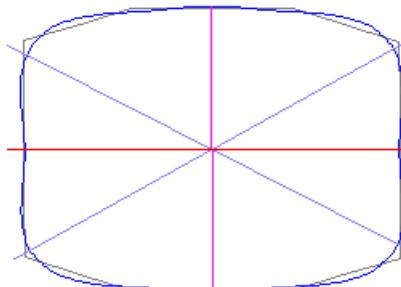


Boundaries directions

$A = 0,2599$ cm $b = 0,1990$ cm $R = 1,306$, 89,98°

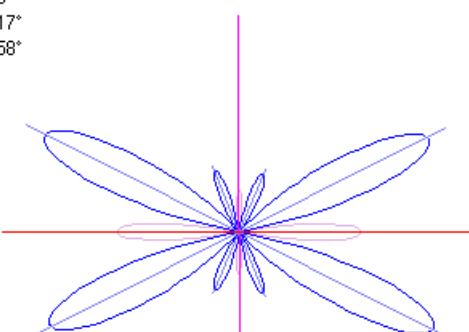
(1) 61°
(2) 117°

$$R = 1,30$$



$A = 0,9649$ cm $b = 0,7434$ cm $R = 1,298$, 89,98°

(1) 23°
(2) 63°
(3) 117°
(4) 158°



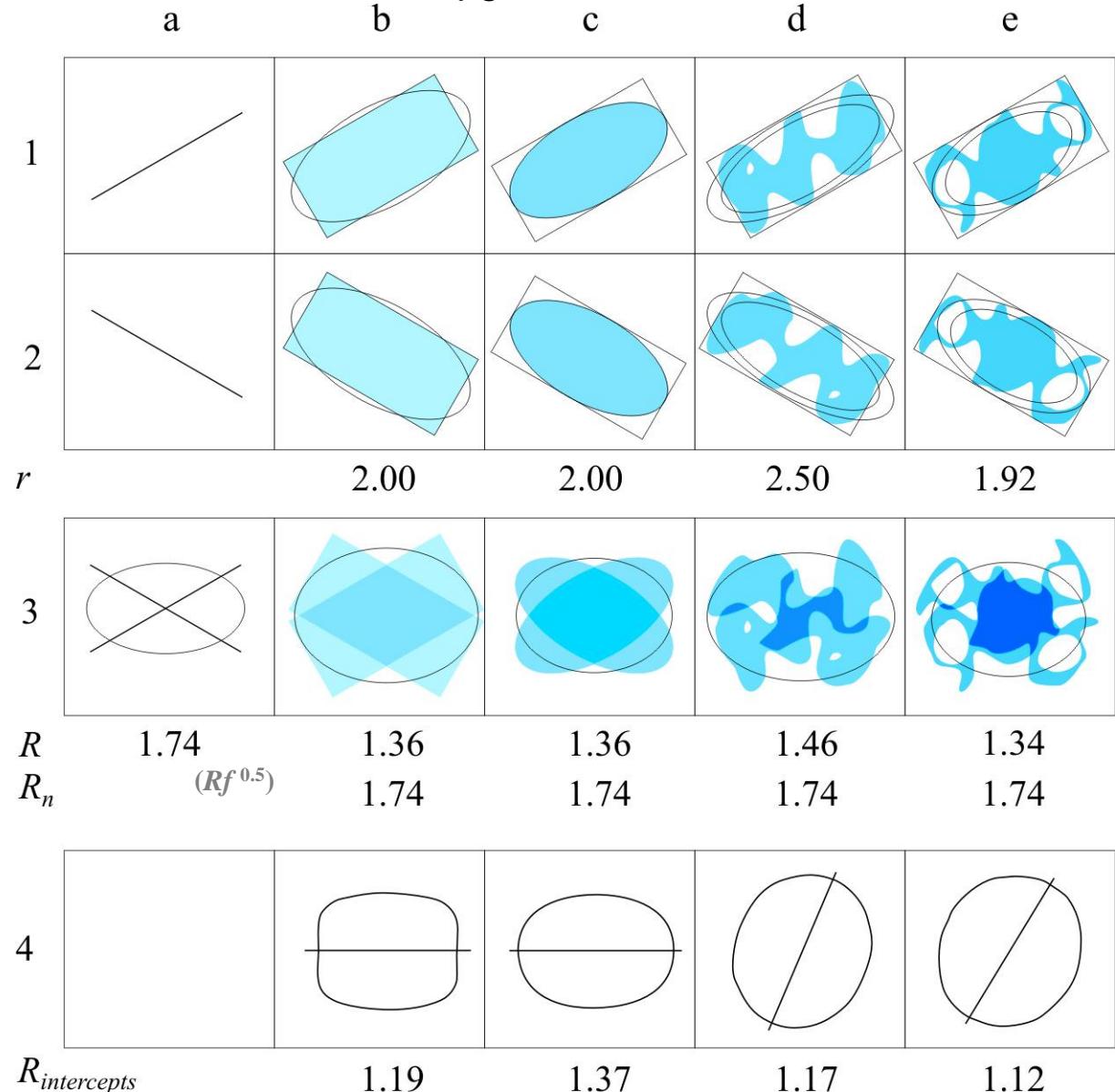
Intercepts are sensitive to the boundary

Launeau, P. and Robin, P.-Y.F. (1996). –
Fabric analysis using the intercept method.
Tectonophysics 267, 91-119

Quantitative Image Analysis of Minerals and Rocks

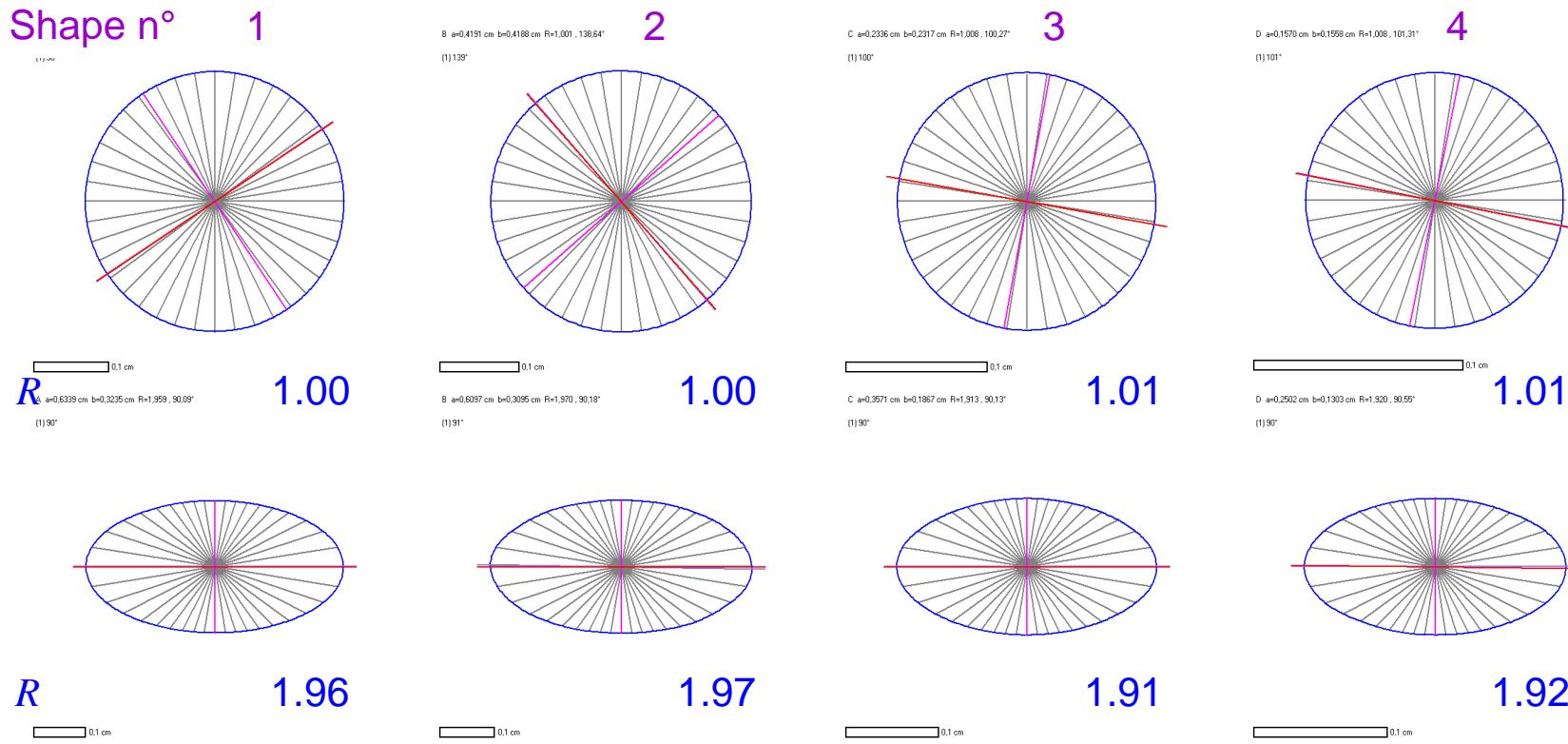
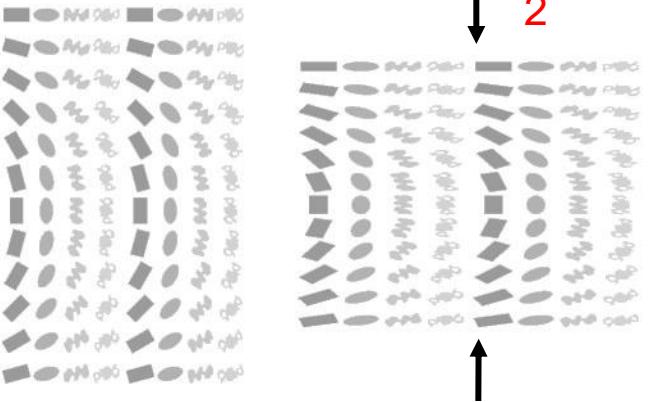
<http://www.sciences.univ-nantes.fr/lpgnantes/SPO>

SPO of
objects
with any
boundary
geometry



Quantitative Image Analysis of Minerals and Rocks

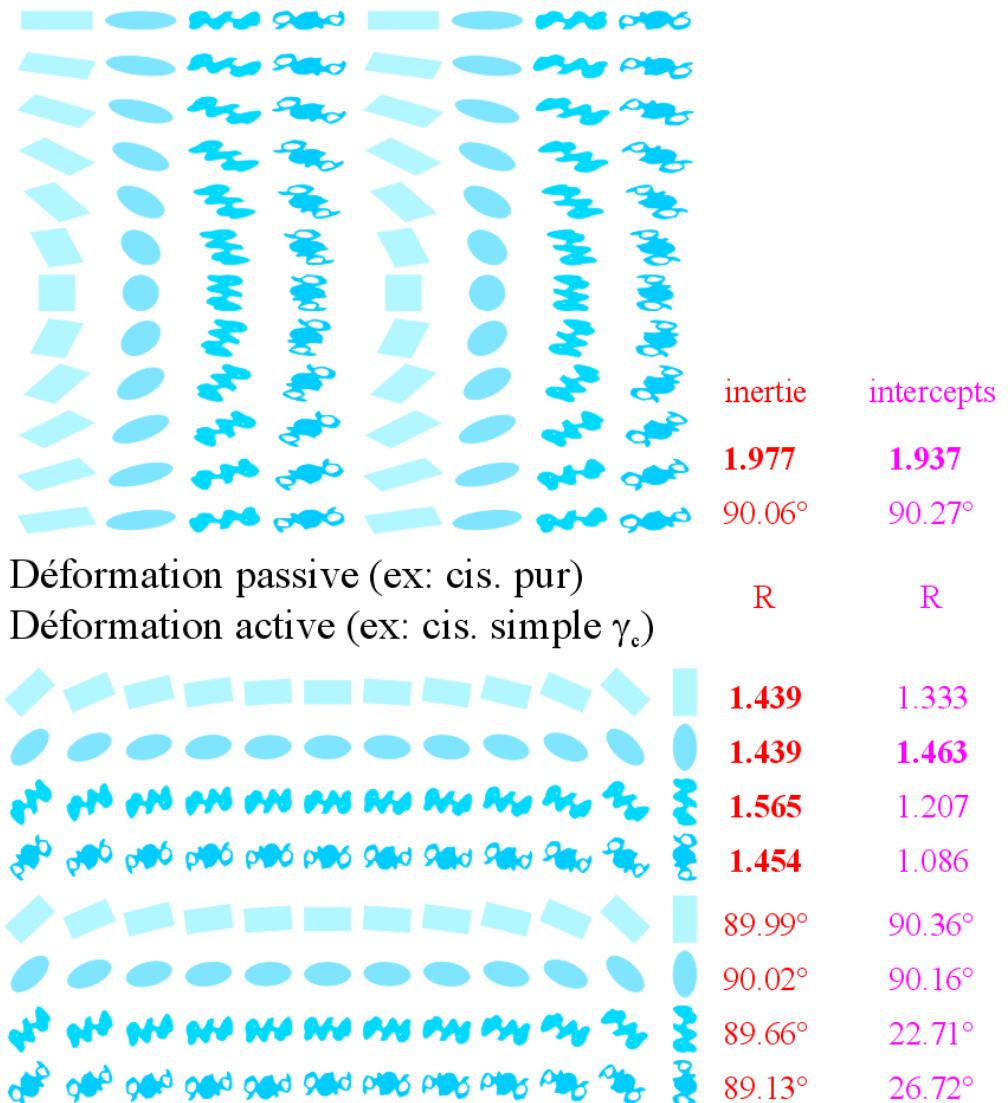
<http://www.sciences.univ-nantes.fr/lpgnantes/SPO>



P. Launeau (2004) "Mise en évidence des écoulements magmatiques par analyse d'images 2-D des distributions 3-D d'Orientations Préférentielles de Formes". *Bull. Soc. Géol. Fr.*, 175, 331-350

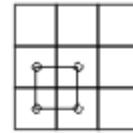
Passive Deformation no viscosity contrast

Passive / Active Deformation

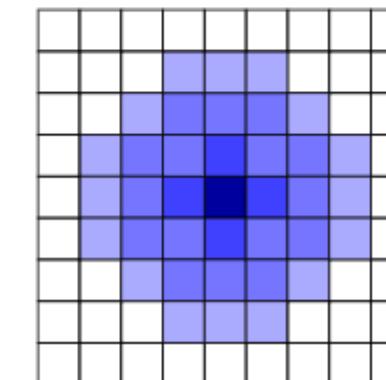
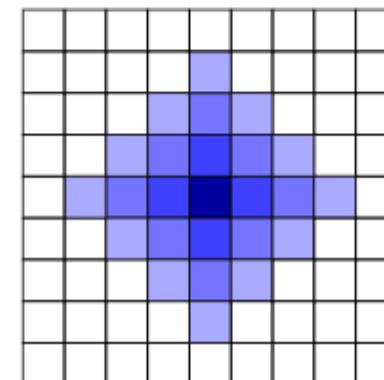
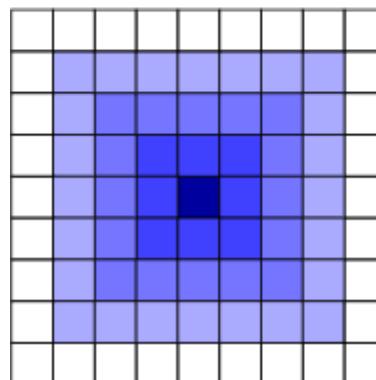
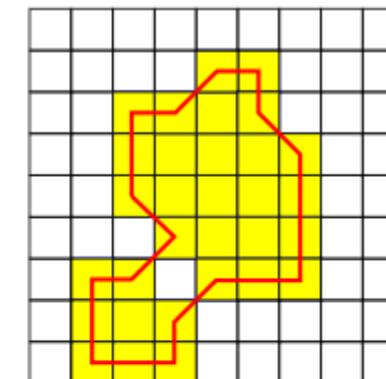
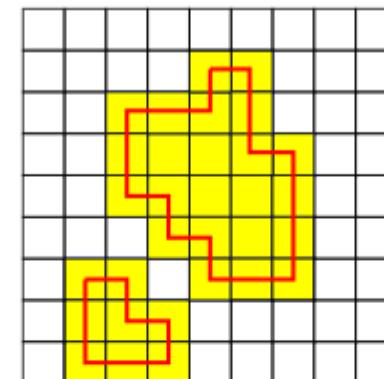
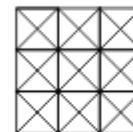
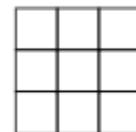
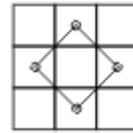


Application to digital images Connexity

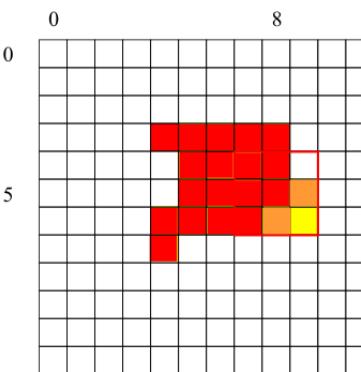
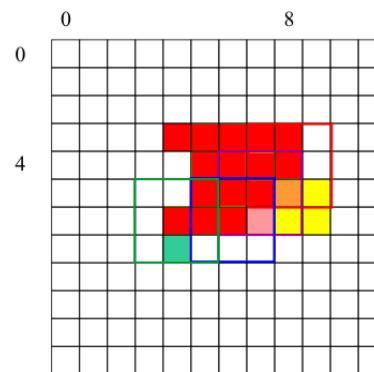
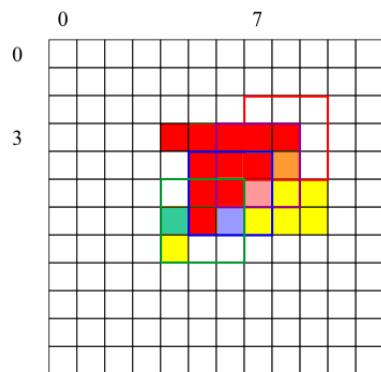
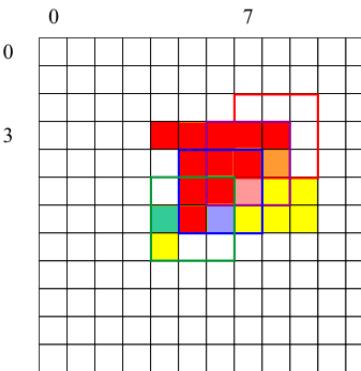
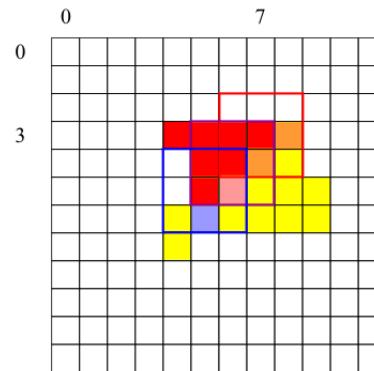
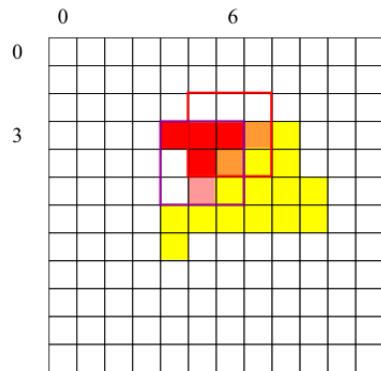
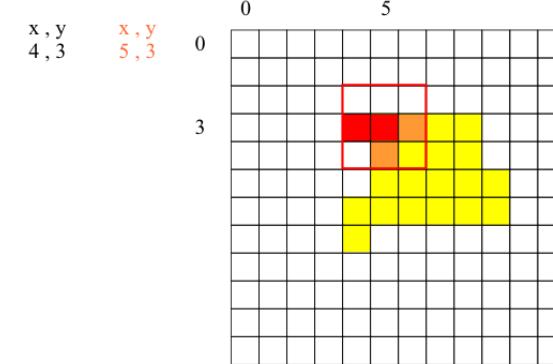
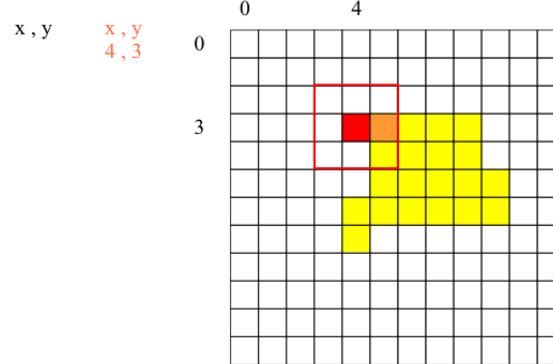
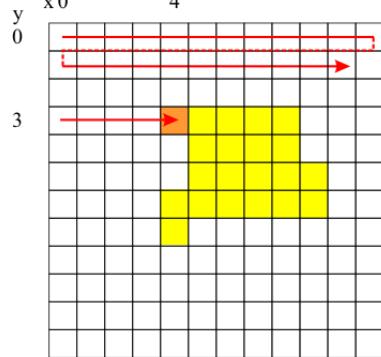
4



8



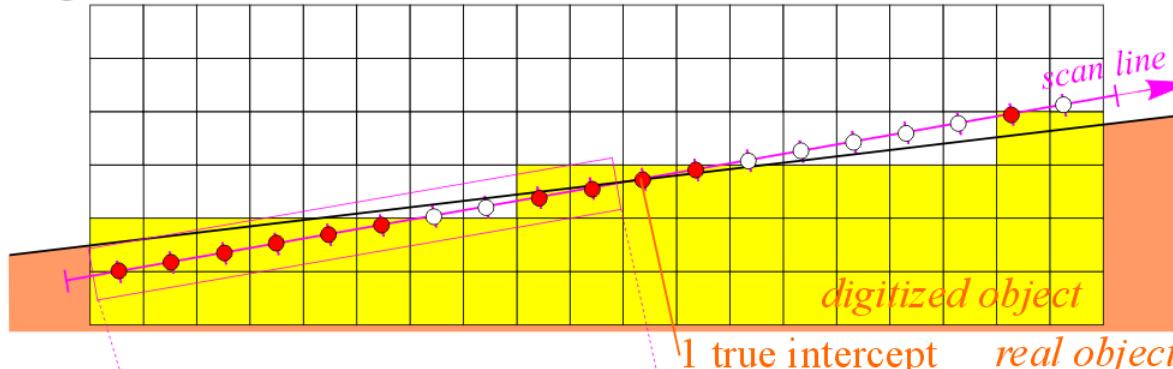
Scanning of one object



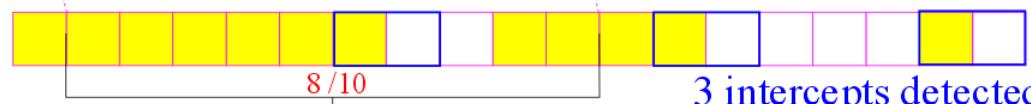
Inertia tensor

Density of intercept count

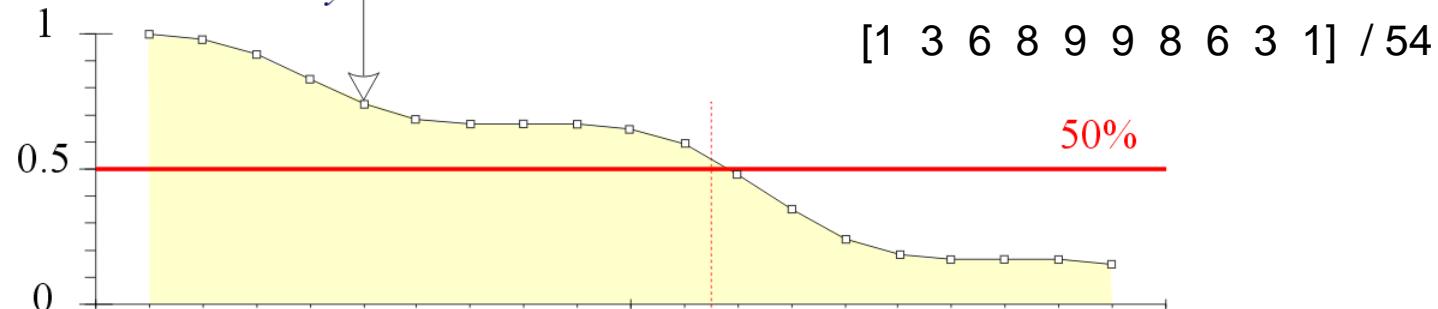
Image



Line



Smoothed local density



New line

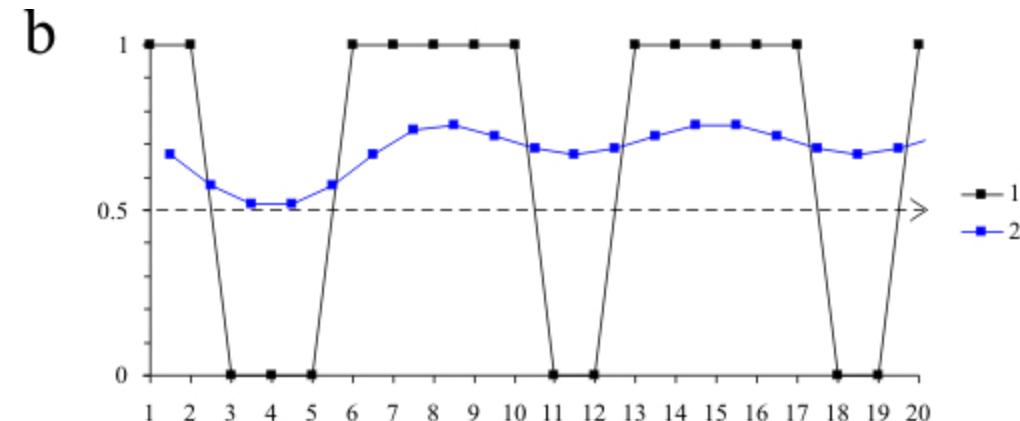
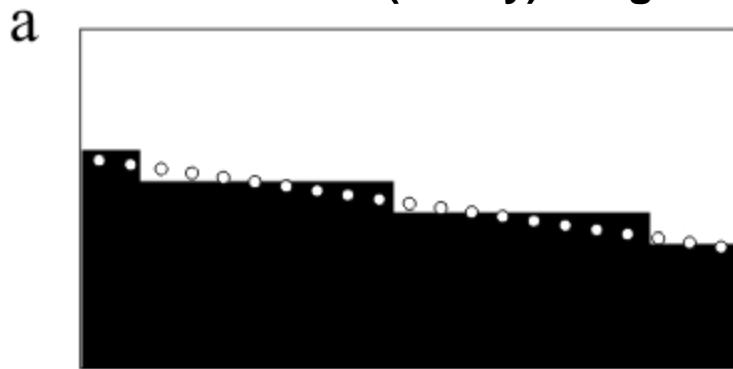


1 intercept detected

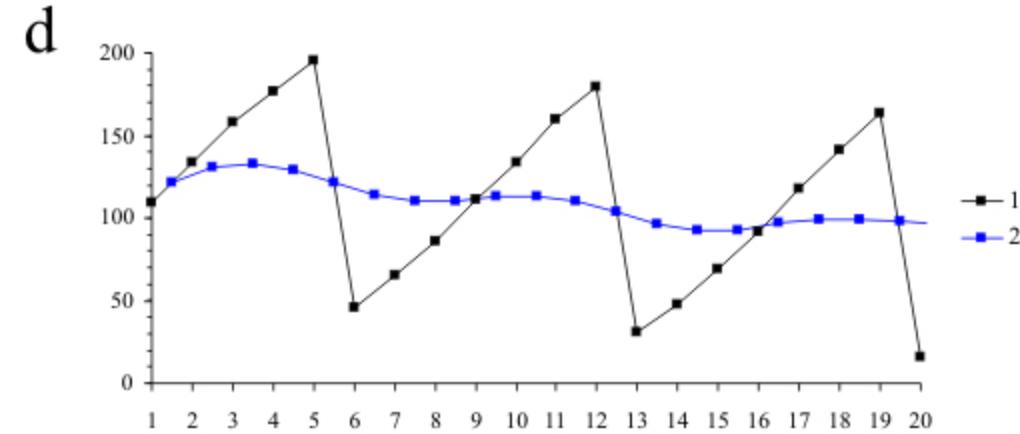
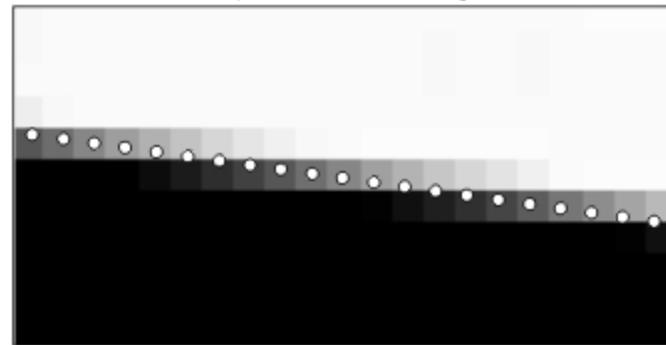
1 – resampling along any line direction in grey level

2 – smoothing sampled line with [1 3 6 8 9 9 8 6 3 1] / 54
minimized over counts of intercepts

Thresholded (binary) image

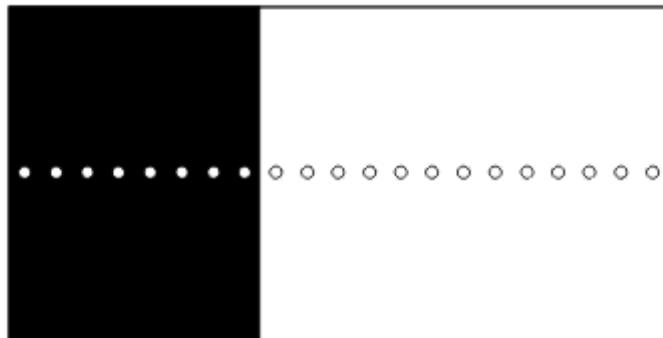


Grey level image

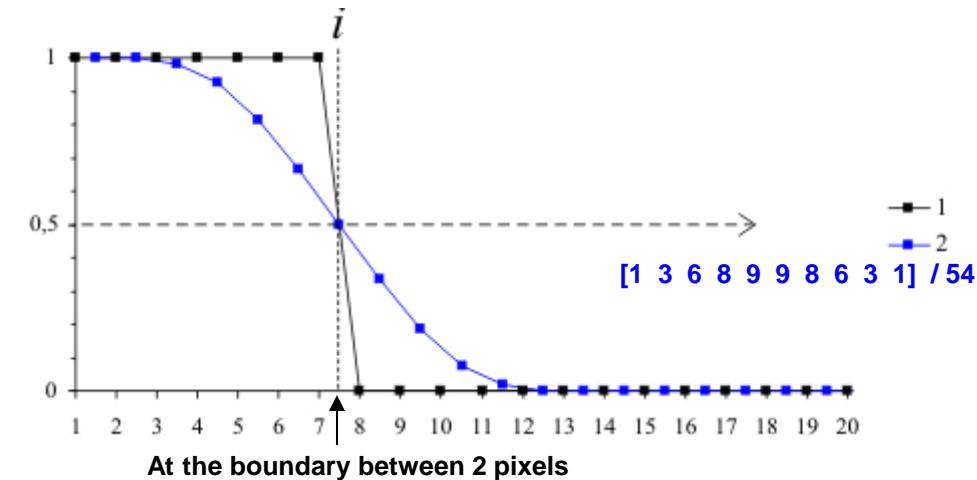


Intercept detection i in grey levels

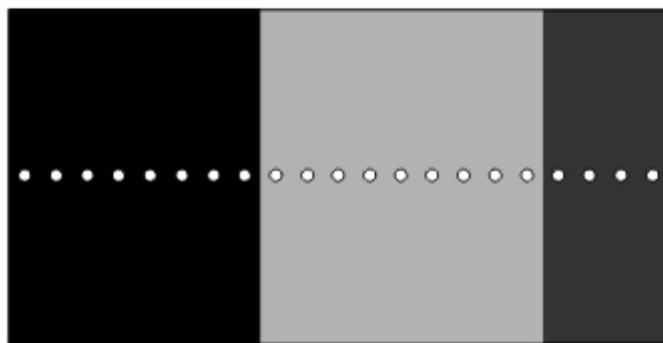
e



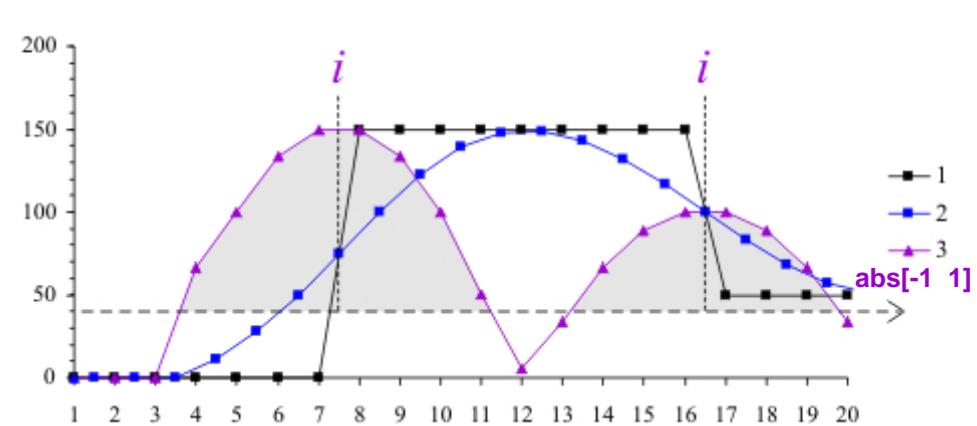
f



g

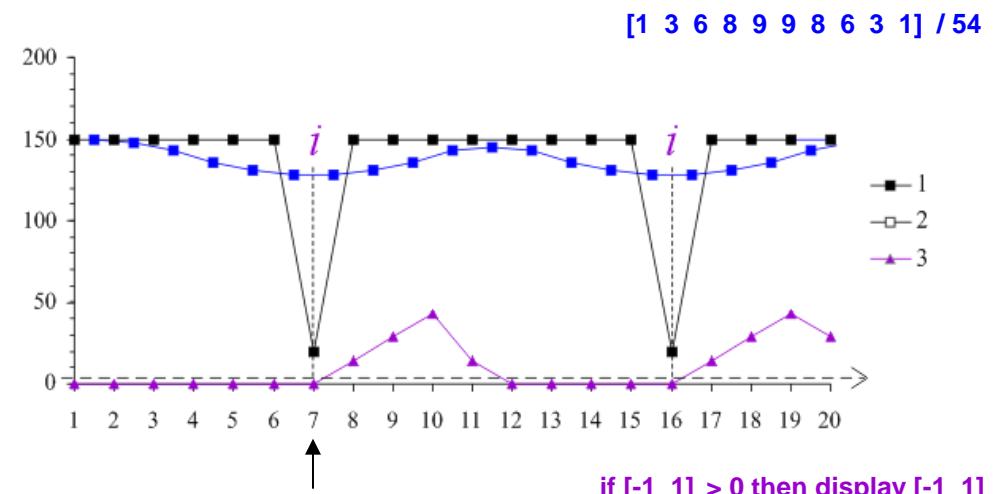
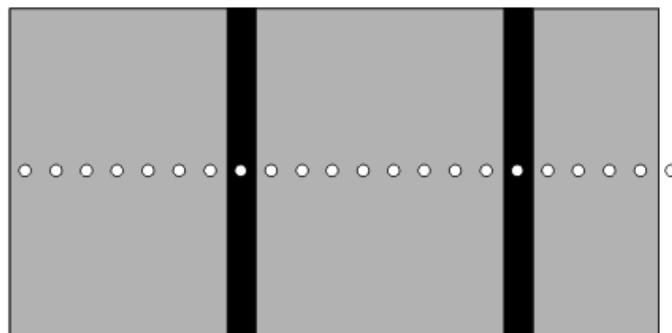


h



Intercept detection *i* in grey levels (sp. case)

In case of one grey phase with dark boundaries the intercepts detection occurs only from a dark pixel to a bright pixel to avoid double detection on both sides of each boundary

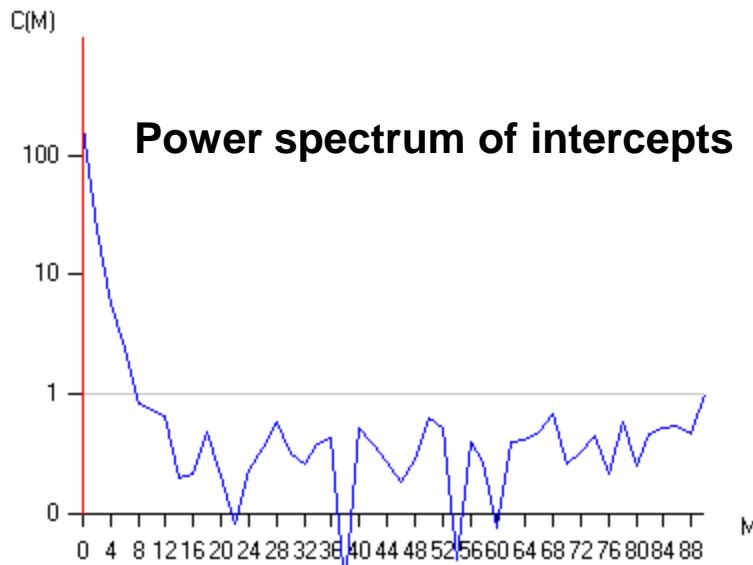


Fourier series of intercepts

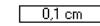
Hilliard, J. E., 1962. Specification and measurement of microstructural anisotropy. Trans. of the Metallurgical Society of AIME, 224: 1201-1211.

$$N_1(\alpha) = \sum_i N_1(j, \alpha)$$

$$\text{Fourier series of intercepts counts : } A_{2m} = \frac{2}{K} \sum_{k=0}^{K-1} N_L(k\delta\alpha) \cos 2mk\delta\alpha, \quad B_{2m} = \frac{2}{K} \sum_{k=0}^{K-1} N_L(k\delta\alpha) \sin 2mk\delta\alpha$$



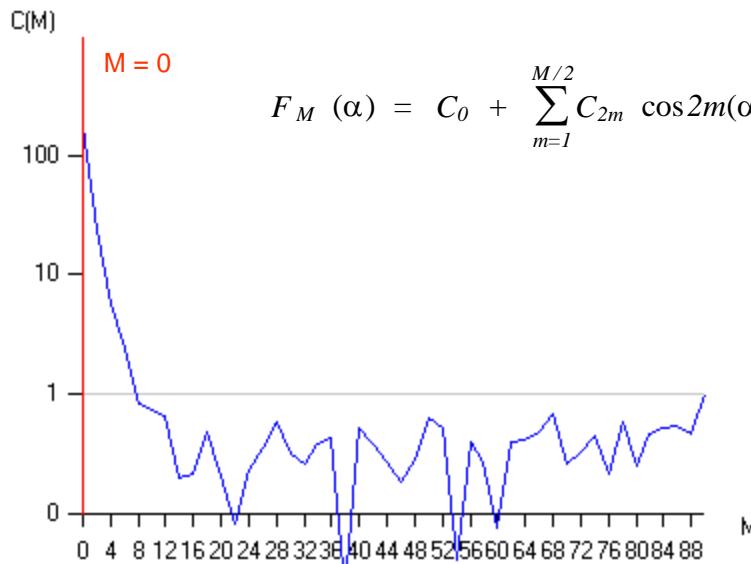
$$\tan 2m\varphi_{2m} = \frac{B_{2m}}{A_{2m}}, \quad C_{2m}^2 = A_{2m}^2 + B_{2m}^2$$



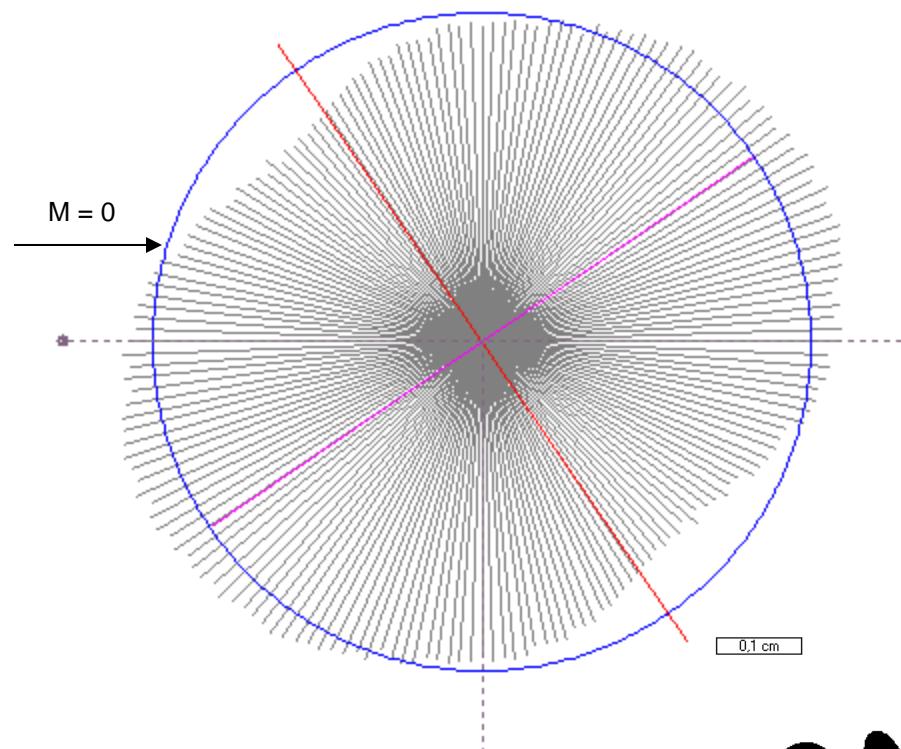
Fourier series : rose of intercepts

A a=0,2592 cm b=0,2592 cm R=1,000 , 145,62° , angle X: 124,38°

$$N_1(\alpha) = \sum_j N_1(j, \alpha)$$



$$F_M(\alpha) = C_0 + \sum_{m=1}^{M/2} C_{2m} \cos 2m(\alpha - \varphi_{2m})$$



$$F_M(\alpha) = C_0 + \sum_{m=1}^{M/2} (A_{2m} \cos 2m\alpha + B_{2m} \sin 2m\alpha)$$

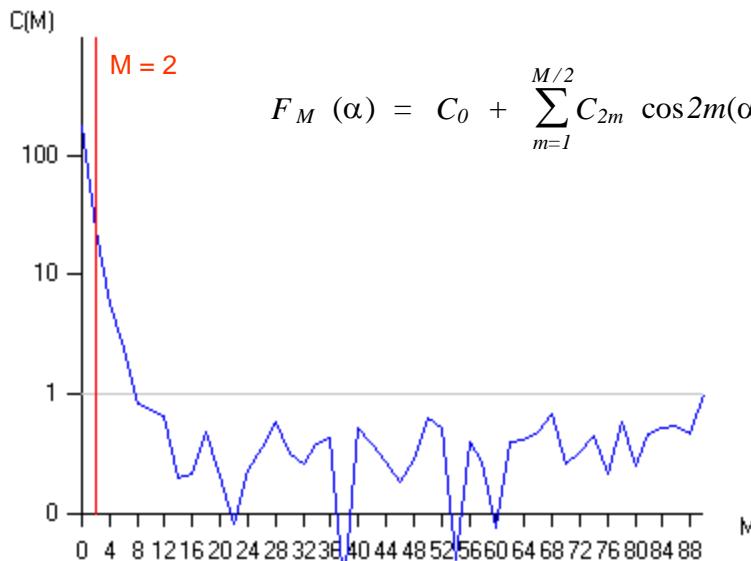


Fourier series : rose of intercepts

A a=0,2942 cm b=0,2242 cm R=1,312 , 145,62° , angle X: 124,38°

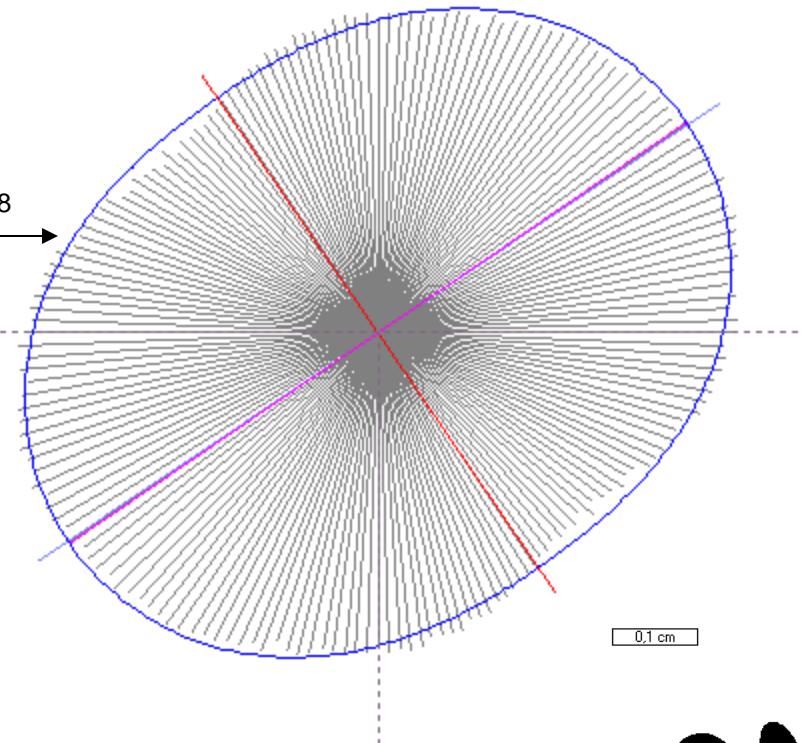
(1) 56°

$$N_1(\alpha) = \sum_j N_1(j, \alpha)$$



$$F_M(\alpha) = C_0 + \sum_{m=1}^{M/2} C_{2m} \cos 2m(\alpha - \varphi_{2m})$$

$\rightarrow M = 8$



$$F_M(\alpha) = C_0 + \sum_{m=1}^{M/2} (A_{2m} \cos 2m\alpha + B_{2m} \sin 2m\alpha)$$

0,1 cm

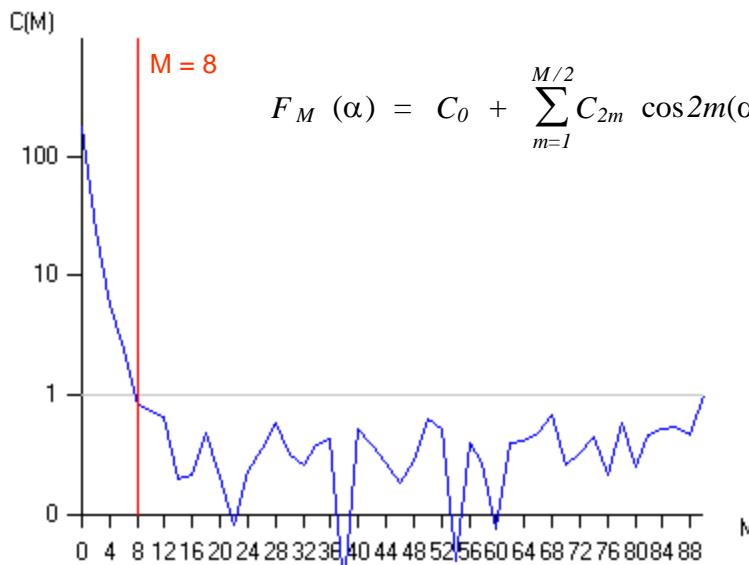


Fourier series : rose of intercepts

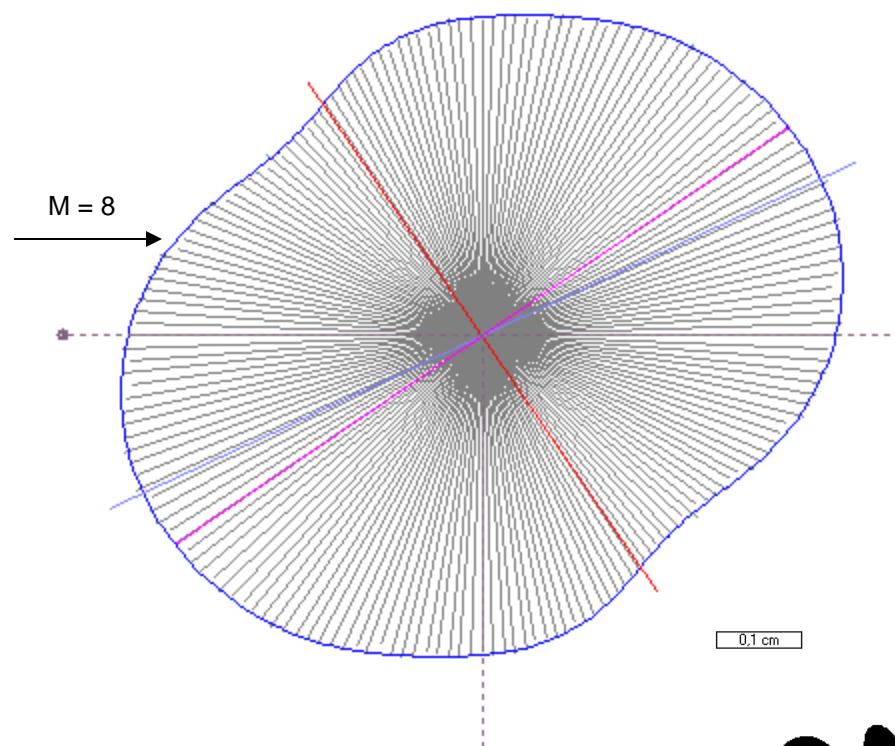
A a=0,2924 cm b=0,2212 cm R=1,322 , 145,62° , angle X: 124,38°

(1) 65°

$$N_1(\alpha) = \sum_j N_1(j, \alpha)$$



$$F_M(\alpha) = C_0 + \sum_{m=1}^{M/2} C_{2m} \cos 2m(\alpha - \varphi_{2m})$$



$$F_M(\alpha) = C_0 + \sum_{m=1}^{M/2} (A_{2m} \cos 2m\alpha + B_{2m} \sin 2m\alpha)$$



Fourier series : rose of traverses (mean length)

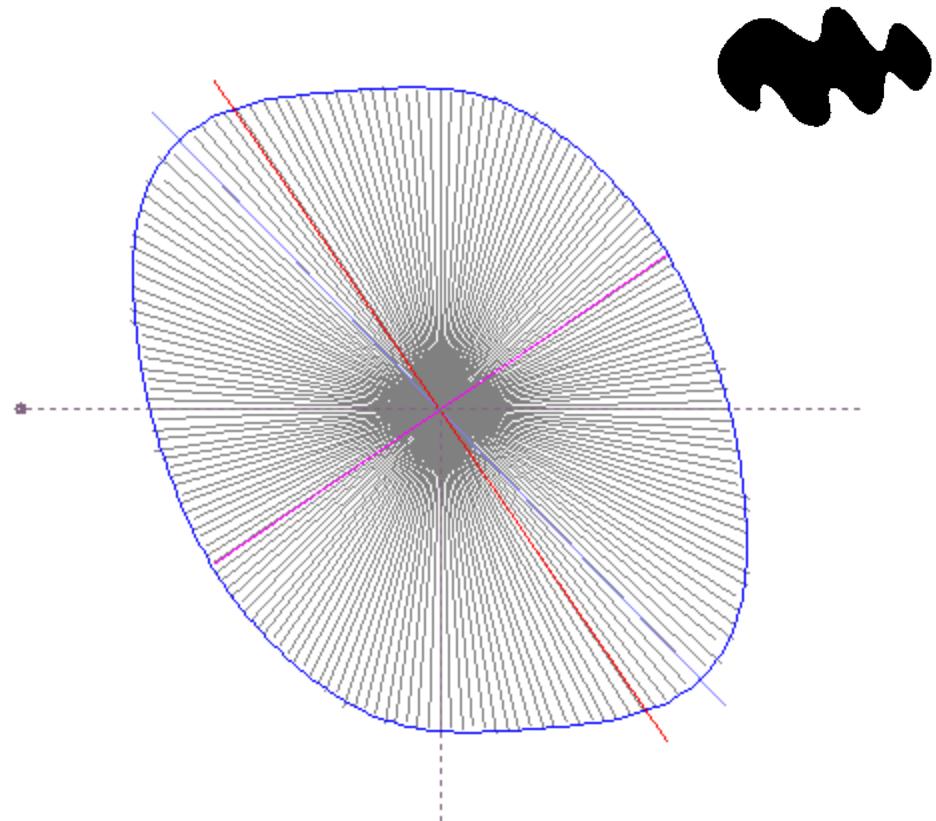
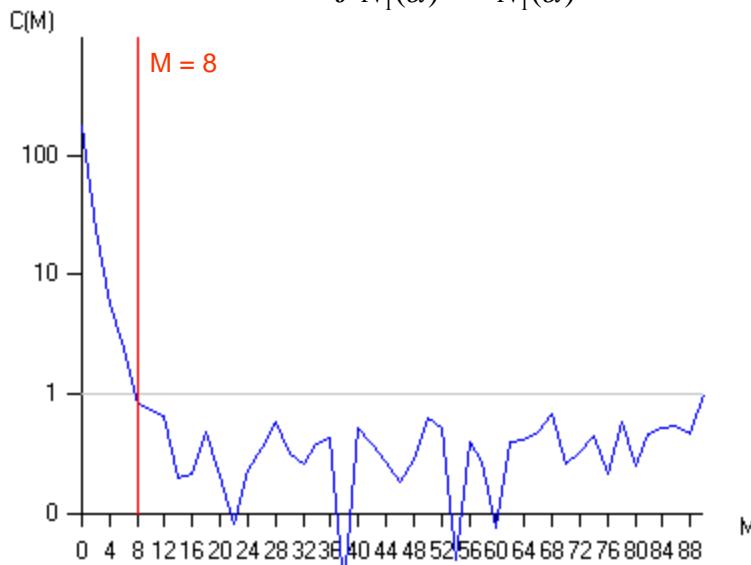
A a=0,1171 cm b=0,0886 cm R=1,322 , 145,62° , angle X: 124,38°

0,1 cm

(1) 136°

$$N_1(\alpha) = \sum_j N_1(j, \alpha)$$

$$\bar{L}(\alpha) = \frac{A}{J N_1(\alpha)} = I \frac{N_0}{N_1(\alpha)}$$



$$F_M (\alpha) = C_0 + \sum_{m=1}^{M/2} (A_{2m} \cos 2m\alpha + B_{2m} \sin 2m\alpha)$$

0,1 cm

Fourier series : rose of directions

For one rod :

$$D(\alpha) = L / \sin(\alpha - \psi) = \frac{L}{\sqrt{2}} \sqrt{1 - \cos 2(\alpha - \psi)}$$

$$\begin{aligned} D(\alpha) + D''(\alpha) &= 0 \quad \text{for } \alpha \neq \psi \text{ modulo } \pi \\ &\infty \quad \text{for } \alpha = \psi \text{ modulo } \pi \end{aligned}$$

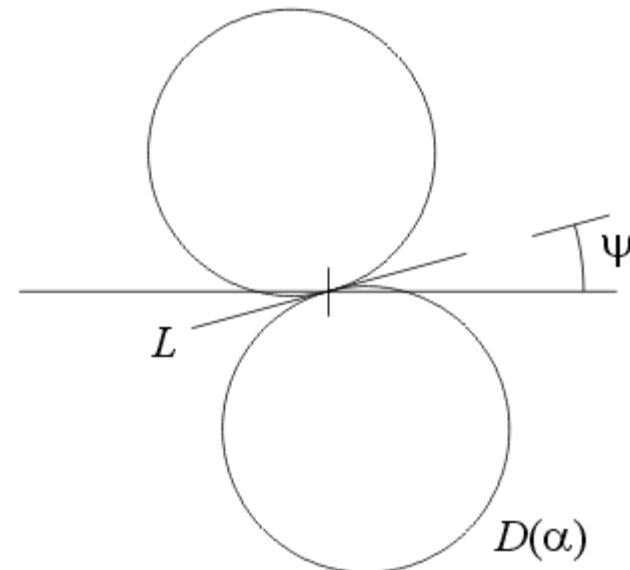
Since the first derivative $D'(\alpha)$ changes from $-L$ to $+L$ on either side of the singularity, $D(\alpha) + D''(\alpha)$ for one rod is a Dirac ‘function’, $2 L \delta(\alpha - \psi)$, i.e. with an integrated value of $2 L$:

$$D(\alpha) + D''(\alpha) = 2 L \delta(\alpha - \psi)$$

For a population of rods :

$$N_L(\alpha) = \frac{D(\alpha)}{A^w} = \frac{1}{A^w} \sum_{i=1}^P L^i / \sin(\alpha - \psi^i) /$$

$$N_L(\alpha) + N_L''(\alpha) = \frac{2}{A^w} \sum_{i=1}^P L^i \delta(\alpha - \psi^i)$$



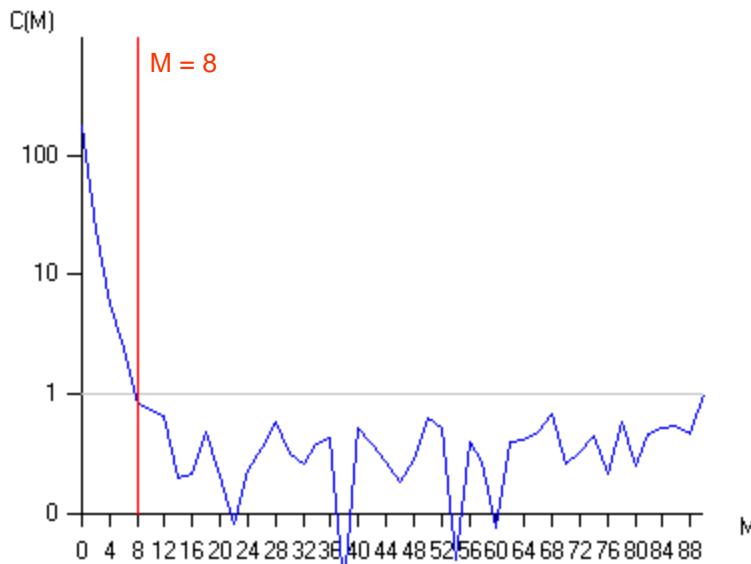
Hilliard, J. E., 1962. Specification and measurement of microstructural anisotropy. Trans. of the Metallurgical Society of AIME, 224: 1201-1211.

Fourier series : rose of directions

For a population of rods :

$$N_L(\alpha) = \frac{D(\alpha)}{A^w} = \frac{I}{A^w} \sum_{i=1}^P L^i / \sin(\alpha - \psi^i) /$$

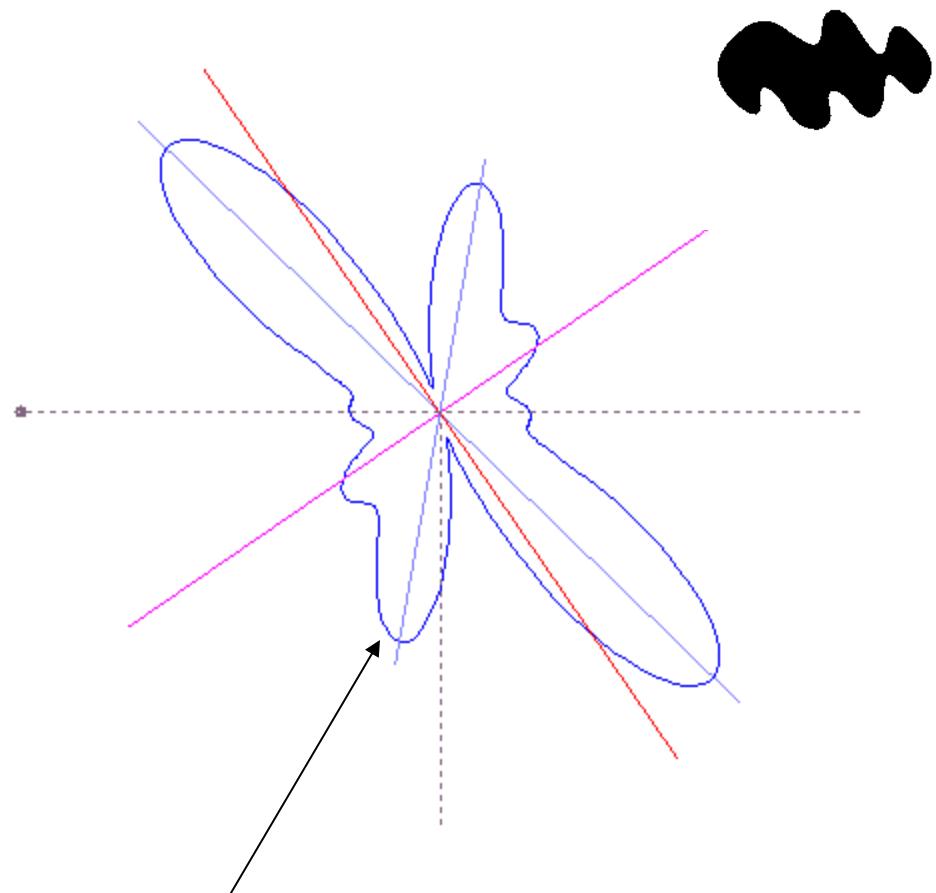
$$N_L(\alpha) + N_L''(\alpha) = \frac{2}{A^w} \sum_{i=1}^P L^i \delta(\alpha - \psi^i)$$



A Nla=2,069 cm⁻¹ Nlb=2,735 cm⁻¹ R=1,322 , 145,62° , angle X: 124,38°

0.1 cm

(1) 134°
(2) 10°



$$N_L''(\alpha) \approx F_M''(\alpha) = - \sum_{m=1}^{M/2} (2m)^2 (A_{2m} \cos 2m\alpha + B_{2m} \sin 2m\alpha)$$

$$\frac{2l(\alpha)}{A^w} \approx F_M(\alpha) + F_M''(\alpha) = C_0 + \sum_{m=1}^{M/2} [1 - (2m)^2] (A_{2m} \cos 2m\alpha + B_{2m} \sin 2m\alpha)$$

Fourier series : characteristic shape

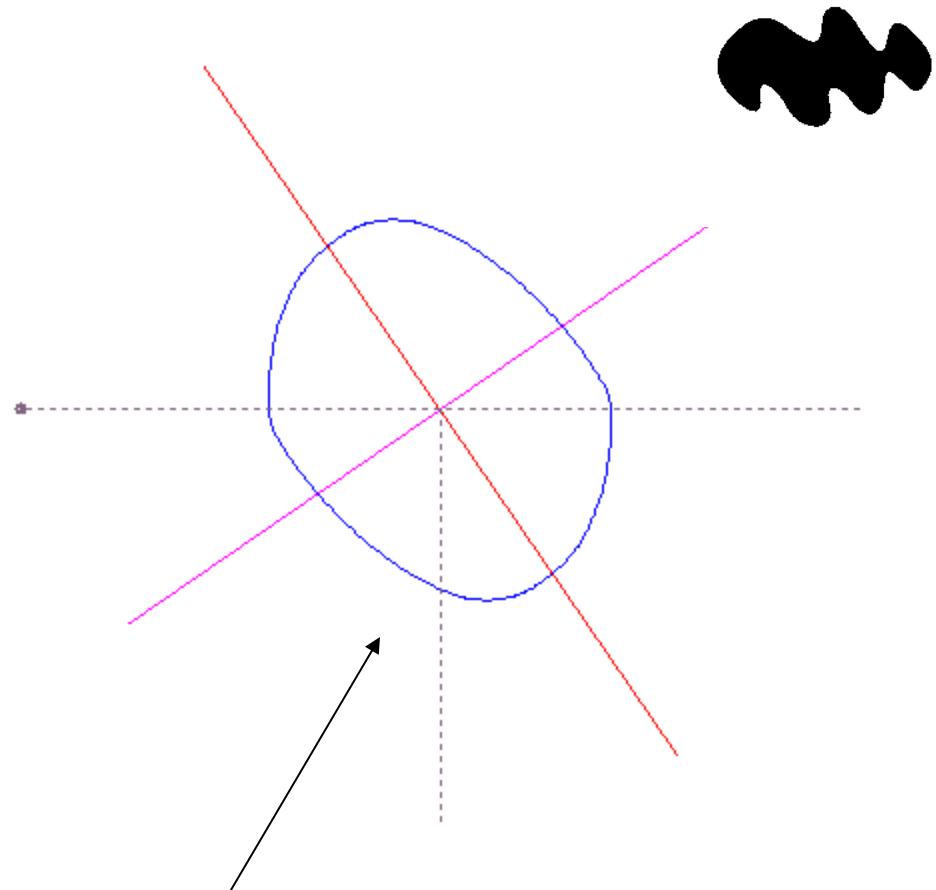
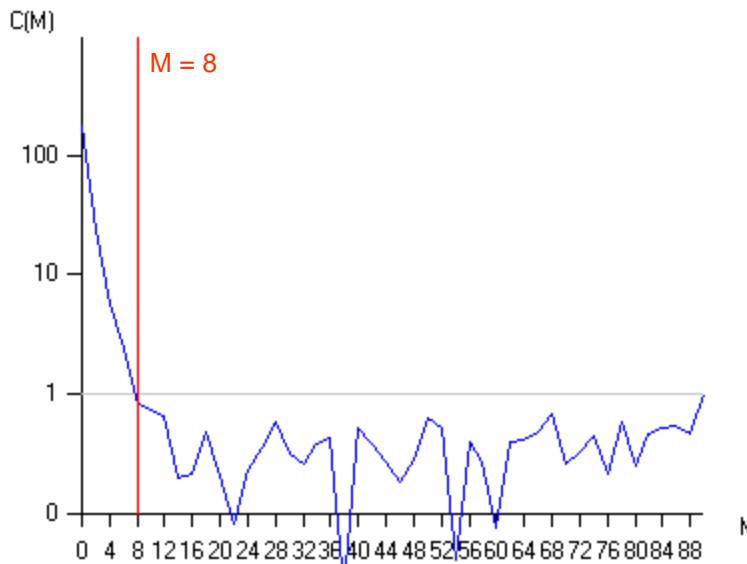
A R=1,322, 145,62°, angle X: 124,38°

0,1 cm

For a population of rods :

$$N_L(\alpha) = \frac{D(\alpha)}{A^w} = \frac{I}{A^w} \sum_{i=1}^P L^i / \sin(\alpha - \psi^i) /$$

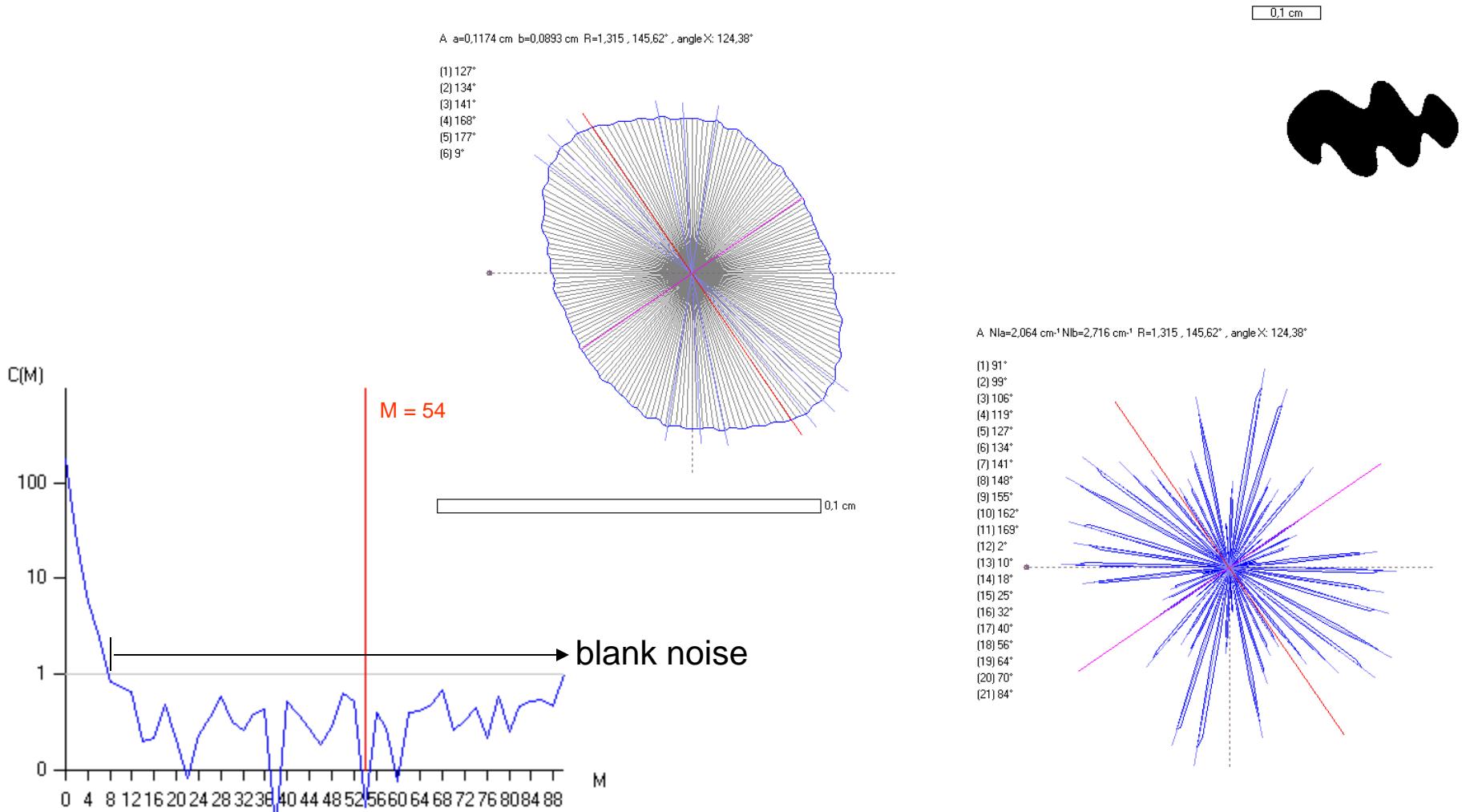
$$N_L(\alpha) + N_L''(\alpha) = \frac{2}{A^w} \sum_{i=1}^P L^i \delta(\alpha - \psi^i)$$



$$N_L''(\alpha) \approx F_M''(\alpha) = - \sum_{m=1}^{M/2} (2m)^2 (A_{2m} \cos 2m\alpha + B_{2m} \sin 2m\alpha)$$

$$\frac{2l(\alpha)}{A^w} \approx F_M(\alpha) + F_M''(\alpha) = C_0 + \sum_{m=1}^{M/2} [1 - (2m)^2] (A_{2m} \cos 2m\alpha + B_{2m} \sin 2m\alpha)$$

Fourier series : noise

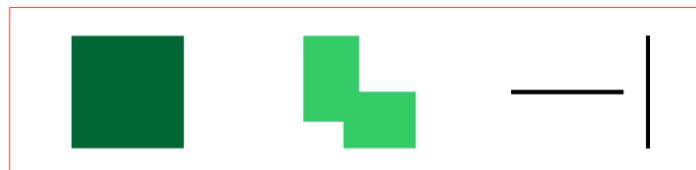


Case of 1 object

Symmetry analysis with the Fourier power spectrum

Theoretical shapes of roses

Shapes



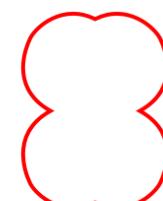
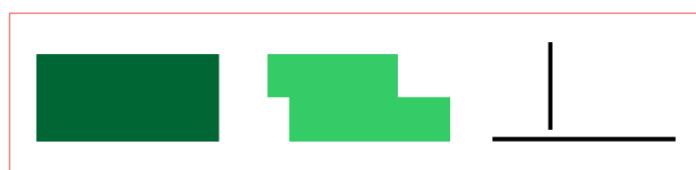
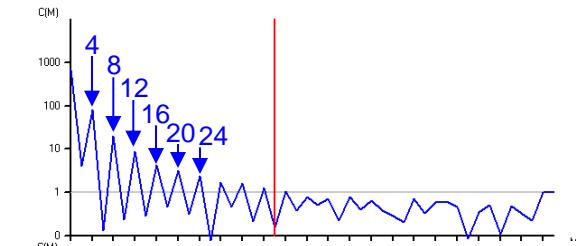
Intercept counts



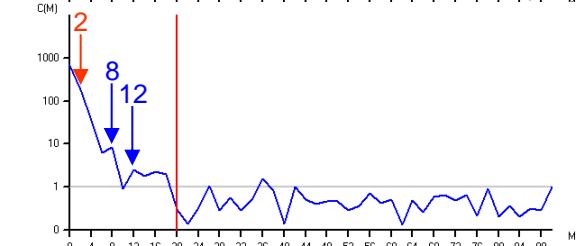
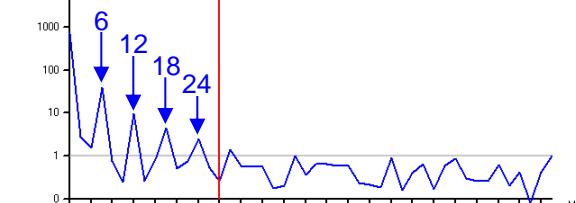
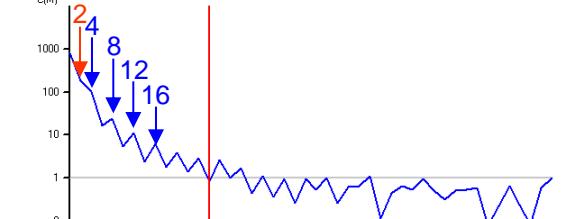
Mean intercept lengths



C_2 = anisotropy
 C_M = symmetry



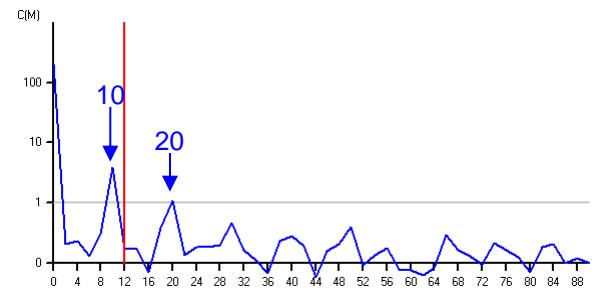
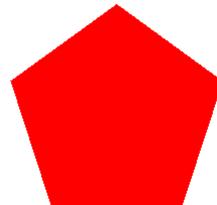
shapes without scale



Case of 1 object

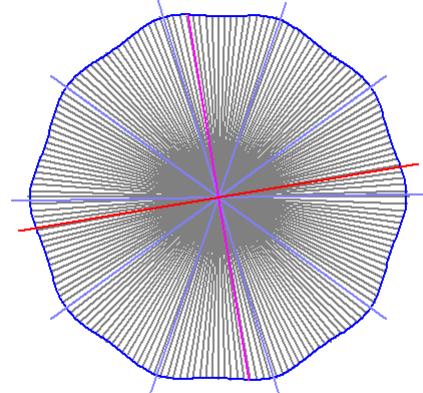
Symmetry analysis with the Fourier power spectrum

C_2 = anisotropy
 C_M = symmetry



A. $a=0.5060 \text{ cm}$ $b=0.5072 \text{ cm}$ $R=0.998 , 80.42^\circ$

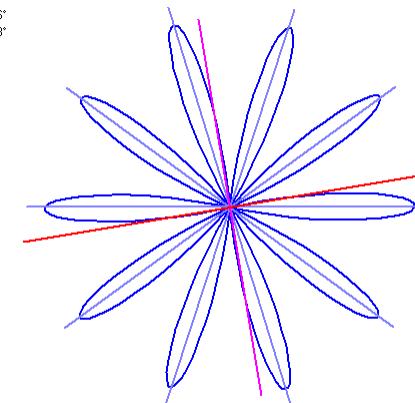
- (1) 126°
- (2) 163°
- (3) 19°
- (4) 54°
- (5) 89°



0.1 cm

A. $Nla=0.138 \text{ cm}^{-1}$ $Nlb=0.137 \text{ cm}^{-1}$ $R=0.998 , 80.42^\circ$

- (1) 90°
- (2) 126°
- (3) 163°
- (4) 18°
- (5) 54°

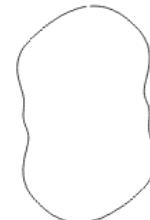


Shape analysis with boundary directions

Shapes

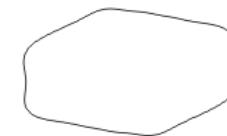


intercept
counts



Roses

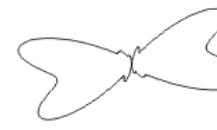
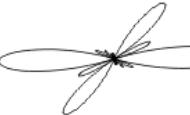
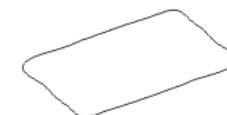
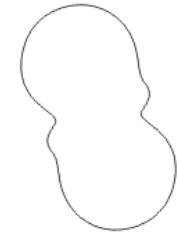
mean
lengths



boundary
directions



characteristic
shapes

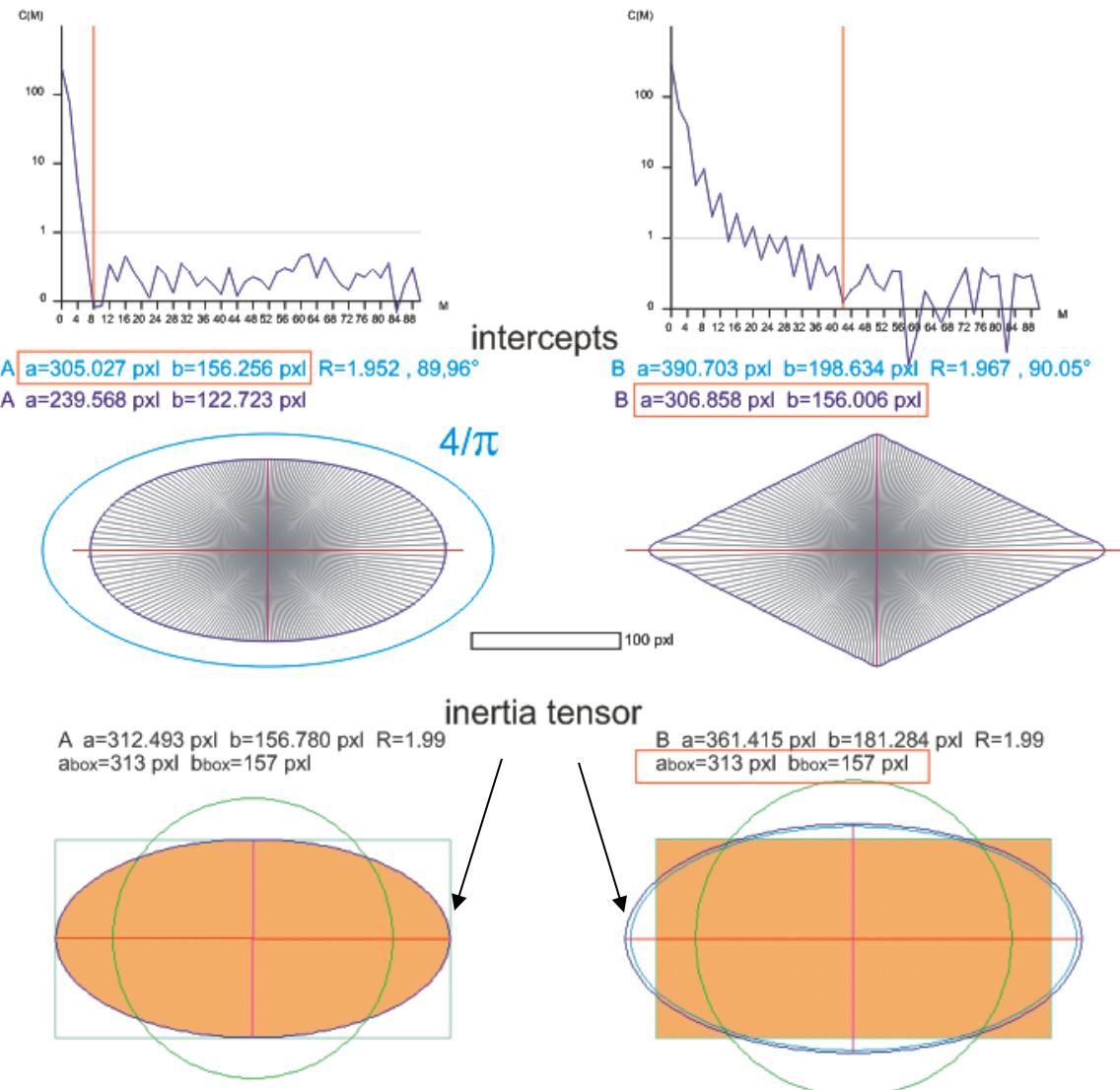


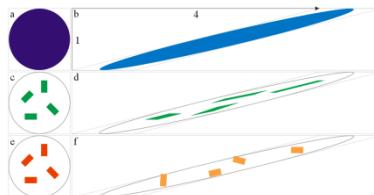
4/ π size correction

This is the ratio between the mean intercept length and the long axis of an ellipse

Therefore the correction is :

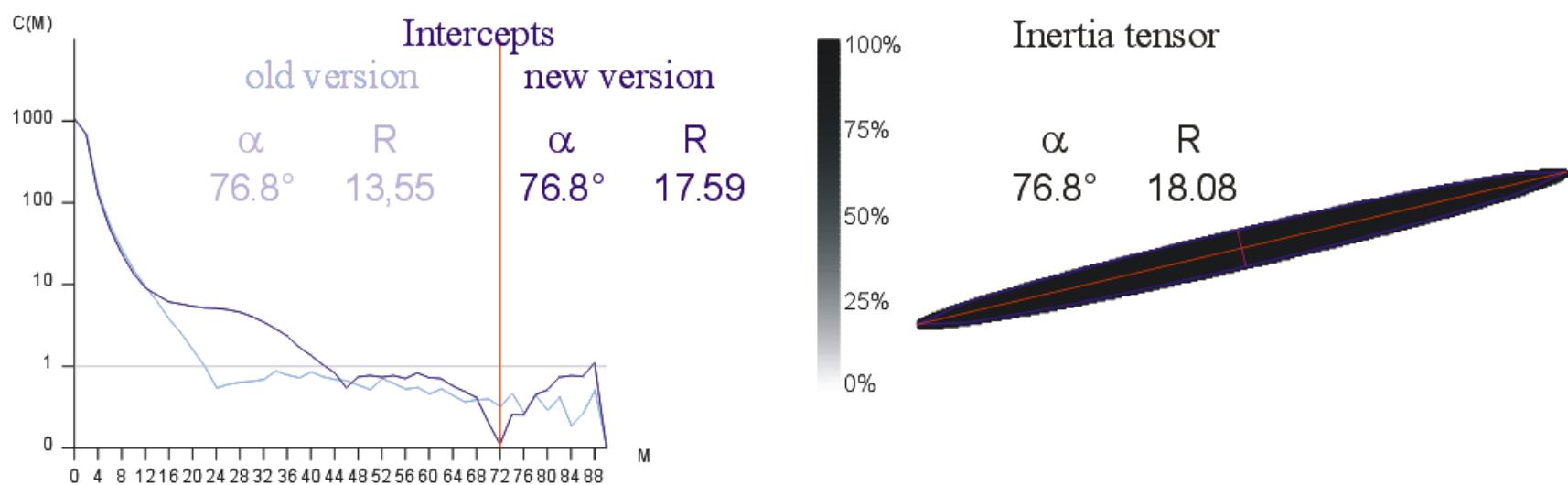
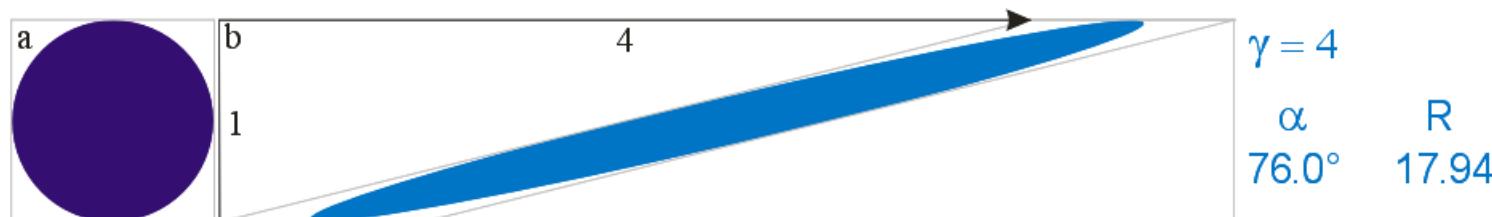
- valid only for ellipses and population of objects giving an elliptical mean shape
- not valid on rectangles parallel to each other

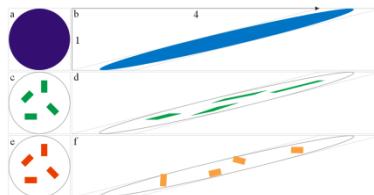




Passive / active deformation of object population

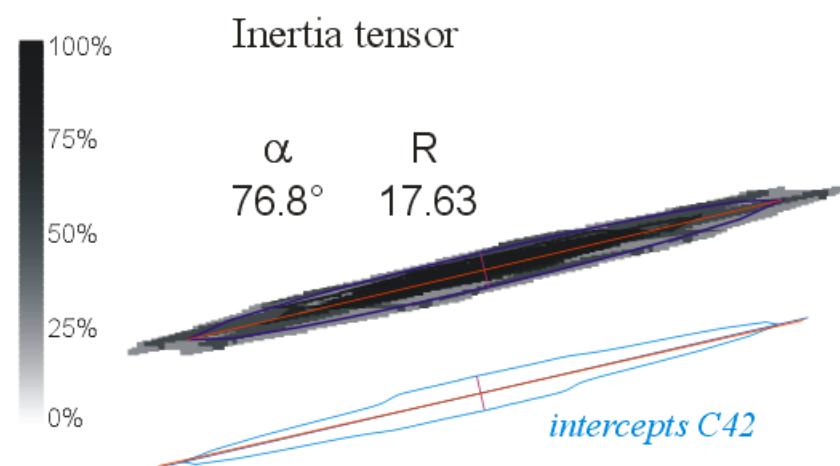
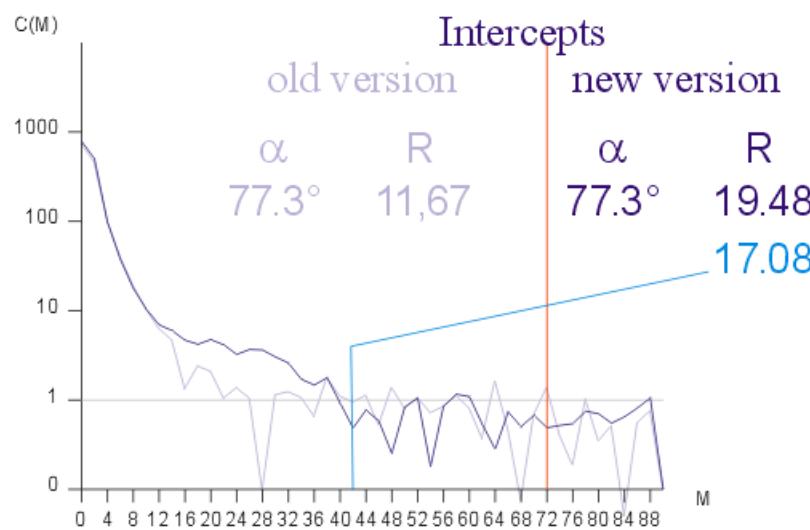
1 passive object

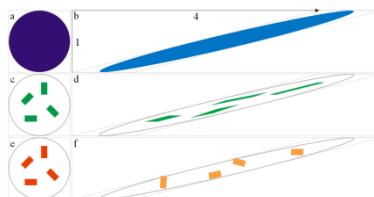




Passive / active deformation of object population

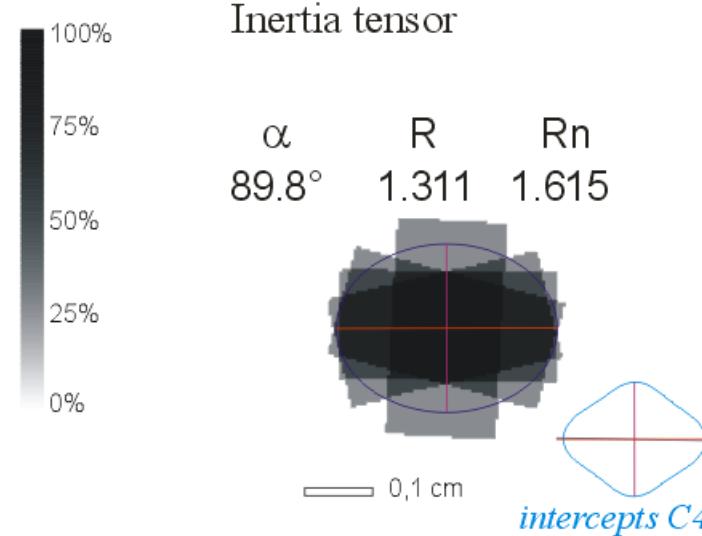
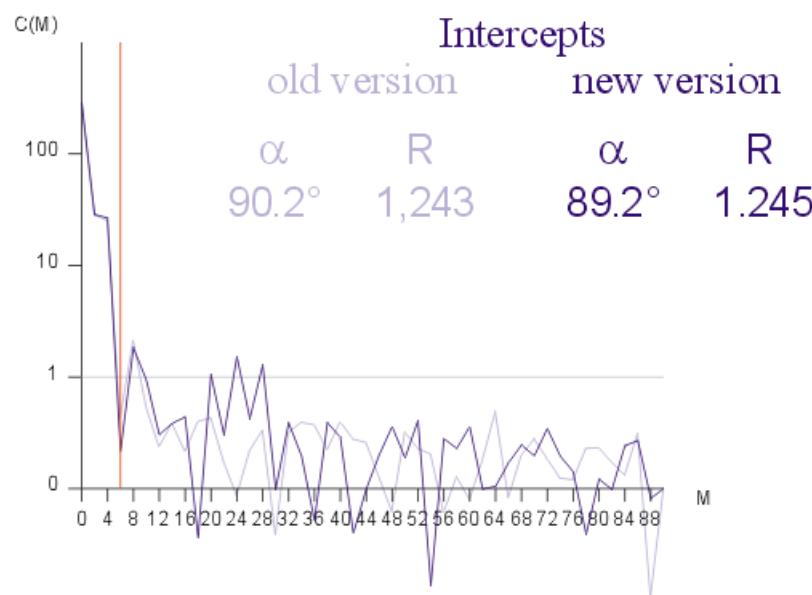
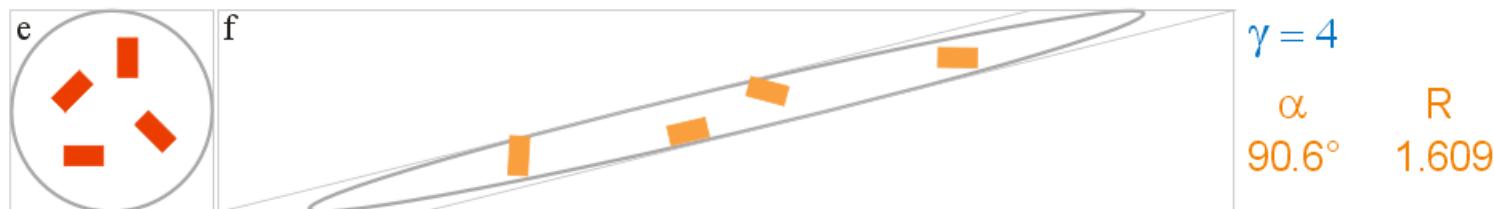
4 passive rectangles





Passive / active deformation of object population

4 active rectangles



Passive deformation of Spirifers

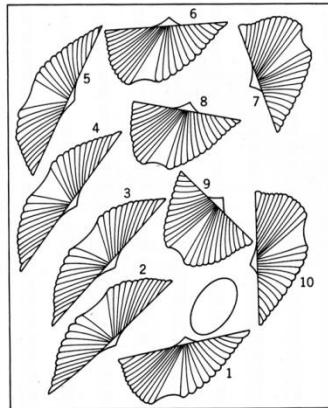


Fig. 5.6. Slab of deformed fossil shells.

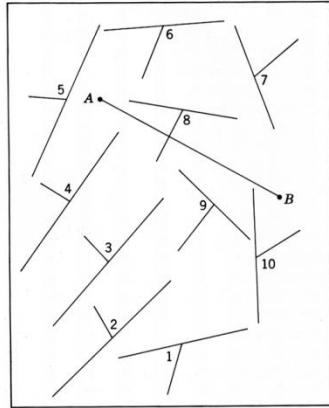


Fig. 5.8. Preliminary steps in construction of strain ellipse by Wellman's method.

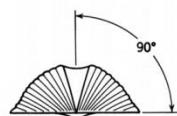


Fig. 5.7. Form of individual shell in undeformed state.

1.62

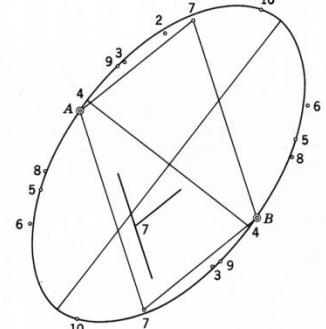
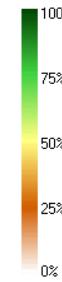
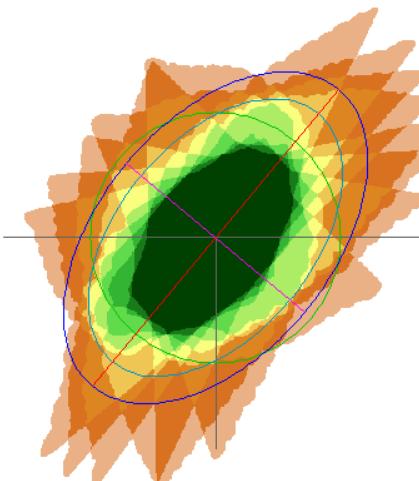


Fig. 5.9. Strain ellipse.

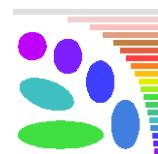
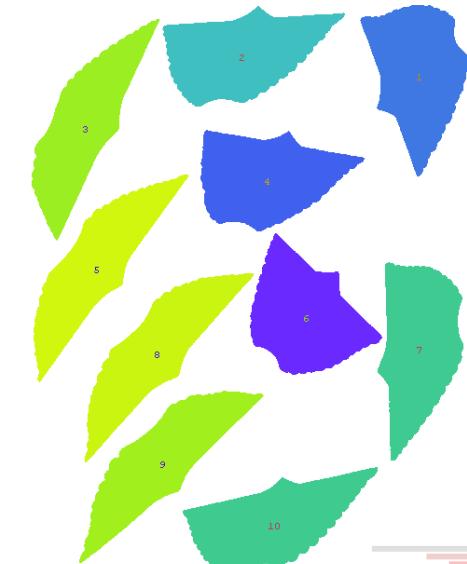
A=10 a=3.9526 cm b=2.3792 cm R=1.661 [2.105]n [2.120]b , 39.73° , angle X : 129.73°
K=0.468, Kn=0.632 (0.741), Kbn=0.636 (0.736)



□ 0,1 cm



1.66

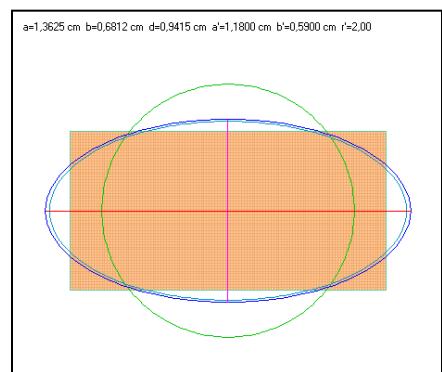


Window of measurement

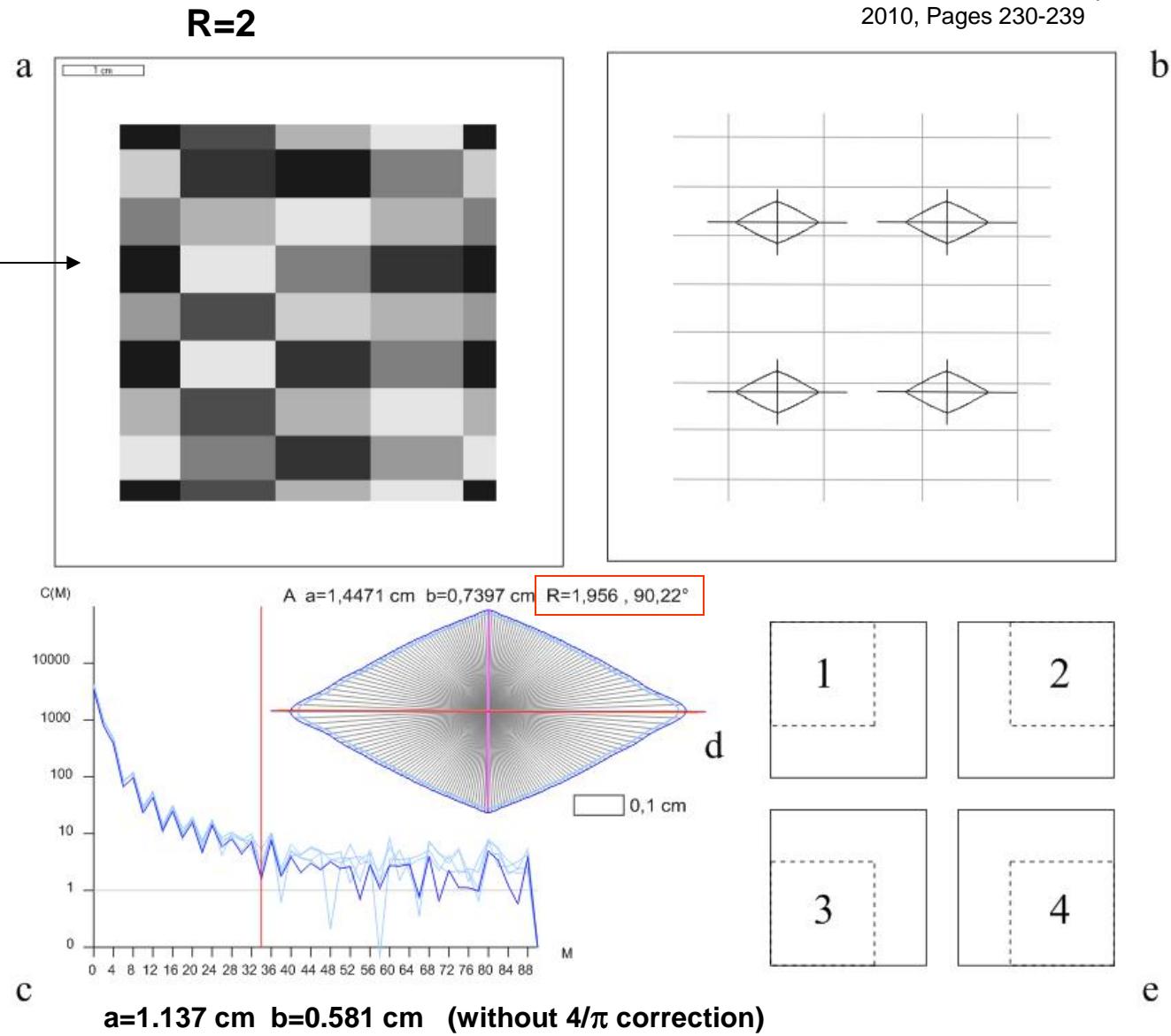
with mask

Means no intercept
detection along grey
levels of the mask

$a_{\text{box}}=1,18\text{cm}$ $b_{\text{box}}=0,69\text{cm}$



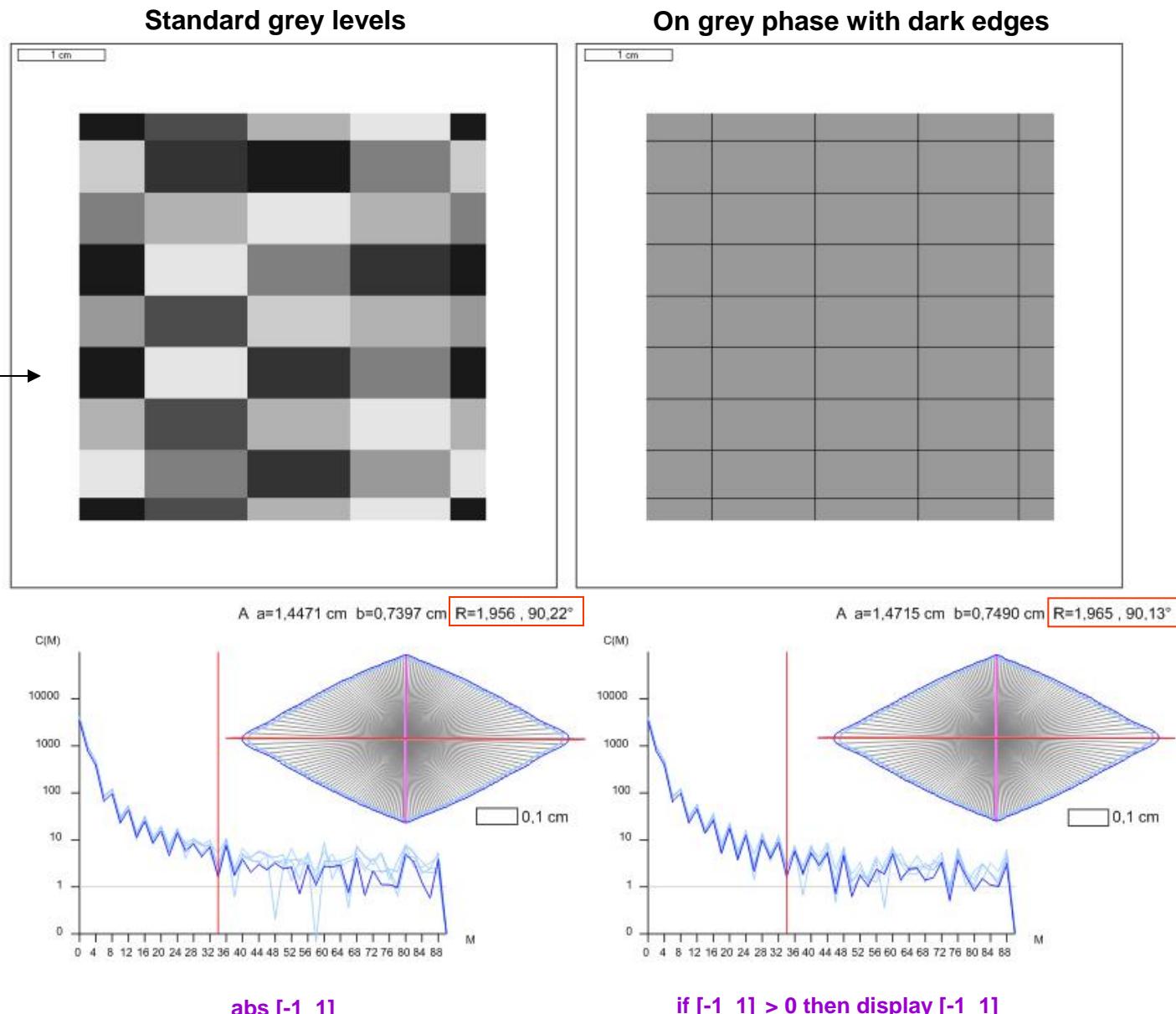
Inertia tensor of one rectangle



Window of measurement

with mask

Means no intercept detection along grey levels of the mask



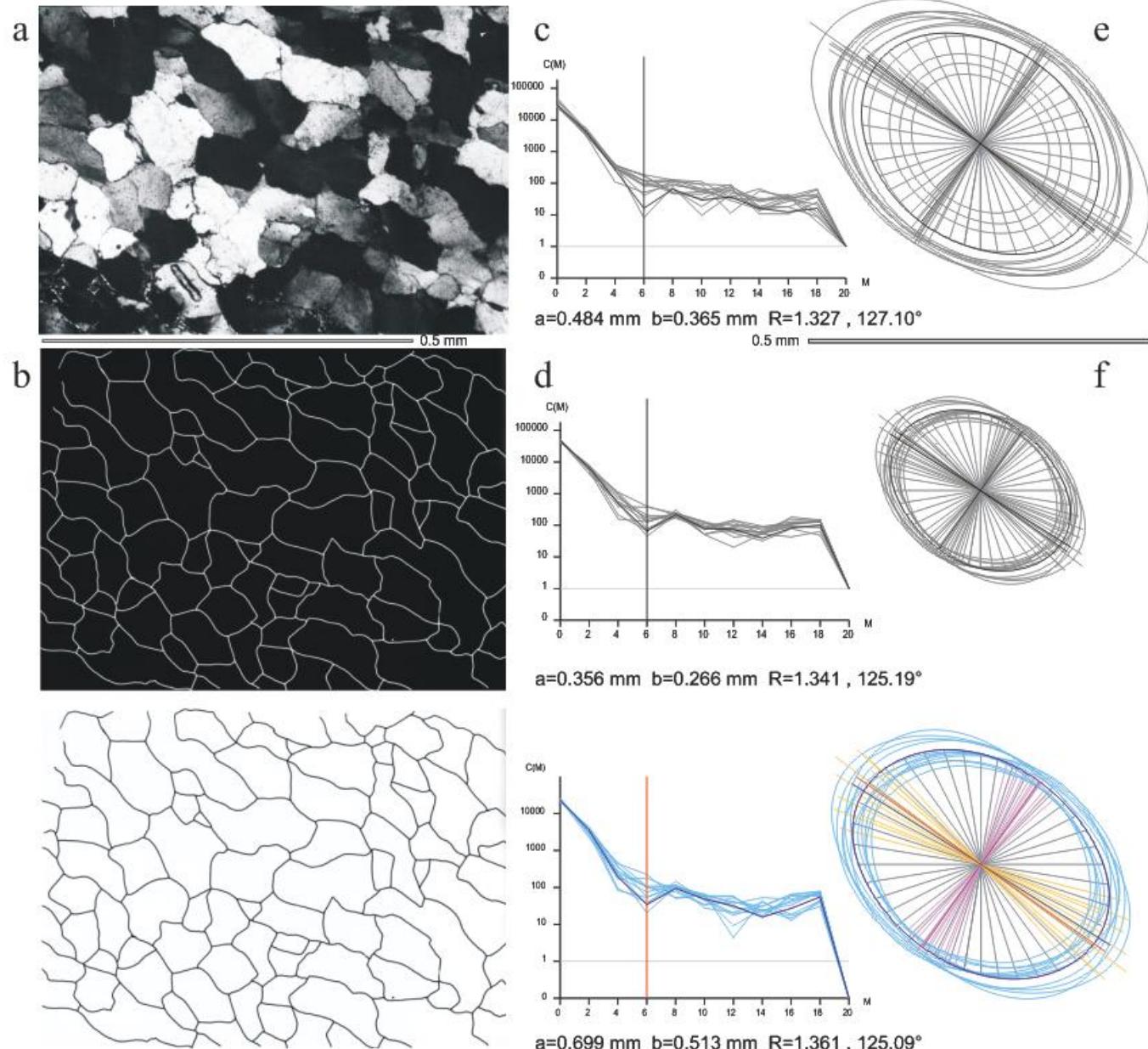
Launeau P., Archanjo C. J., Picard D., Arbaret L., Robin P.Y. (2010). Two- and three-dimensional shape fabric analysis by the intercept method in grey levels. Tectonophysics, Volume 492, Issues 1-4, 20 September 2010, Pages 230-239

Direct analysis versus drawing

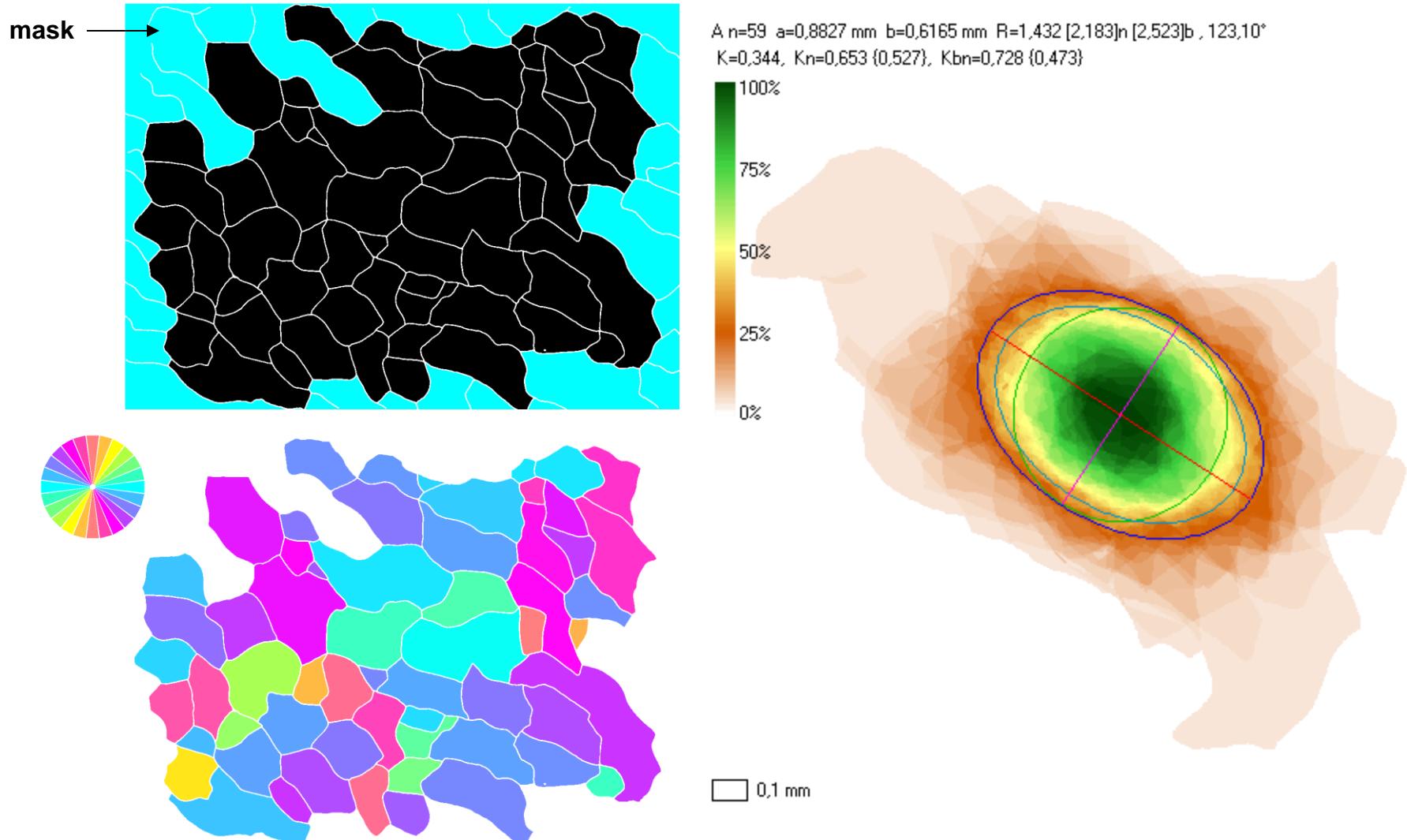
Launeau P., Archanjo C. J.,
Picard D., Arbaret L., Robin
P.Y. (2010). Two- and three-
dimensional shape fabric analysis
by the intercept method in grey
levels. Tectonophysics, Volume
492, Issues 1-4, 20 September
2010, Pages 230-239

abs [-1 1]

Standard grey levels

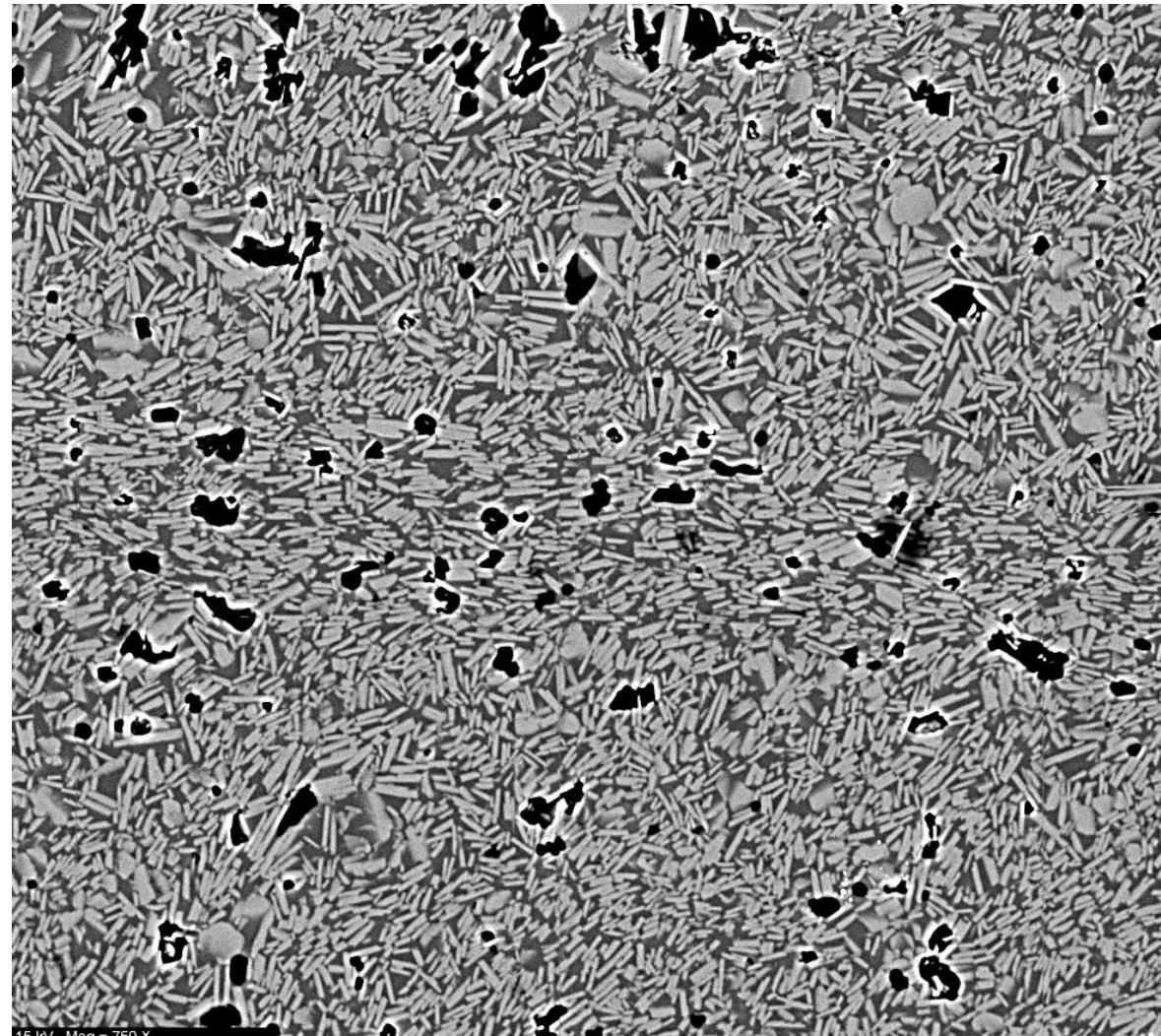


Direct analysis versus drawing with inertia tensor



Application to the BSE image of a synthetic magma

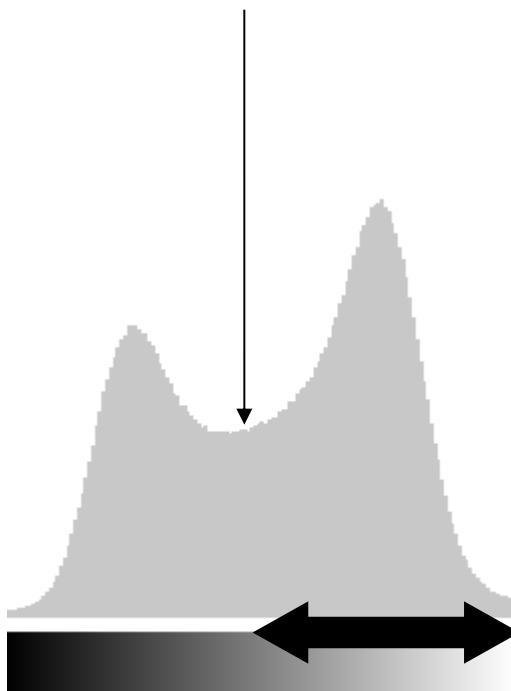
Plagioclase-bearing suspension composed of 52% of crystals was synthesized and then deformed using a Paterson HP-HT apparatus at a confining pressure of 300 MPa, a temperature of 850 C and a shear strain $\gamma = 3.5$



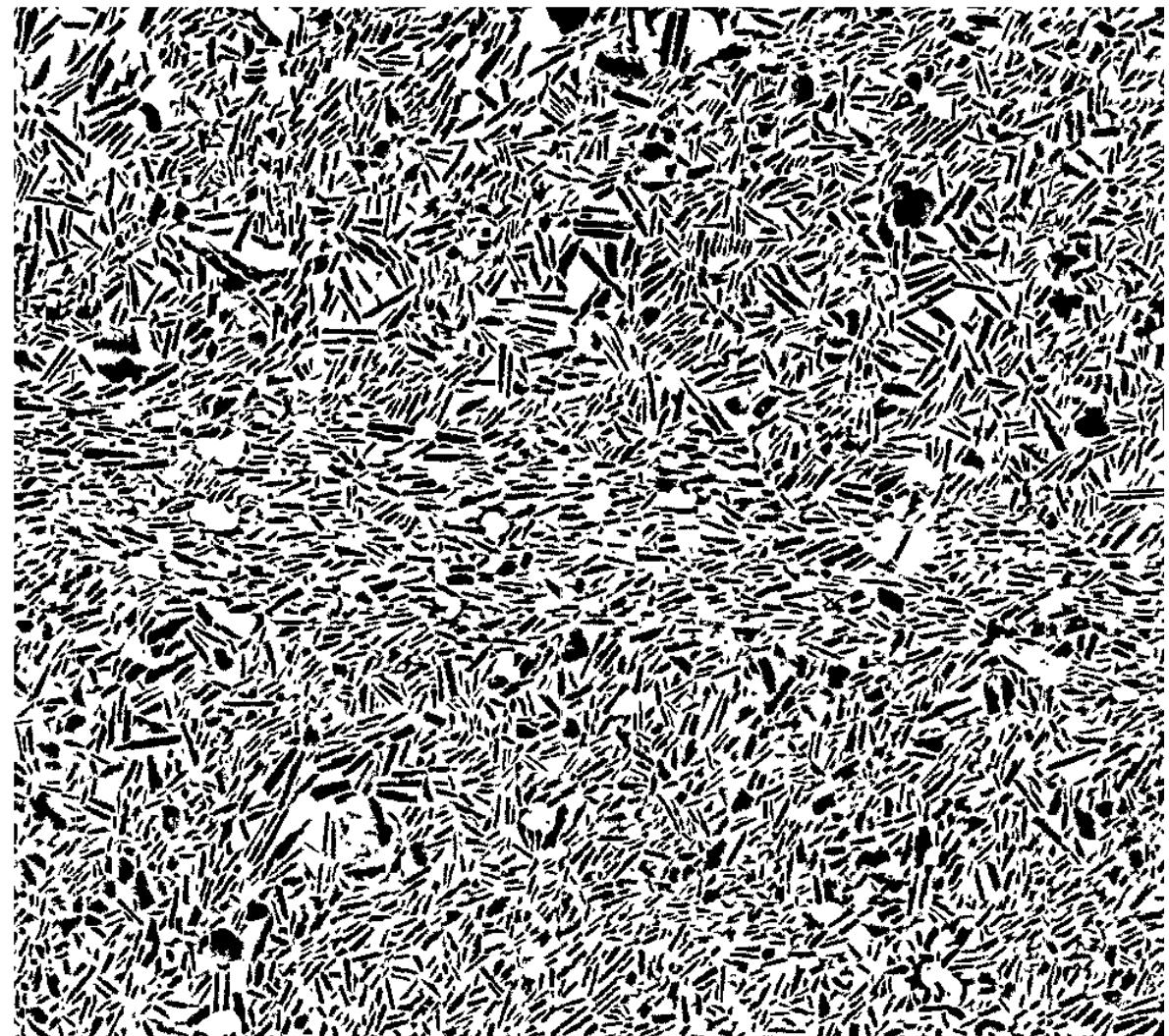
Application to the BSE image of a synthetic magma

Threshold at grey level

140

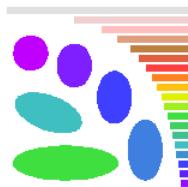


+ drawing a lot of
missing boundaries

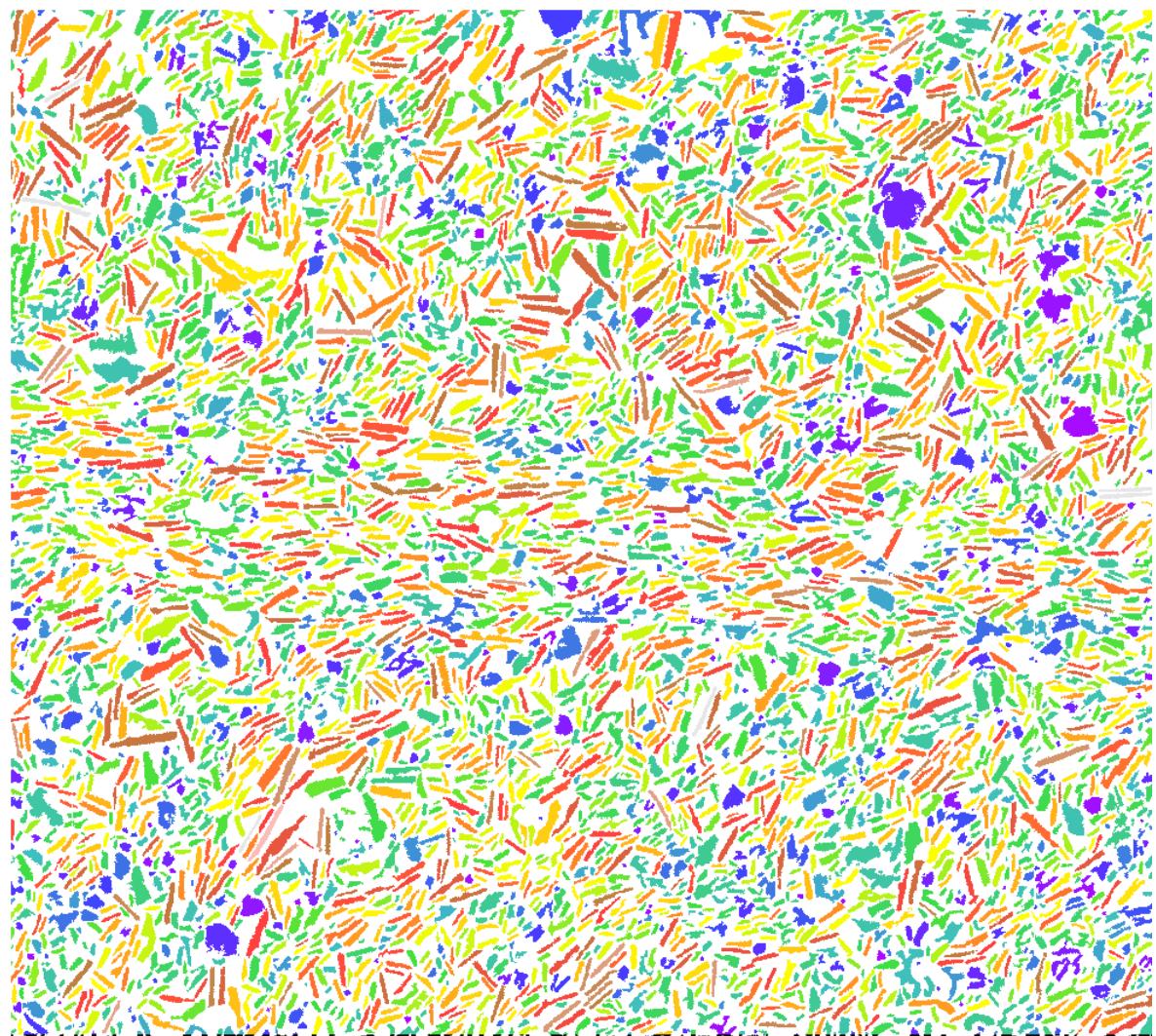
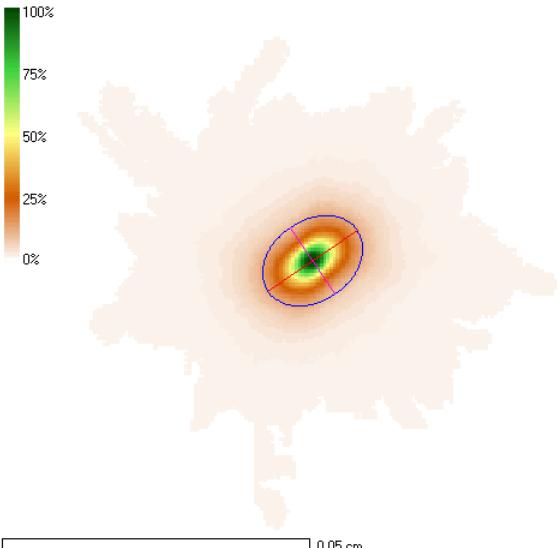


Application to the BSE image of a synthetic magma

Inertia tensor : R

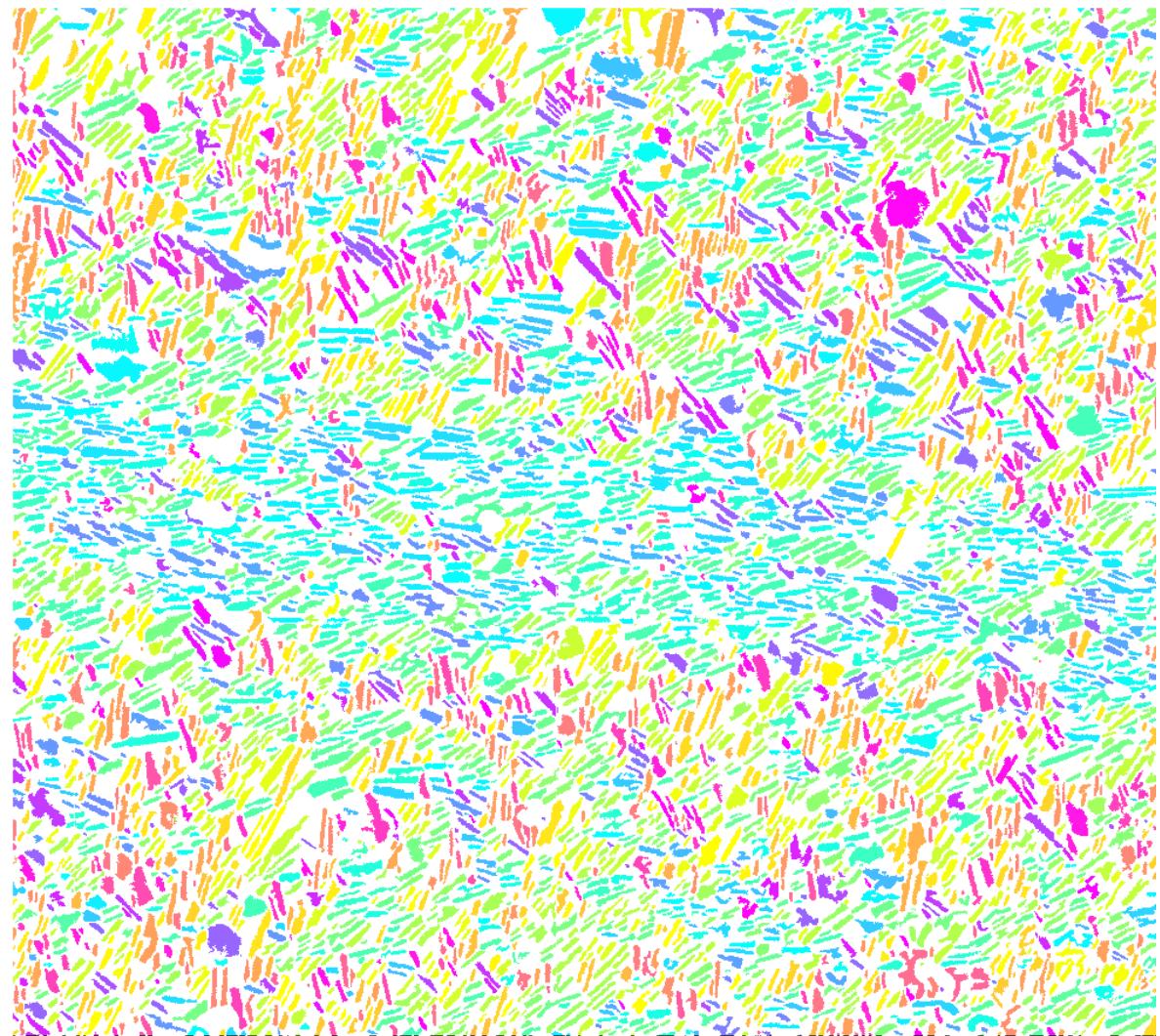


$A = 4973$ $a = 0.0176 \text{ cm}$ $b = 0.0131 \text{ cm}$ $R = 1.346 [1.419]n [1.464]b , 55.51^\circ$
 $K = 0.288$, $Kn = 0.336 (0.858)$, $Kbn = 0.364 (0.793)$

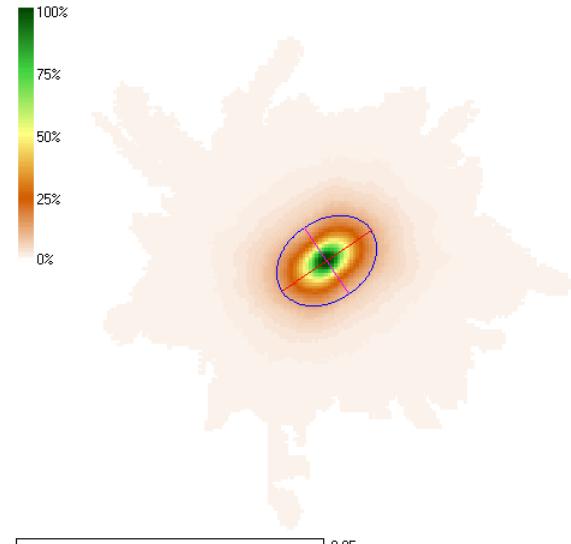


Application to the BSE image of a synthetic magma

Inertia tensor : α

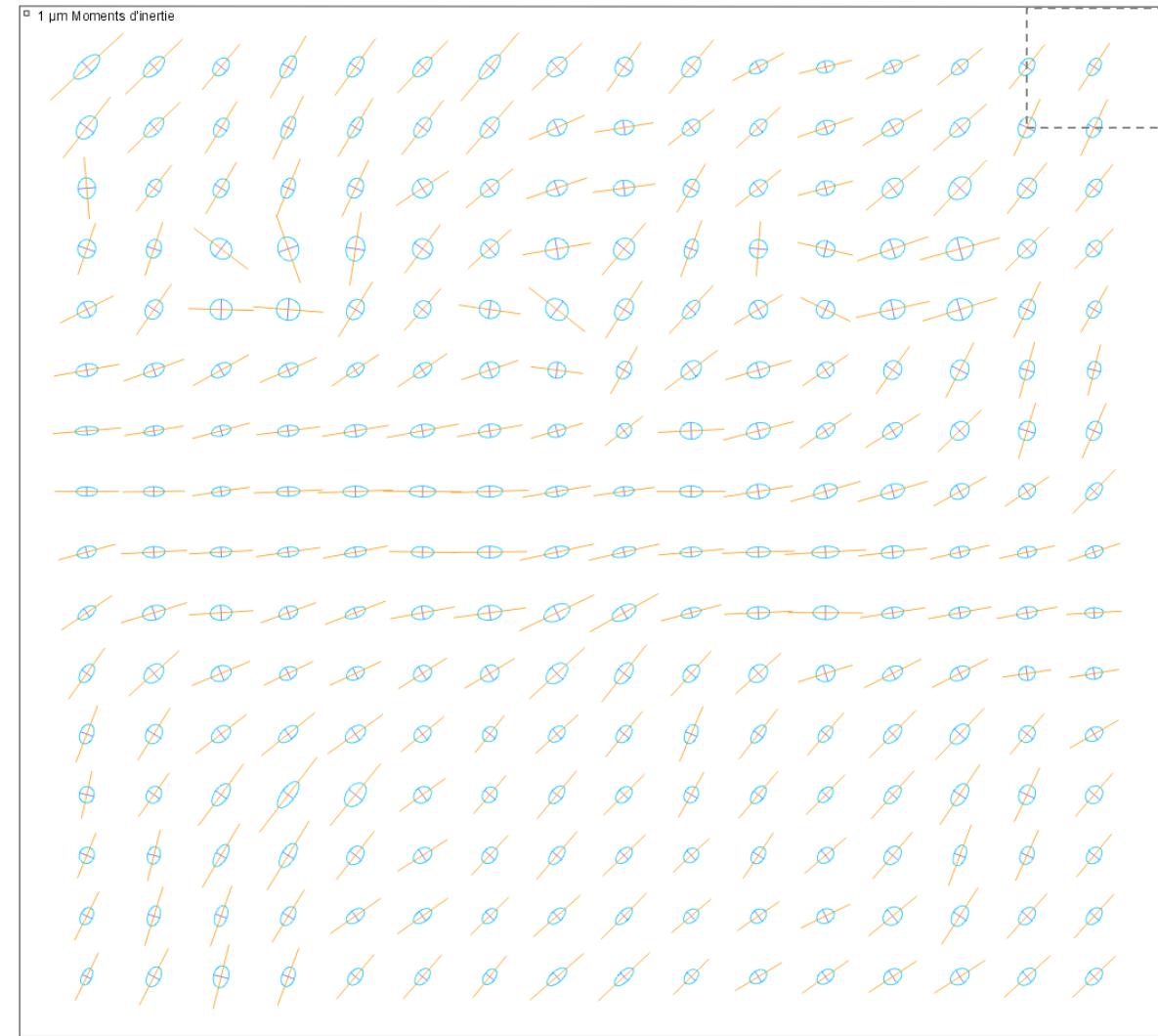


A n=4973 a=0,0176 cm b=0,0131 cm R=1,346 [1,419]n [1,464]b , 55,51°
K=0,288, Kn=0,336 (0,858), Kbn=0,364 (0,793)



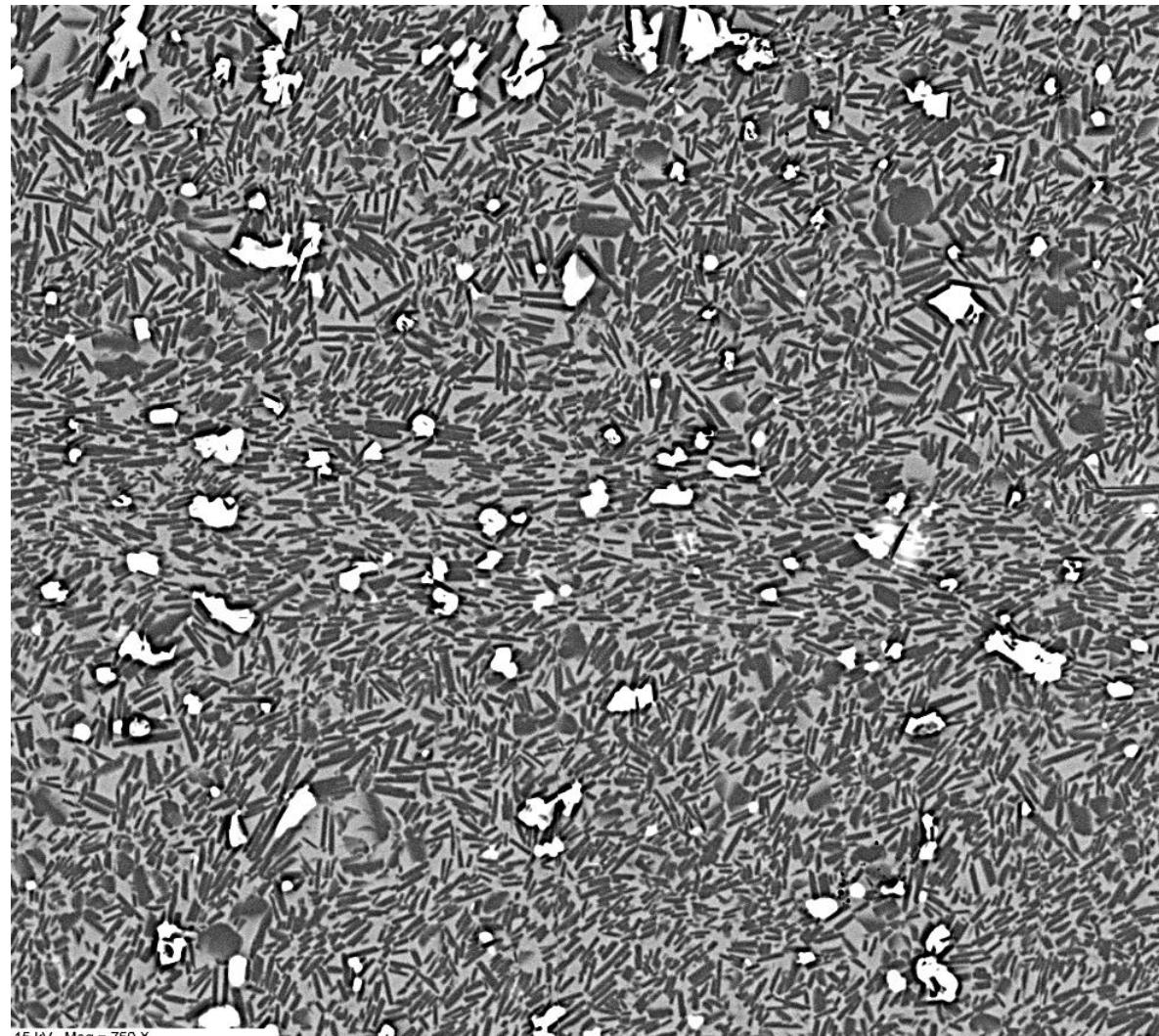
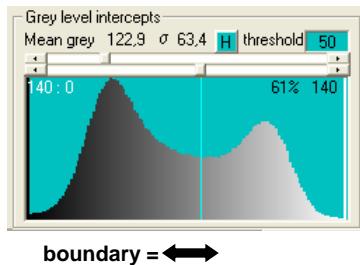
Application to the BSE image of a synthetic magma

Inertia tensor : R, α



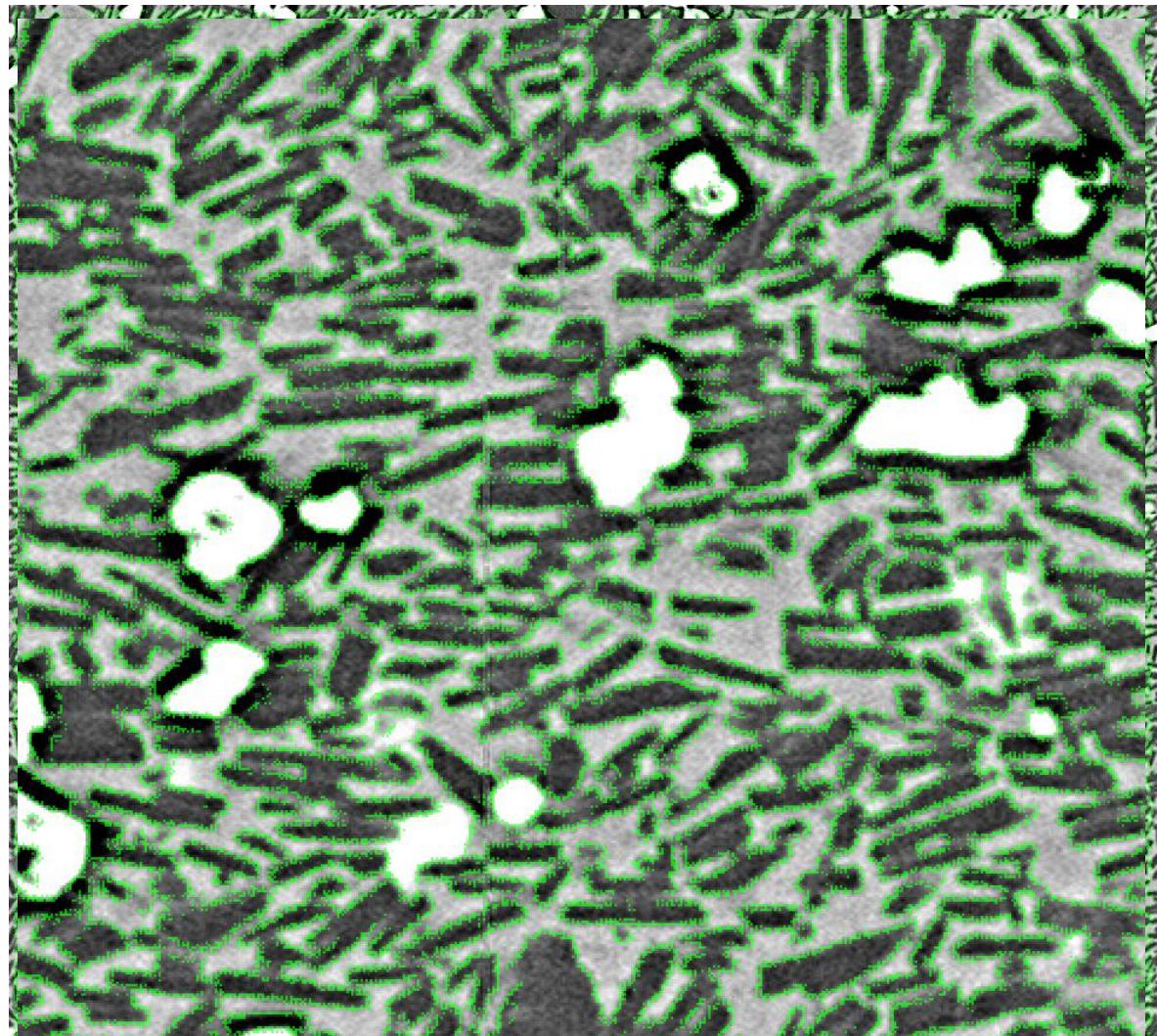
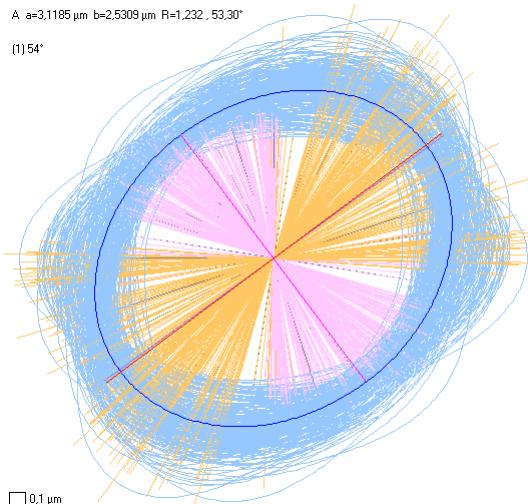
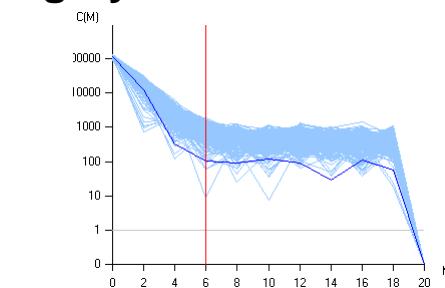
Application to the BSE image of a synthetic magma

Analysis by intercept in
grey level



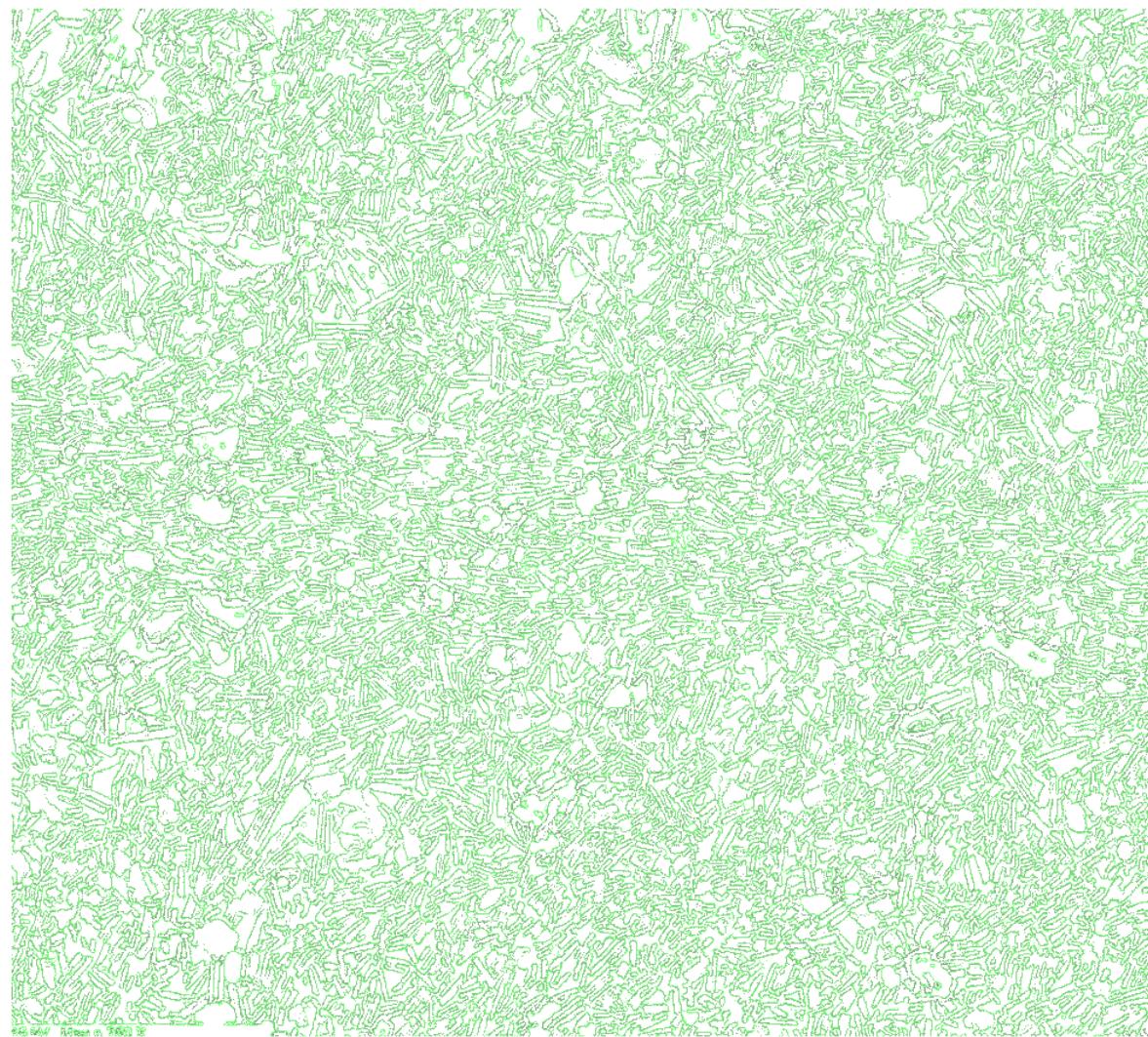
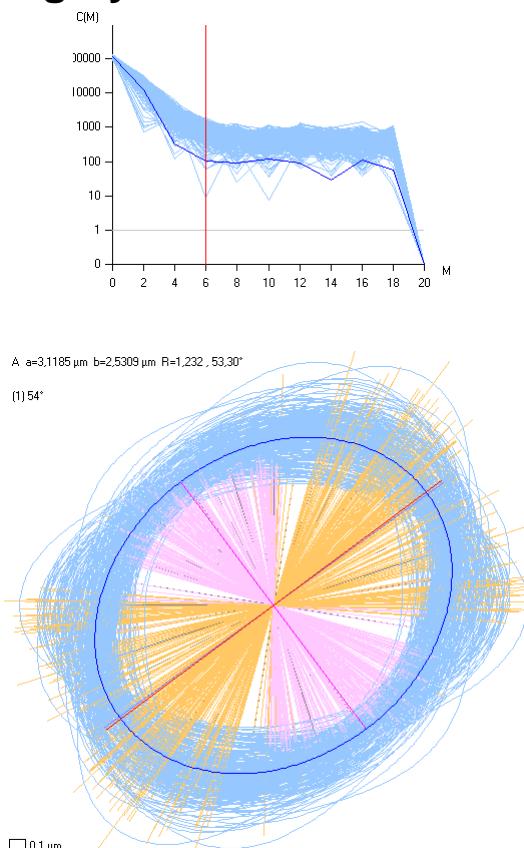
Application to the BSE image of a synthetic magma

Analysis by intercept in grey level



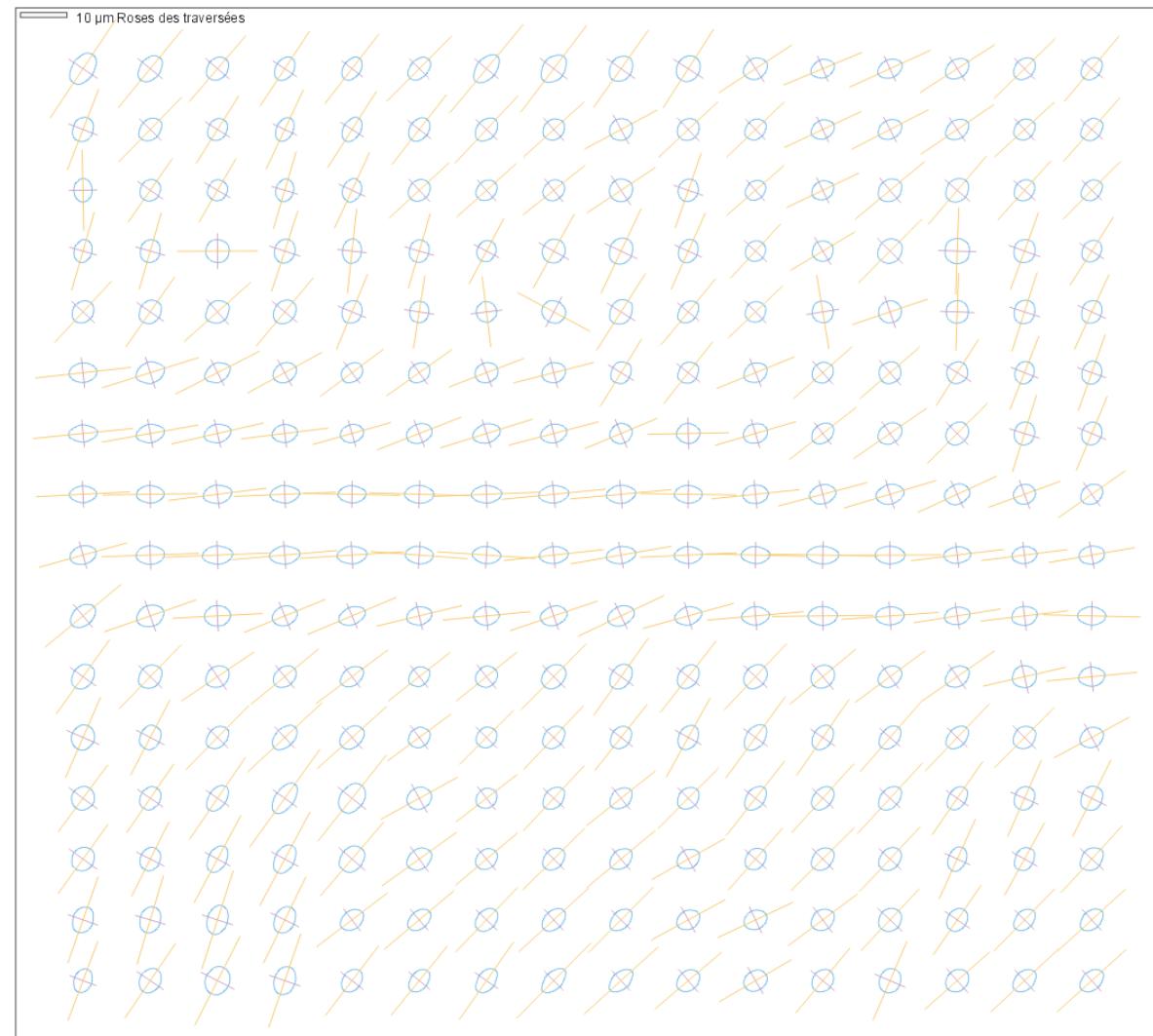
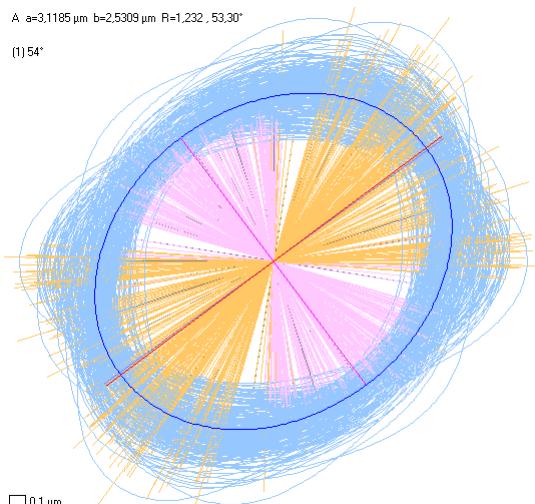
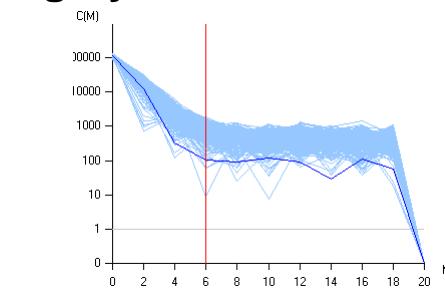
Application to the BSE image of a synthetic magma

Analysis by intercept in grey level



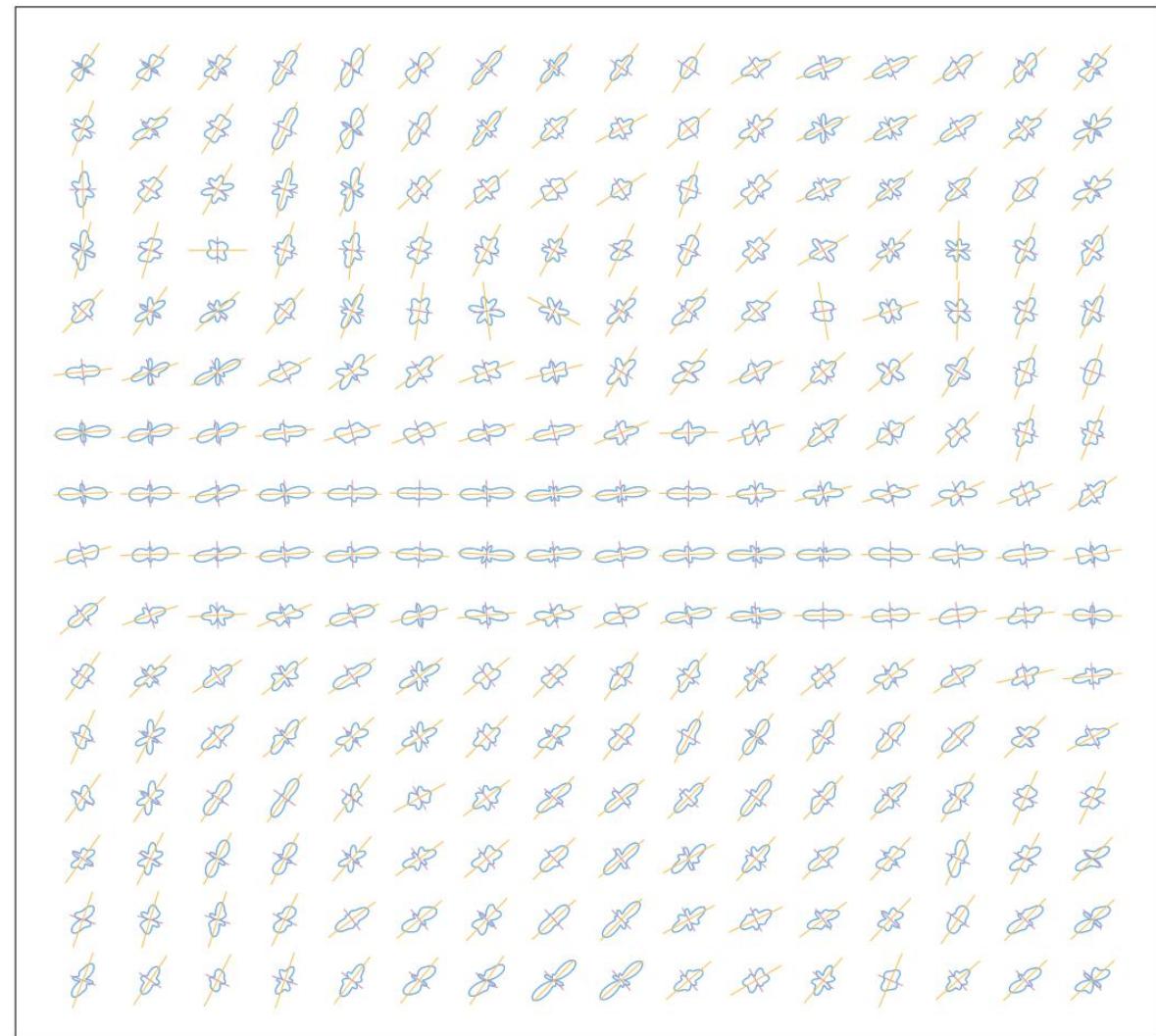
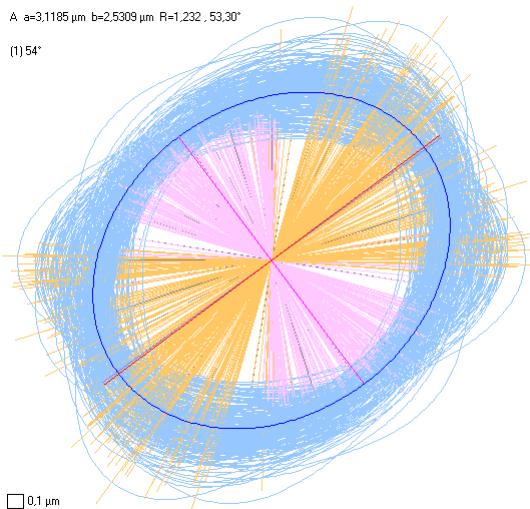
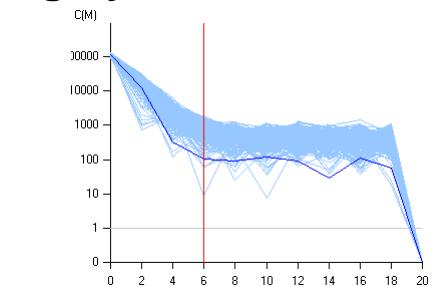
Application to the BSE image of a synthetic magma

Analysis by intercept in grey level



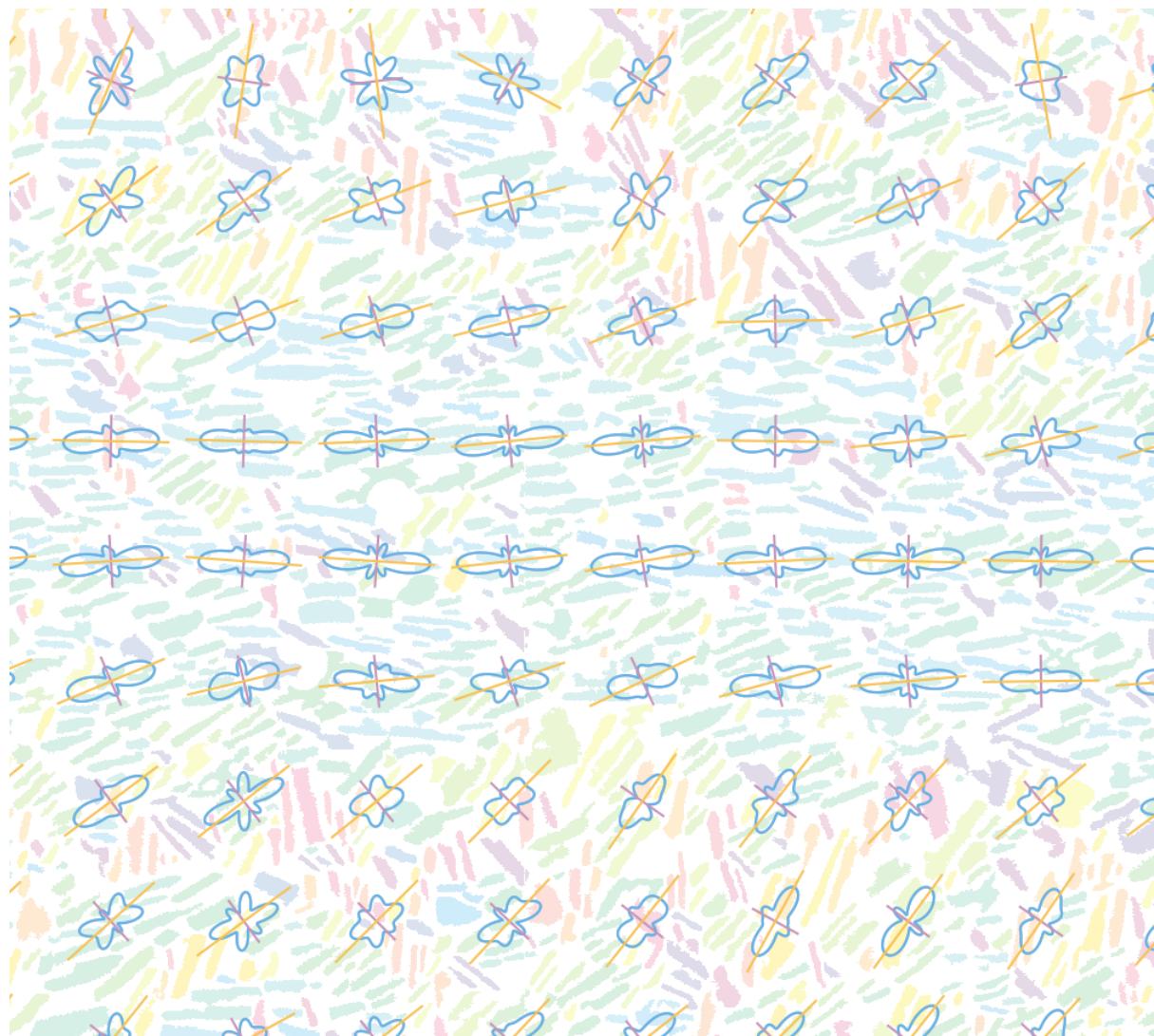
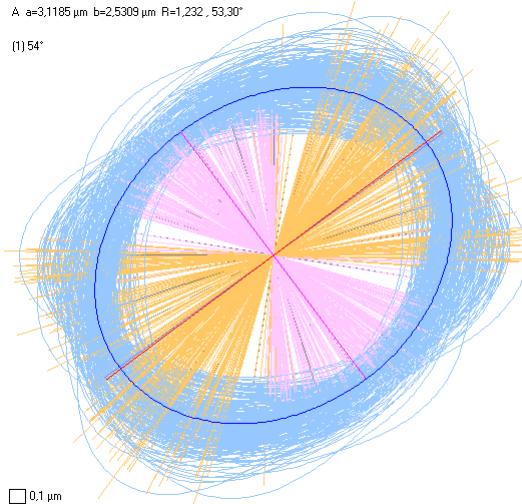
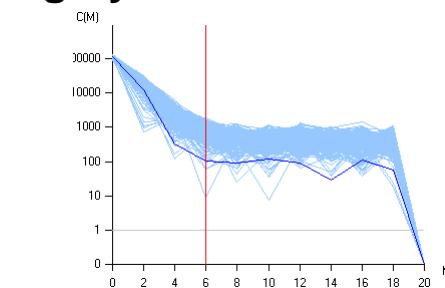
Application to the BSE image of a synthetic magma

Analysis by intercept in grey level



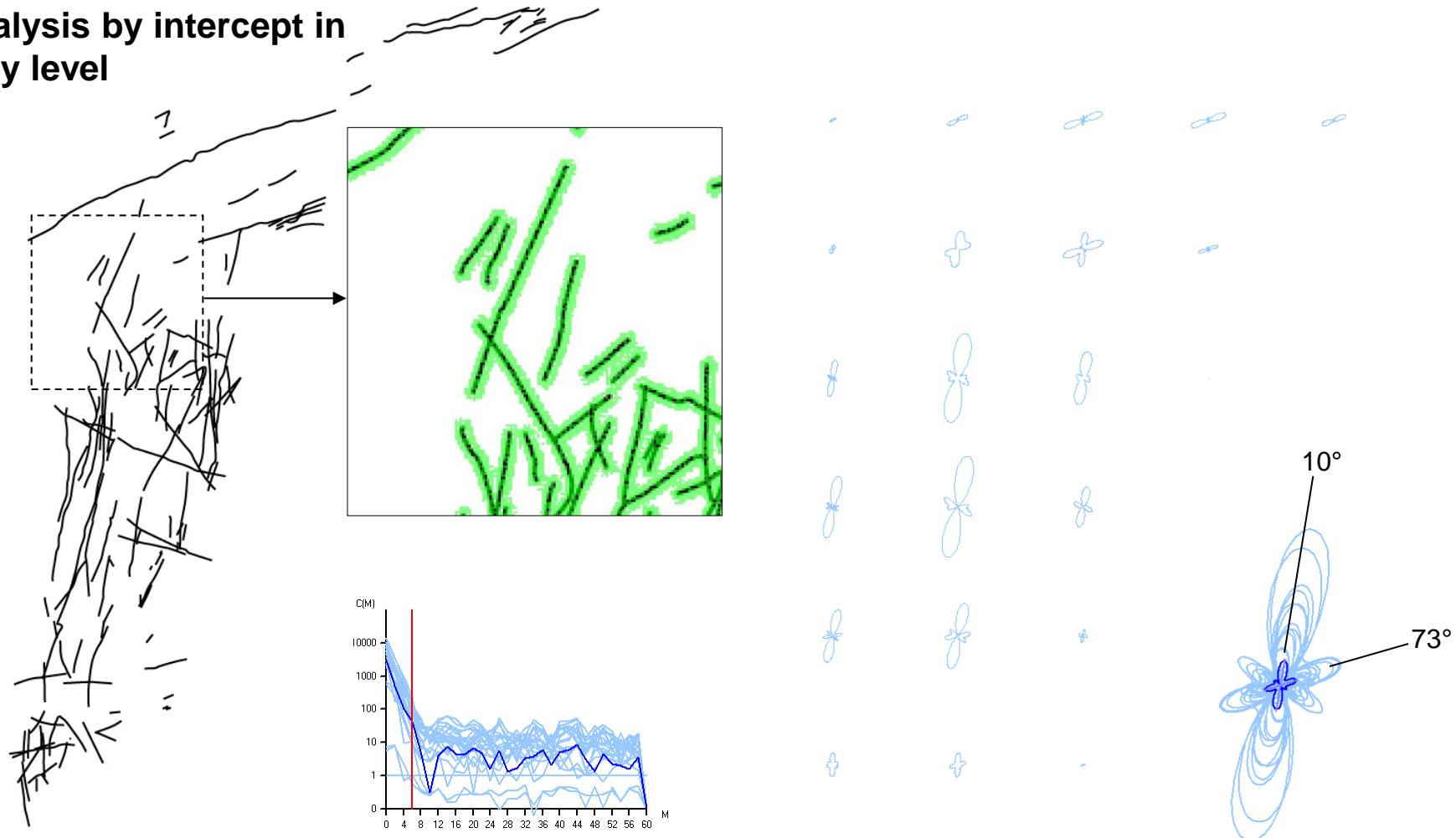
Application to the BSE image of a synthetic magma

Analysis by intercept in grey level



Application to lineaments (Kaapvaal dikes - South Africa)

Analysis by intercept in grey level



Patrick Launeau

Quantitative Image Analysis off Minerals and Rocks

Fabric analysis

- Shape preferred orientation (SPO) vs. strain quantification.
- Intercepts in digital images : a tool to analyze interconnection of grains in rocks vs. inertia tensor of individualized grains
- SPO vs Spatial distribution (Fry)
- Ellipsoid of SPO and strain by combining 3 \perp images.



Ellipsoid2003



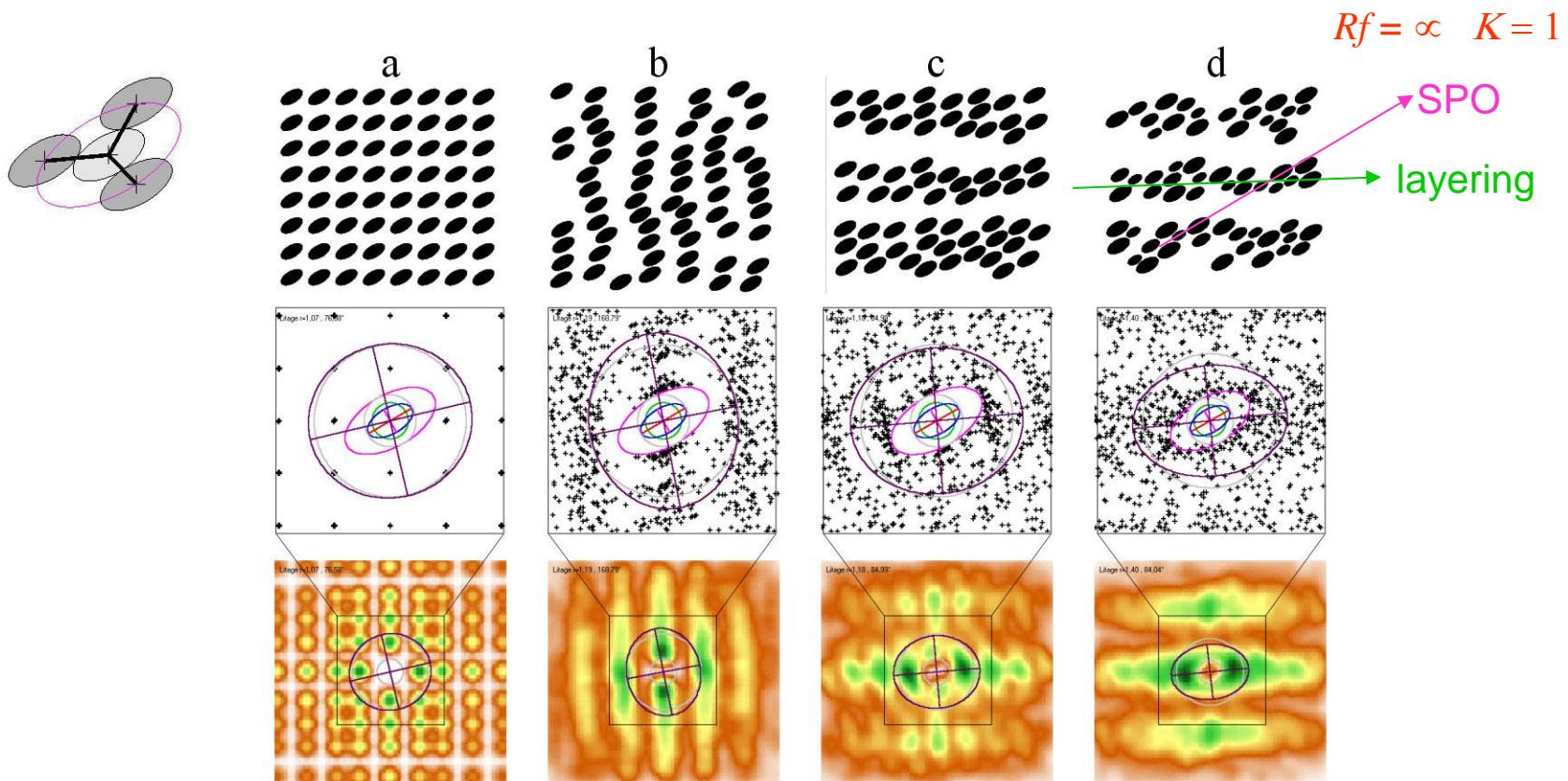
Intercepts2003



SPO2003

Inertia tensor = Fry diagram

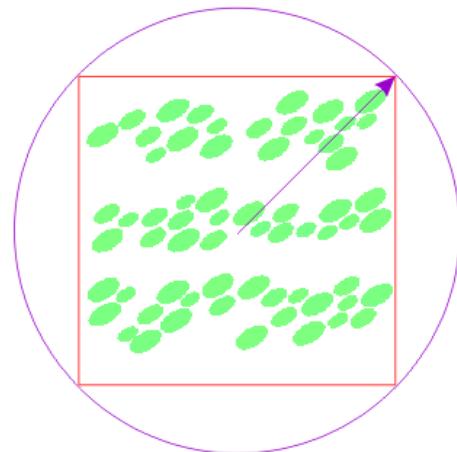
Centre to centre method = Spatial distribution



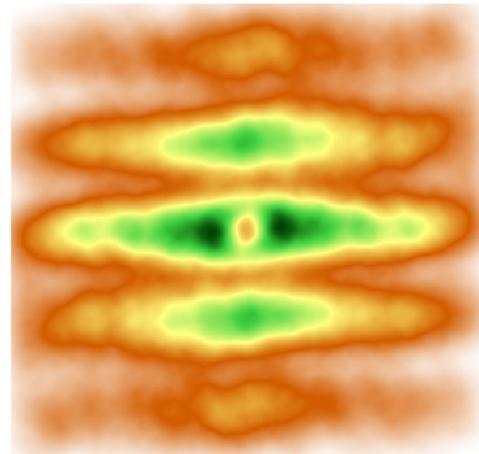
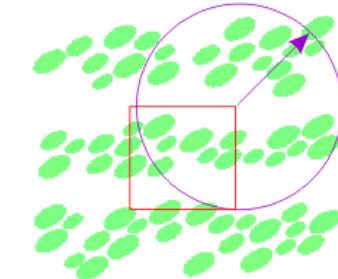
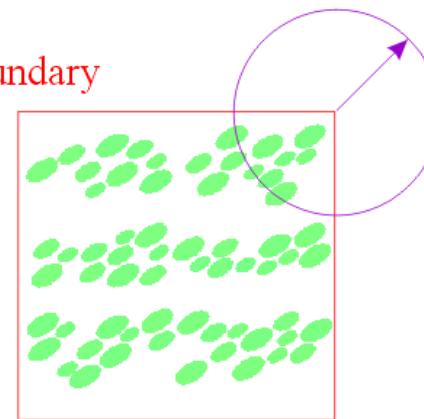
Quantitative Image Analysis of Minerals and Rocks

<http://www.sciences.univ-nantes.fr/lpgnantes/SPO>

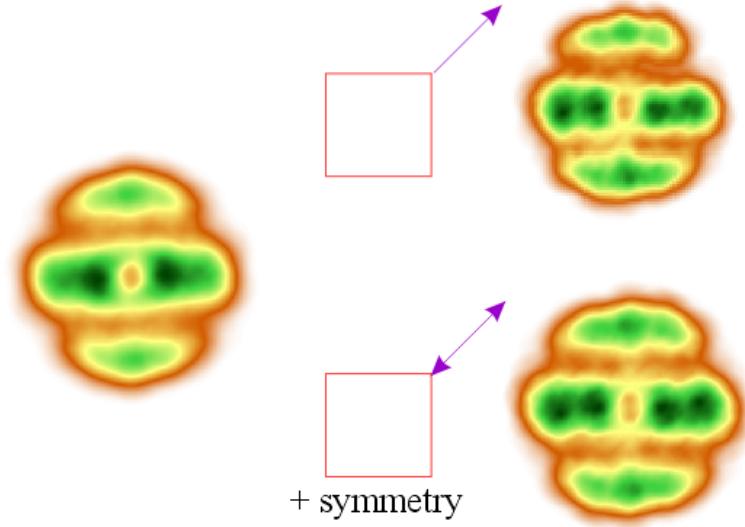
minimum distance



boundary



sensitive to boundary

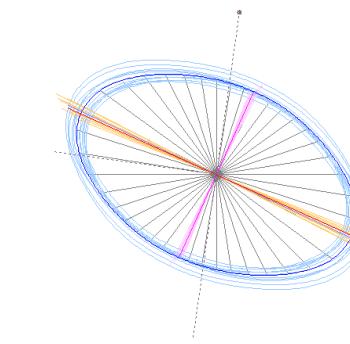
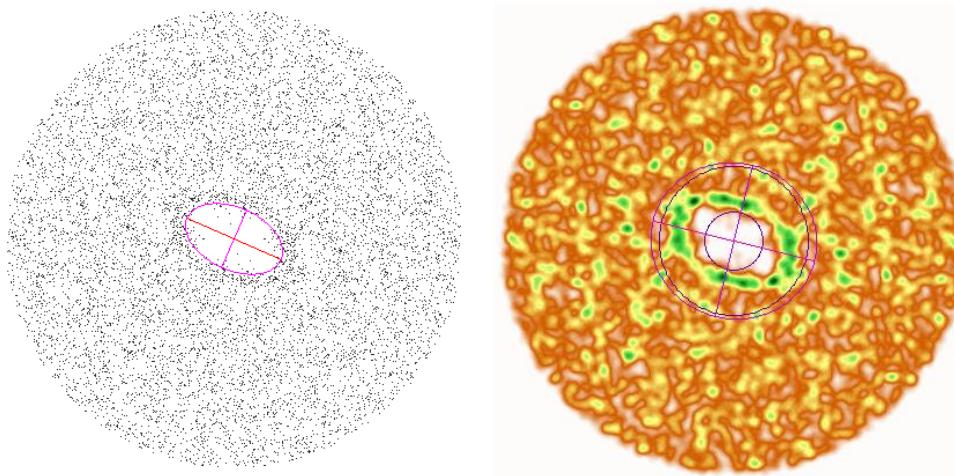
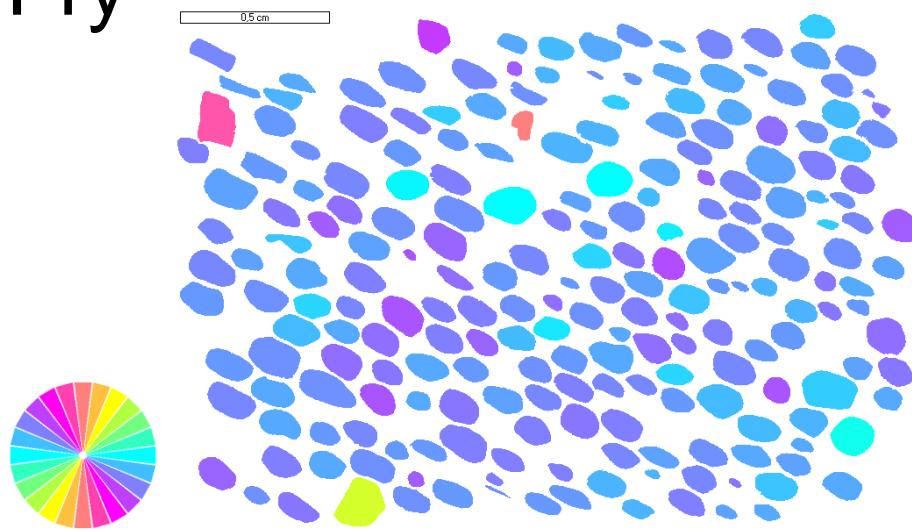


not
sensitive to boundary

Quantitative Image Analysis of Minerals and Rocks

<http://www.sciences.univ-nantes.fr/lpgnantes/SPO>

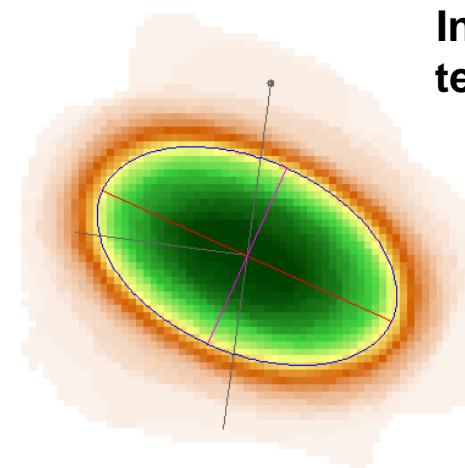
Fry



$R = 1.630 @ 114,3^\circ$

1mm

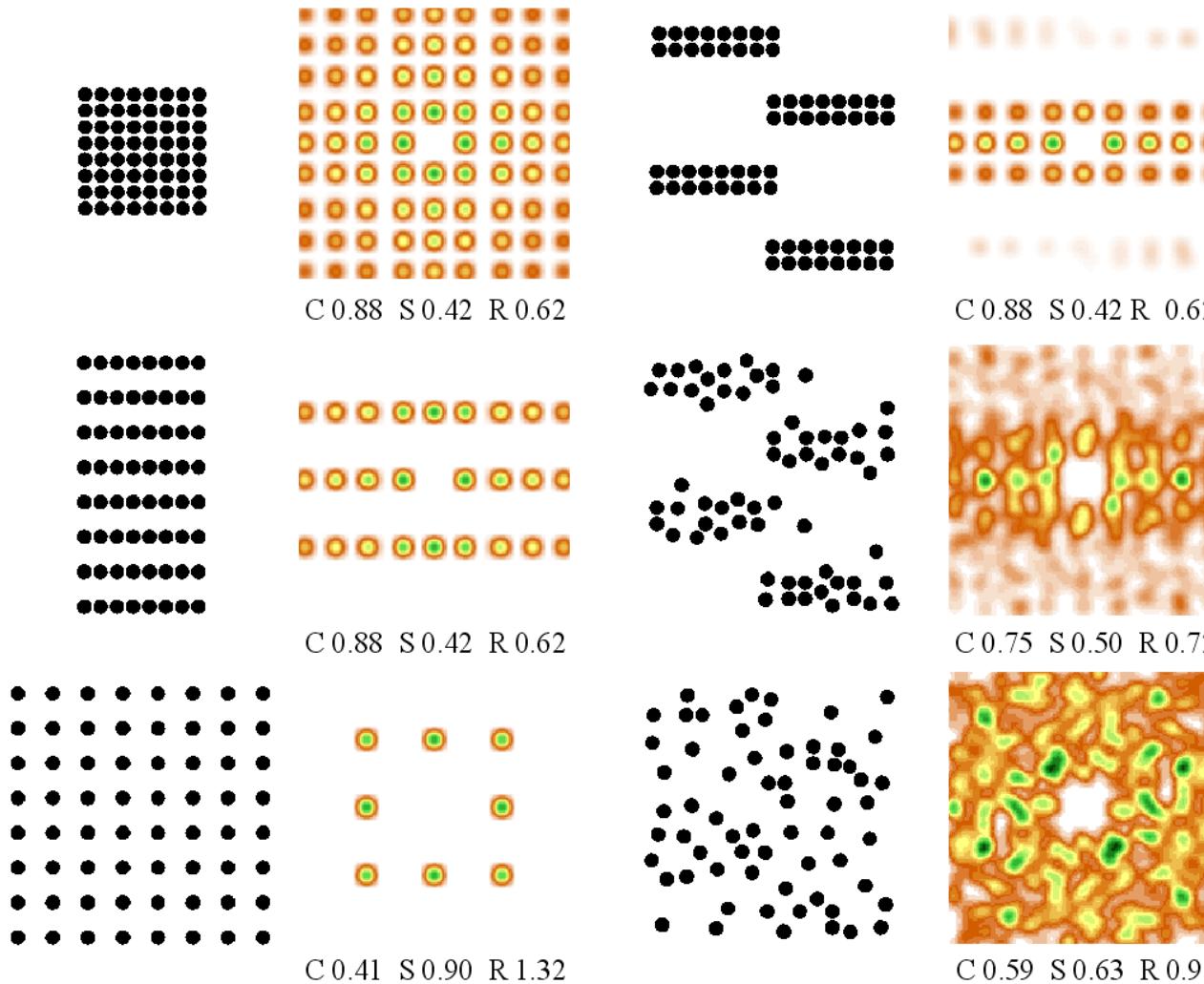
$R = 1.645 @ 114,1^\circ$



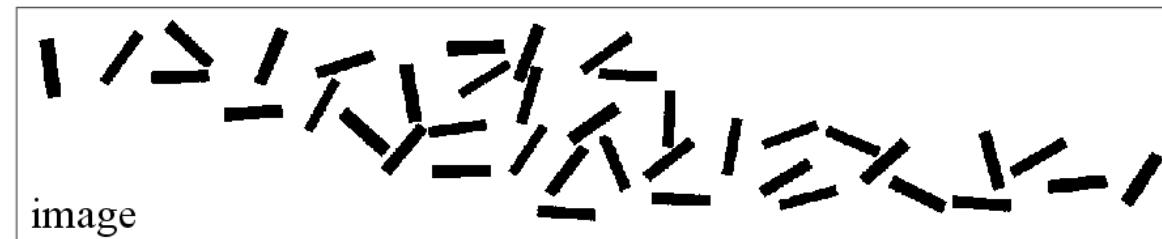
Mean
intercept
length

Inertia
tensor

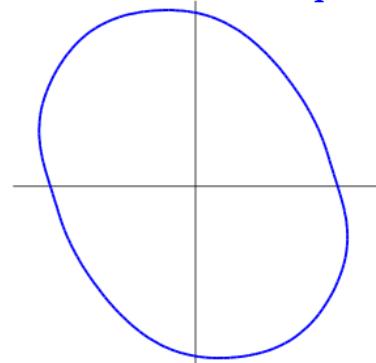
Spatial distribution and Compaction



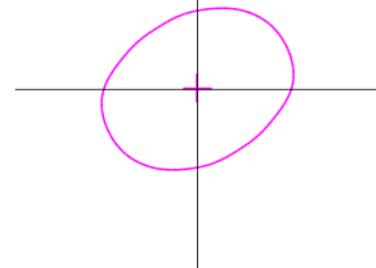
Autocorrelation / intercepts



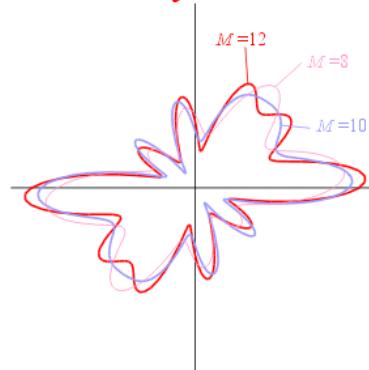
count of intercept



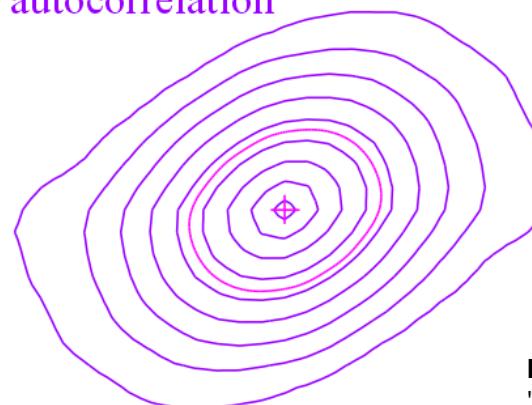
mean intercept length



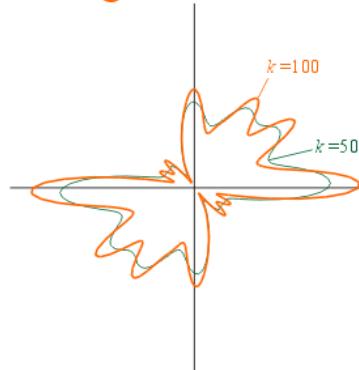
boundary direction



autocorrelation

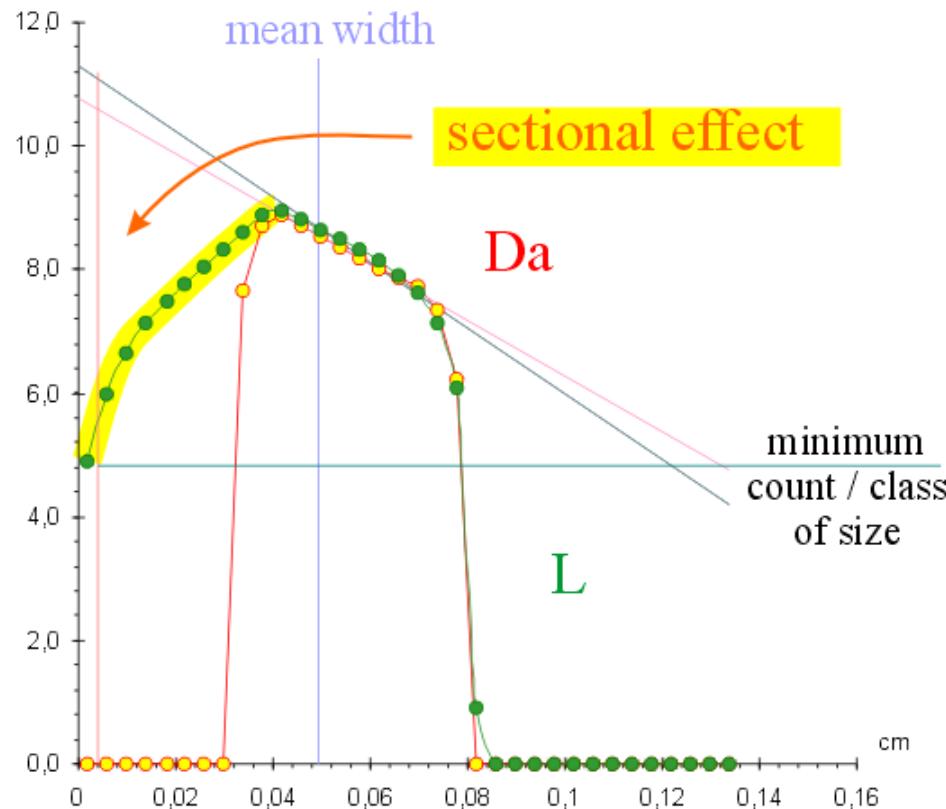


long axis direction

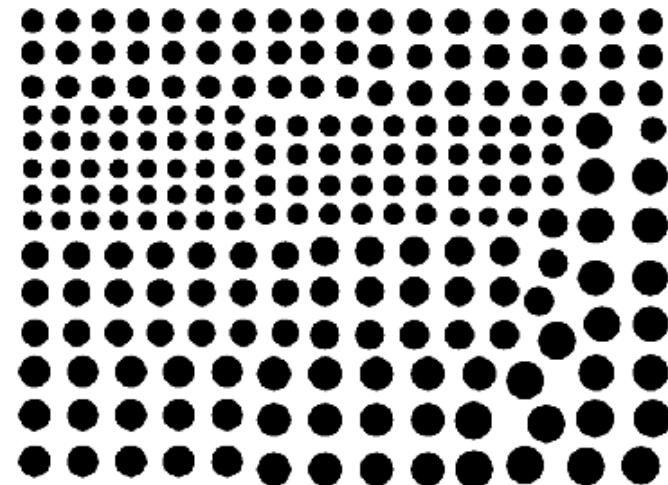


CSD

$\ln(n)$ with n =number of object / surface area / class length

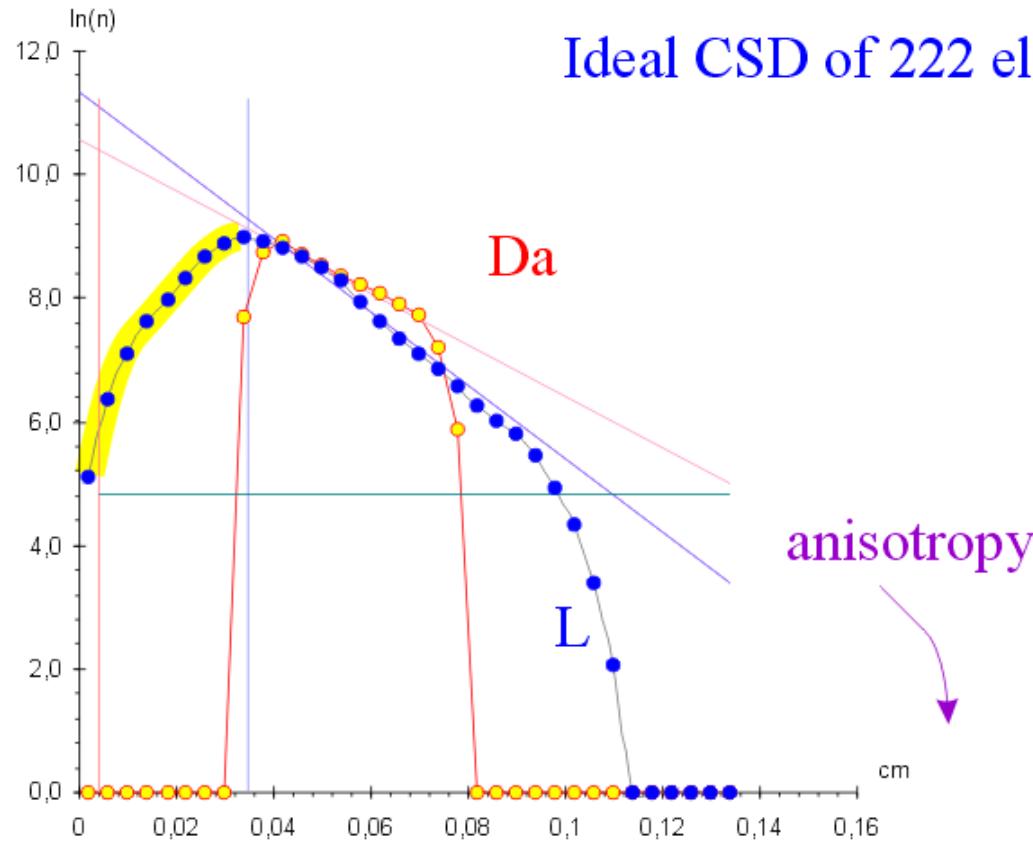


Ideal CSD of 222 circles

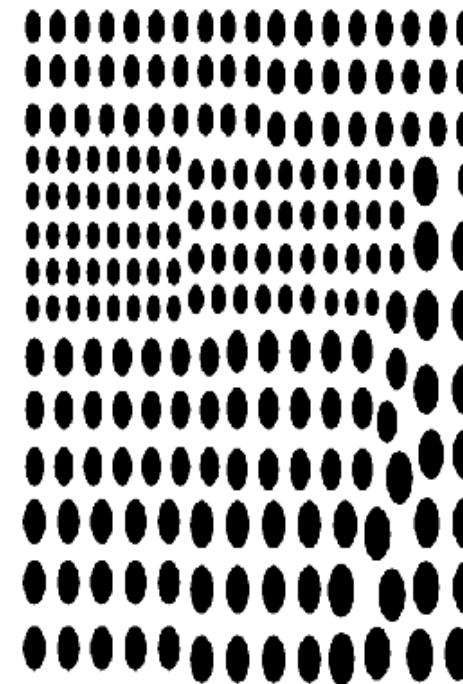


CSD

$\ln(n)$ with n =number of object / surface area / class length

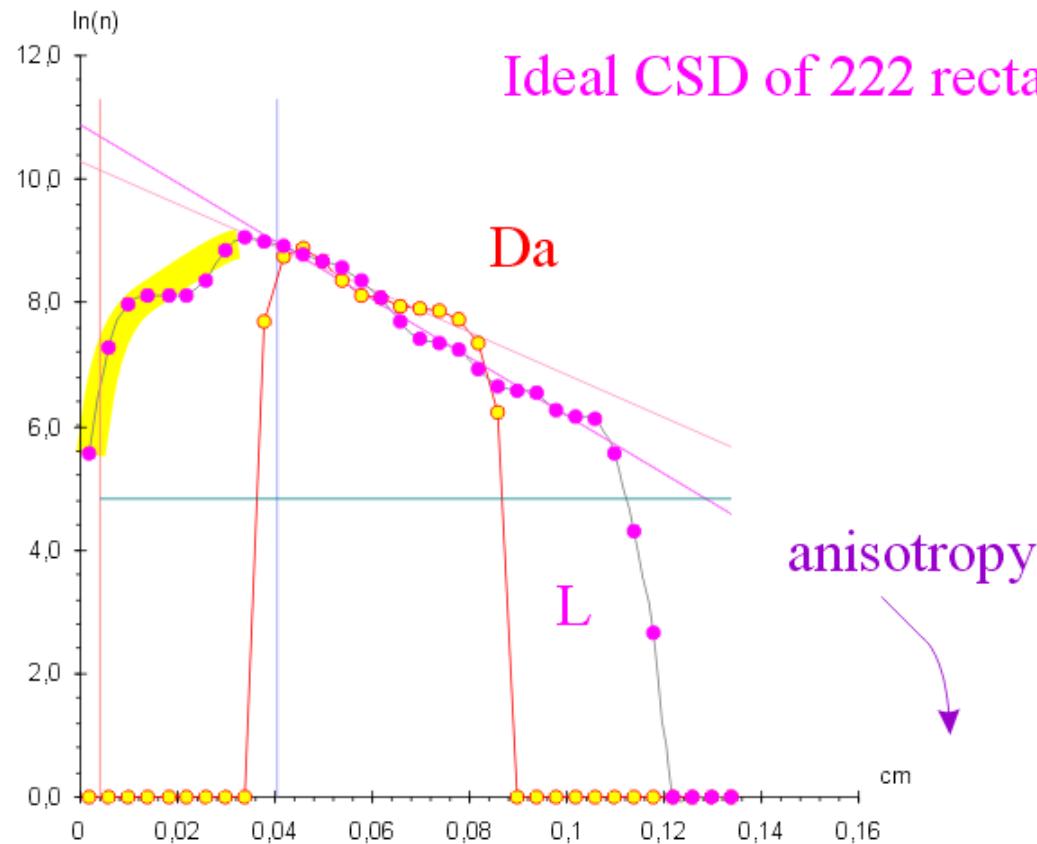


Ideal CSD of 222 ellipses with shape ratio = 2

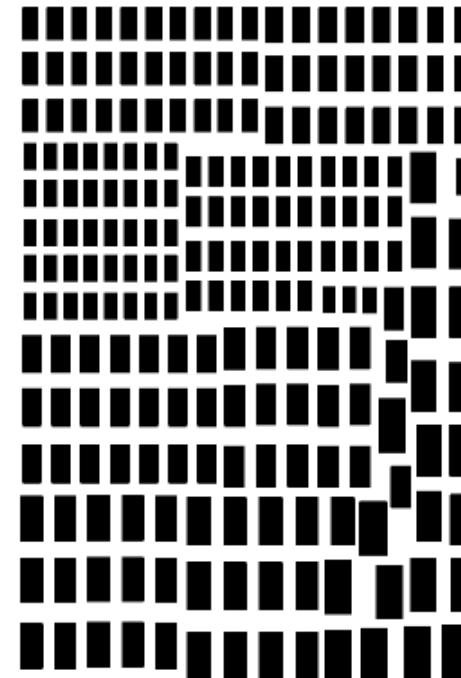


CSD

$\ln(n)$ with n =number of object / surface area / class length

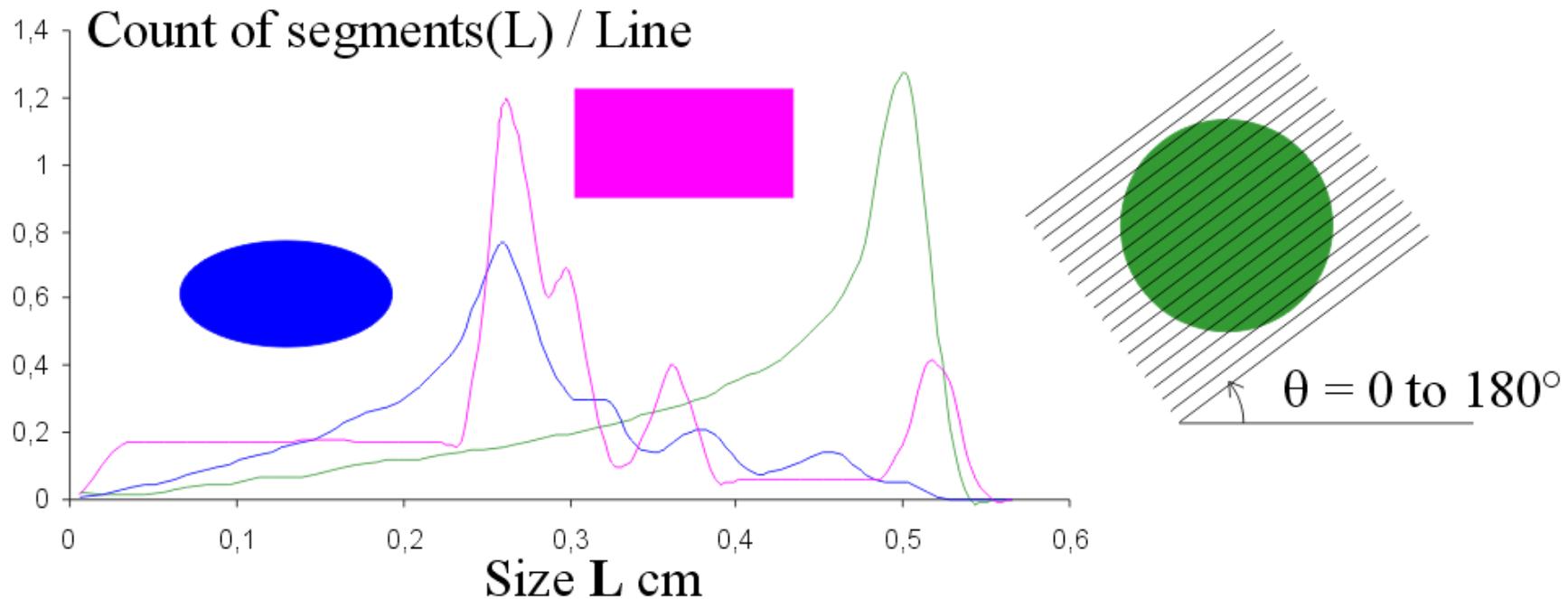


Ideal CSD of 222 rectangles with shape ratio = 2



CSD

Histogram of segments cut by a set of lines between 0 and 180°



Patrick Launeau

Quantitative Image Analysis off Minerals and Rocks

Fabric analysis

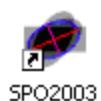
- Shape preferred orientation (SPO) vs. strain quantification.
- Intercepts in digital images : a tool to analyze interconnection of grains in rocks vs. inertia tensor of individualized grains
- SPO vs Spatial distribution (Fry)
- Ellipsoid of SPO and strain by combining 3 \perp images.



Ellipsoid2003



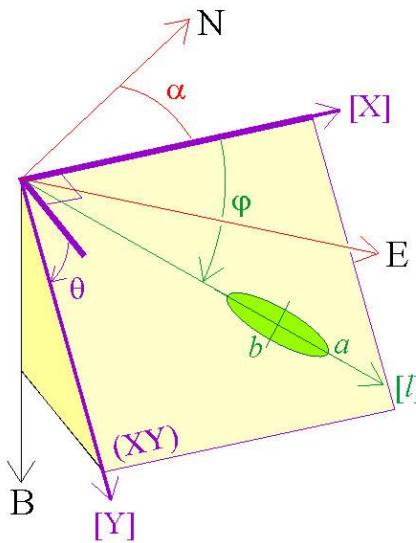
Intercepts2003



SPO2003

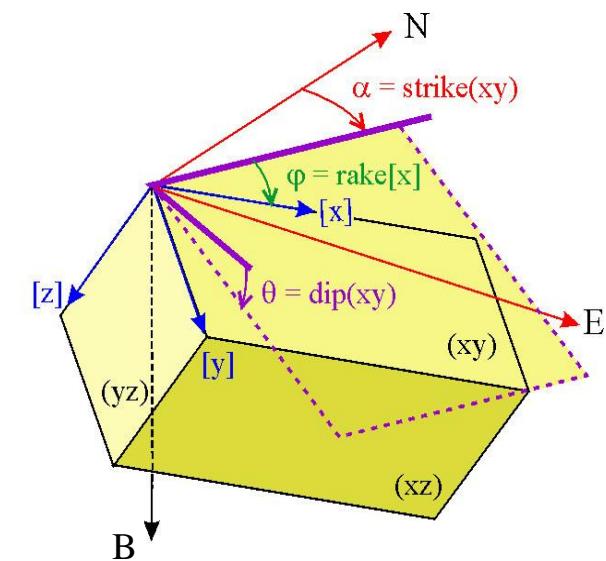
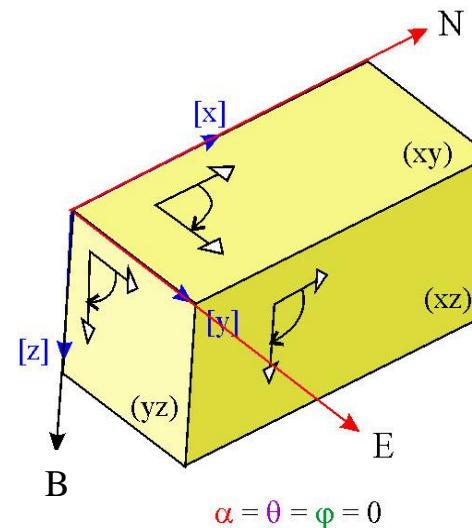
From 2D image analysis to 3D ellipsoid construction

$$\mathbf{R}_L = \begin{bmatrix} \cos\alpha\cos\varphi - \sin\alpha\cos\theta\sin\varphi \\ \sin\alpha\cos\varphi + \cos\alpha\cos\theta\sin\varphi \\ \sin\theta\sin\varphi \end{bmatrix}$$



North,East,Down convention

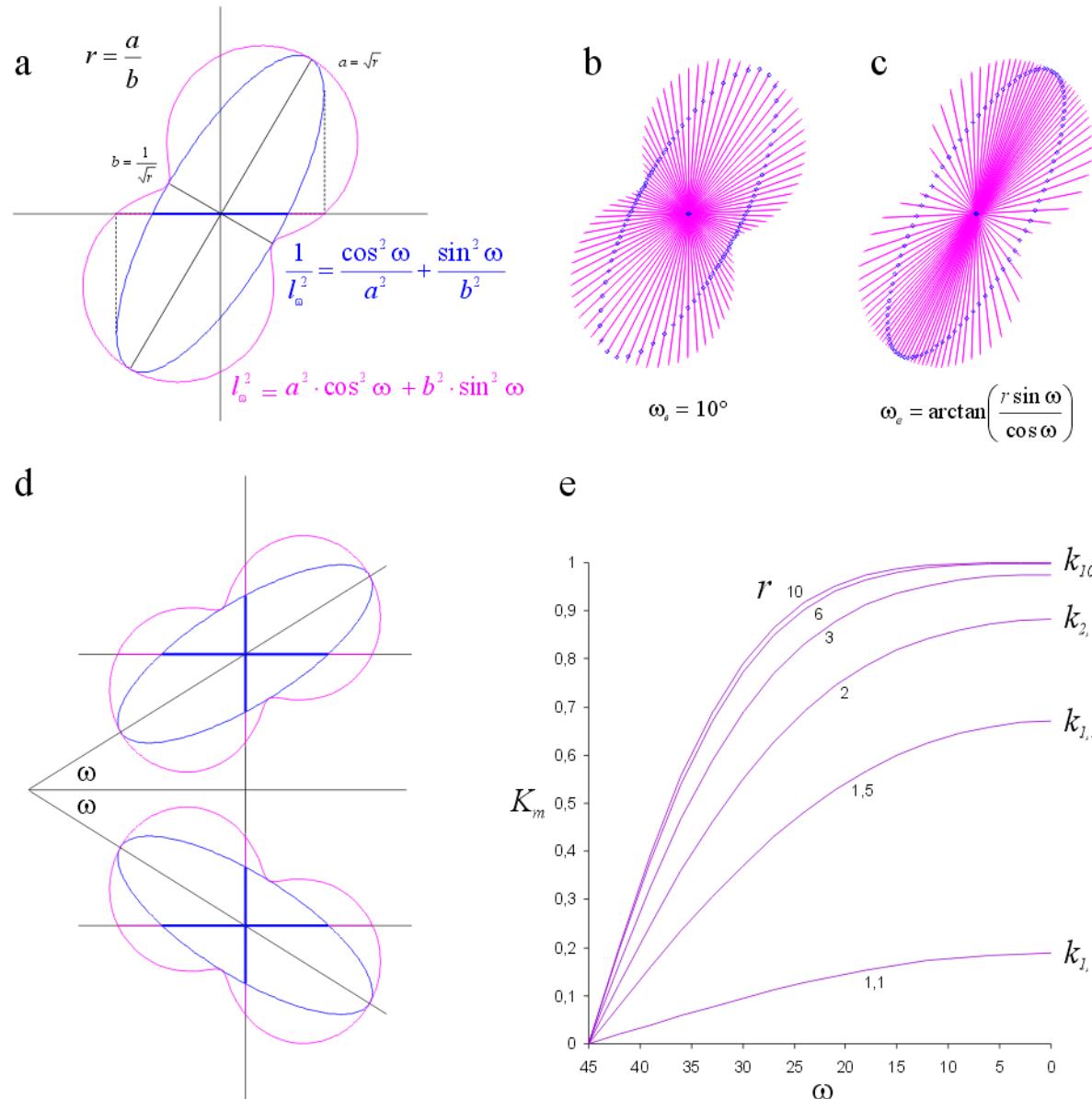
Euler angles : strike α , dip θ , pitch or rake φ



P. Launeau, P.-Y. F. Robin
(2005) "Determination of fabric and strain ellipsoids from measured sectional ellipses—implementation and applications". *Journal of Structural Geology*, 27, 2223-2233

$$\mathbf{R}_V = \begin{bmatrix} \cos\alpha\cos\varphi - \sin\alpha\cos\theta\sin\varphi & -\cos\alpha\sin\varphi - \sin\alpha\cos\theta\cos\varphi & \sin\alpha\sin\theta \\ \sin\alpha\cos\varphi + \cos\alpha\cos\theta\sin\varphi & -\sin\alpha\sin\varphi + \cos\alpha\cos\theta\cos\varphi & -\cos\alpha\sin\theta \\ \sin\theta\sin\varphi & \sin\theta\cos\varphi & \cos\theta \end{bmatrix}$$

From 2D ellipse to 3D ellipsoid



P. Launeau (2004) "Mise en évidence des écoulements magmatiques par analyse d'images 2-D des distributions 3-D d'Orientations Préférentielles de Formes". *Bull. Soc. Géol. Fr.*, 175, 331-350

$$\frac{1}{l(\omega)^2} = \frac{\cos^2 \omega}{a^2} + \frac{\sin^2 \omega}{b^2}$$

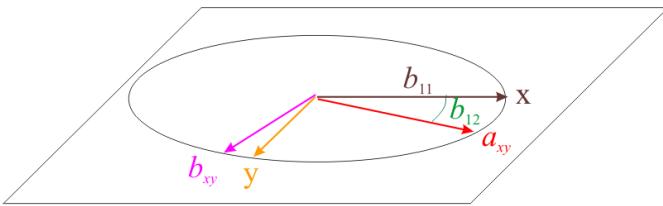
2D ellipses

$$\mathbf{B}_{12} = \begin{bmatrix} b_{11} & b_{12} \\ b_{12} & b_{22} \end{bmatrix} = \begin{bmatrix} \cos \varphi & \sin \varphi \\ -\sin \varphi & \cos \varphi \end{bmatrix} \cdot \begin{bmatrix} 1/a^2 & 0 \\ 0 & 1/b^2 \end{bmatrix} \cdot \begin{bmatrix} \cos \varphi & -\sin \varphi \\ \sin \varphi & \cos \varphi \end{bmatrix}$$

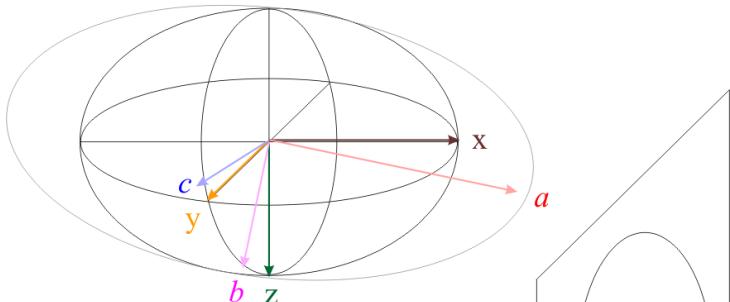
& 3D ellipsoids

$$\frac{1}{l(x, y, z)^2} = \frac{x^2}{a^2} + \frac{y^2}{b^2} + \frac{z^2}{c^2}$$

$$\mathbf{B}_{123} = \begin{bmatrix} b_{11} & b_{12} & b_{13} \\ b_{12} & b_{22} & b_{23} \\ b_{13} & b_{23} & b_{33} \end{bmatrix} = \mathbf{R}_V^T \cdot \begin{bmatrix} 1/a^2 & 0 & 0 \\ 0 & 1/b^2 & 0 \\ 0 & 0 & 1/c^2 \end{bmatrix} \cdot \mathbf{R}_V$$

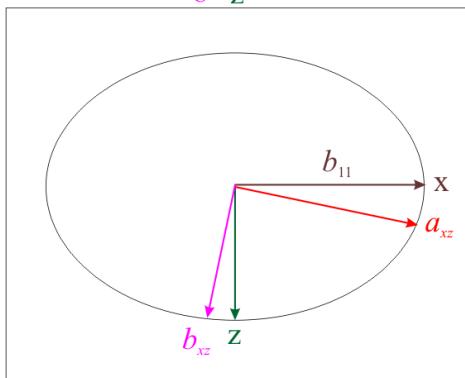


$$\mathbf{B}_{12} = \begin{bmatrix} b_{11} & b_{12} \\ b_{12} & b_{22} \end{bmatrix} = \begin{bmatrix} \cos \varphi_{12} & \sin \varphi_{12} \\ -\sin \varphi_{12} & \cos \varphi_{12} \end{bmatrix} \cdot \begin{bmatrix} 1/a_{12}^2 & 0 \\ 0 & 1/b_{12}^2 \end{bmatrix} \cdot \begin{bmatrix} \cos \varphi_{12} & -\sin \varphi_{12} \\ \sin \varphi_{12} & \cos \varphi_{12} \end{bmatrix}$$



$$\mathbf{B}_{13} = \begin{bmatrix} b_{11} & b_{13} \\ b_{13} & b_{33} \end{bmatrix} = \begin{bmatrix} \cos \varphi_{13} & \sin \varphi_{13} \\ -\sin \varphi_{13} & \cos \varphi_{13} \end{bmatrix} \cdot \begin{bmatrix} 1/a_{13}^2 & 0 \\ 0 & 1/b_{13}^2 \end{bmatrix} \cdot \begin{bmatrix} \cos \varphi_{13} & -\sin \varphi_{13} \\ \sin \varphi_{13} & \cos \varphi_{13} \end{bmatrix}$$

$$\mathbf{B}_{23} = \begin{bmatrix} b_{22} & b_{23} \\ b_{23} & b_{33} \end{bmatrix} = \begin{bmatrix} \cos \varphi_{23} & \sin \varphi_{23} \\ -\sin \varphi_{23} & \cos \varphi_{23} \end{bmatrix} \cdot \begin{bmatrix} 1/a_{23}^2 & 0 \\ 0 & 1/b_{23}^2 \end{bmatrix} \cdot \begin{bmatrix} \cos \varphi_{23} & -\sin \varphi_{23} \\ \sin \varphi_{23} & \cos \varphi_{23} \end{bmatrix}$$



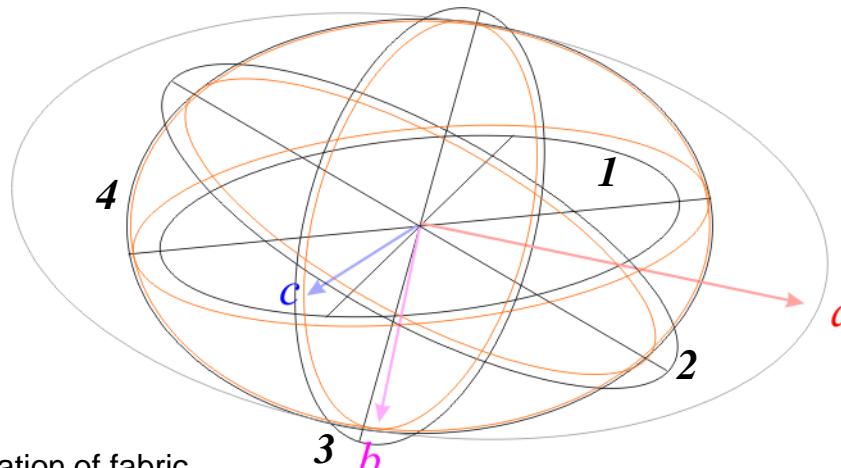
$$\mathbf{B} = \begin{bmatrix} \bar{b}_{11} & \bar{b}_{12} & b_{13} \\ b_{12} & \bar{b}_{22} & b_{23} \\ b_{13} & b_{23} & \bar{b}_{33} \end{bmatrix}$$

(x,y,z) (1,2,3)

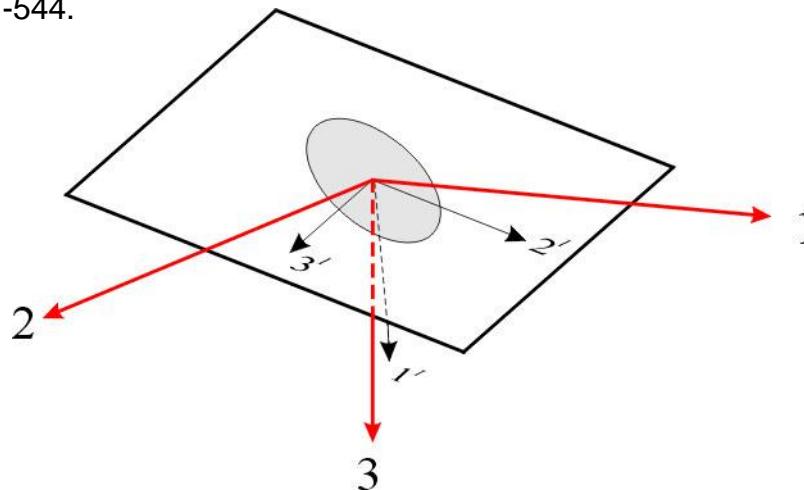
Launeau, P. and Cruden, A.R. (1998). – Magmatic fabric acquisition mechanism in a syenite: Results of a combined anisotropy of magnetic susceptibility and image analysis study. J. Geophys. Res. 103, 5067-5089.

2D ellipses & 3D ellipsoids

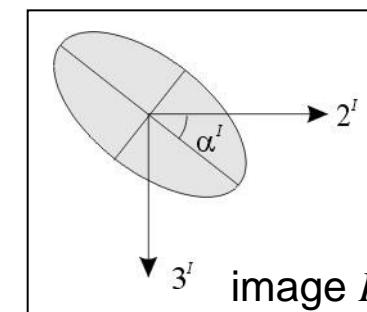
$I = 1 \text{ to } 4$



Robin, P.-Y.F. (2002). – Determination of fabric and strain ellipsoids from measured sectional ellipses – Theory. *Journal of Structural Geology*, 24, 531-544.



$$\mathbf{B}_{23}^I = \mathbf{L}_{23}^{IT} \cdot \mathbf{B}_{123} \cdot \mathbf{L}_{23}^I + \mathbf{X}^I$$



$$\text{with } \mathbf{X}^I = \begin{bmatrix} \chi_{22}^I & \chi_{23}^I \\ \chi_{23}^I & \chi_{33}^I \end{bmatrix}$$

deviations to measurements

$$\rightarrow \mathbf{B}_{23}^I = \begin{bmatrix} b_{22}^I & b_{23}^I \\ b_{23}^I & b_{33}^I \end{bmatrix}$$

of the section

$$\mathbf{L}_{23}^I = \begin{bmatrix} l_{21}^I & l_{31}^I \\ l_{22}^I & l_{32}^I \\ l_{23}^I & l_{33}^I \end{bmatrix}$$

of ellipsoid \mathbf{B}_{123}

Scale factors of the Robin 2002 method

Scale factor available

$$\mathbf{B}_{23}^I = \begin{bmatrix} \cos\varphi & \sin\varphi \\ -\sin\varphi & \cos\varphi \end{bmatrix} \cdot \begin{bmatrix} 1/a^2 & 0 \\ 0 & 1/b^2 \end{bmatrix} \cdot \begin{bmatrix} \cos\varphi & -\sin\varphi \\ \sin\varphi & \cos\varphi \end{bmatrix}$$

$$\frac{1}{\bar{a}^I} \cdot \begin{bmatrix} 1/\rho & 0 \\ 0 & \rho \end{bmatrix}$$

$$\bar{n}^I \cdot \begin{bmatrix} 1/\rho & 0 \\ 0 & \rho \end{bmatrix}$$

length a et b

mean surface area

with $\rho = \frac{a}{b}$ the shape ratio

mean density

$$\bar{a}^I = A^I / j^I$$

$$\bar{n}^I = j^I / A^I$$

Robin, P.-Y.F.
(2002). –

Determination of fabric and strain ellipsoids from measured sectional ellipses – Theory.

Journal of Structural Geology, 24, 531-544.

$$\mathbf{B}_{23}^I = f^I \cdot \begin{bmatrix} \cos\varphi & \sin\varphi \\ -\sin\varphi & \cos\varphi \end{bmatrix} \cdot \begin{bmatrix} 1/\rho & 0 \\ 0 & \rho \end{bmatrix} \cdot \begin{bmatrix} \cos\varphi & -\sin\varphi \\ \sin\varphi & \cos\varphi \end{bmatrix}$$

No scale factor available

$\sqrt{F_{\min}^I}$ Compatibility index of section I

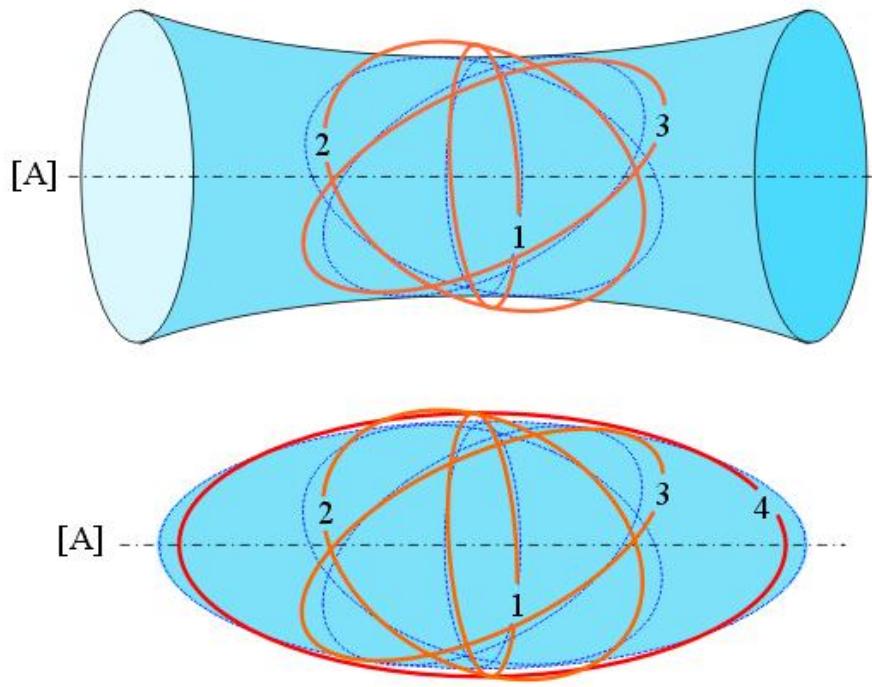
In case of 3 perpendicular sections

$$\begin{bmatrix} b_{11} & b_{12} & b_{13} \\ b_{21} & b_{22} & b_{23} \\ b_{31} & b_{23} & b_{33} \end{bmatrix} = \begin{bmatrix} (b_{33}^2 + b_{22}^2)/2 & b_{23}^3 & b_{23}^2 \\ b_{23}^3 & (b_{22}^1 + b_{33}^3)/2 & b_{23}^1 \\ b_{23}^2 & b_{23}^1 & (b_{33}^1 + b_{22}^2)/2 \end{bmatrix}$$

Compatibility index of all sections

$$\tilde{F} = \frac{1}{6} \frac{(b_{33}^2 - b_{22}^3)^2 + (b_{22}^1 - b_{33}^3)^2 + (b_{33}^1 - b_{22}^2)^2}{(b_1^1)^2 + (b_2^1)^2 + (b_3^1)^2}$$

Convergence of the Robin 2002 method



Robin, P.-Y.F. (2002). – Determination of fabric and strain ellipsoids from measured sectional ellipses – Theory. *Journal of Structural Geology*, 24, 531-544.

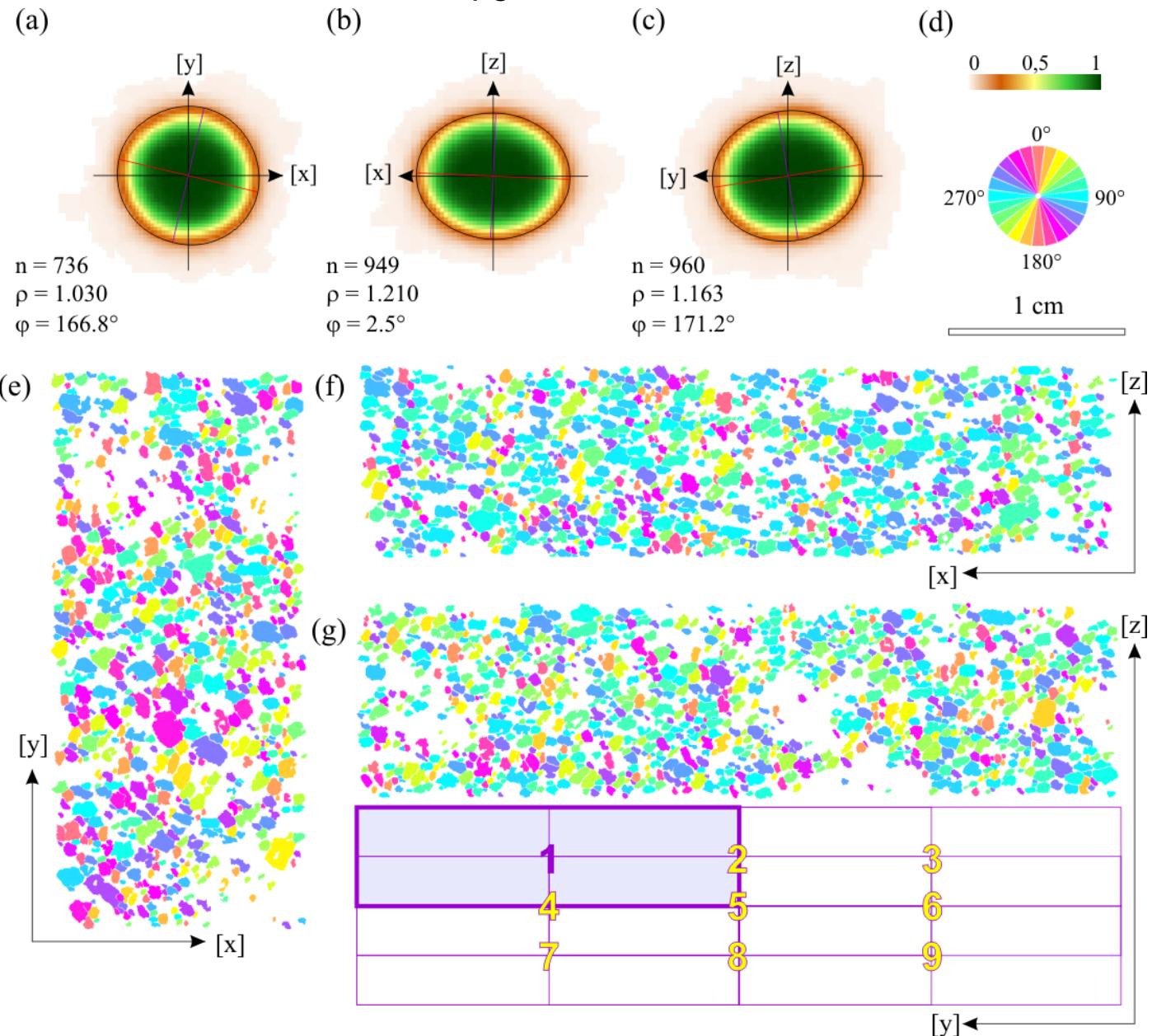
Quantitative Image Analysis of Minerals and Rocks

<http://www.sciences.univ-nantes.fr/lpgnantes/SPO>

Application to gabbronorites of the Bushveld complex

Orthopyroxene

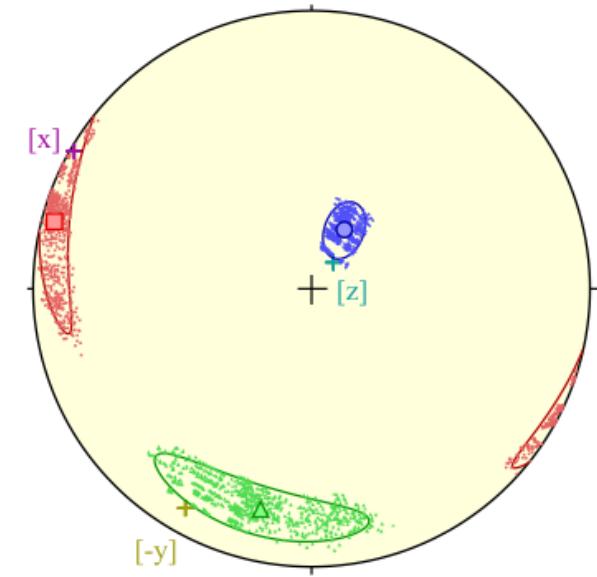
P. Launeau, P.-Y. F. Robin
(2005) "Determination of fabric and strain ellipsoids from measured sectional ellipses—implementation and applications". *Journal of Structural Geology*, 27, 2223-2233



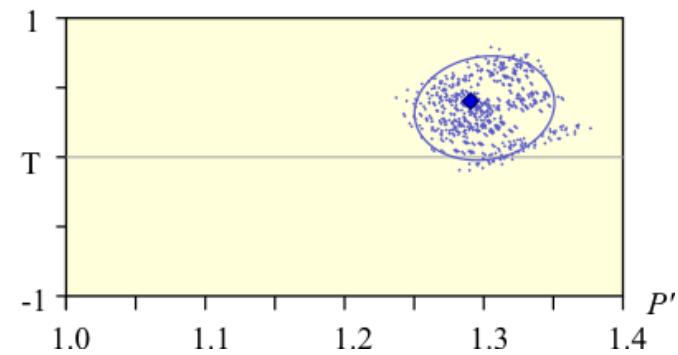
Application to gabbronorites of the Bushveld complex

(h)	A	B	C	$\sqrt{\tilde{F}}$
L. norm	1.113	1.034	0.869	2.5%
trend	284.7	192.9	29.0	
plunge	5.1	19.1	70.2	
A/C	1.282		Flinn	0.405
A/B	1.077		P'	1.291
B/C	1.192		T	0.396
				729 combinations
L. norm	1.120	1.030	0.867	\bar{X}
	0.028	0.033	0.021	2σ
trend	285.9	194.2	29.1	\bar{X}
plunge	4.7	19.0	70.4	\bar{X}
	27.4	27.4	8.8	$2\sigma_1$
	5.6	8.8	5.7	$2\sigma_2$
	\bar{X}	2σ	\bar{X}	2σ
A/C	1.292	0.048	$\sqrt{\tilde{F}}$	2.6%
A/B	1.087	0.057	P'	1.301
B/C	1.189	0.060	T	0.352
				0.375

(i)



(j)



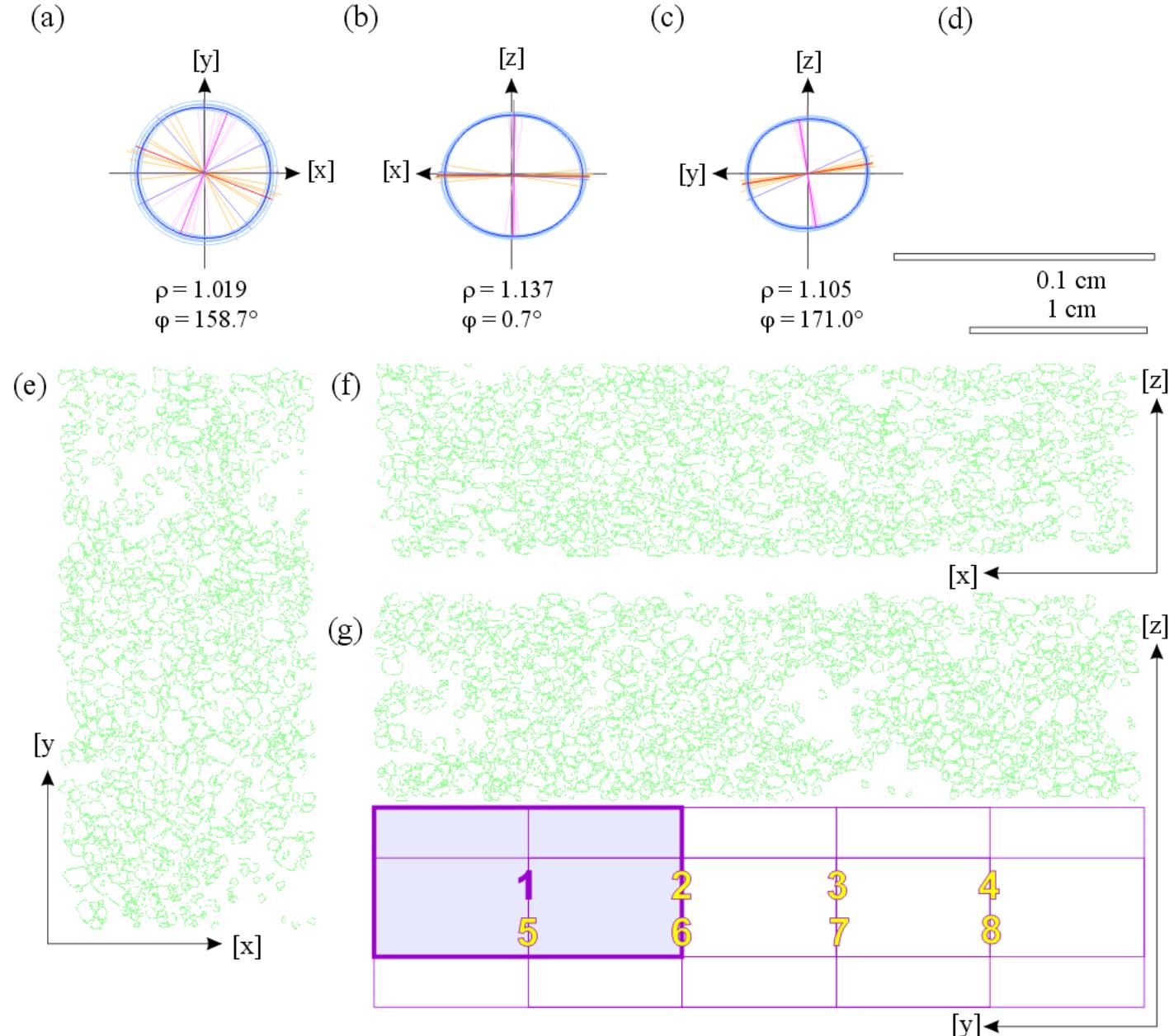
Quantitative Image Analysis of Minerals and Rocks

<http://www.sciences.univ-nantes.fr/lbanantes/SPO>

Application to gabbronorites of the Bushveld complex

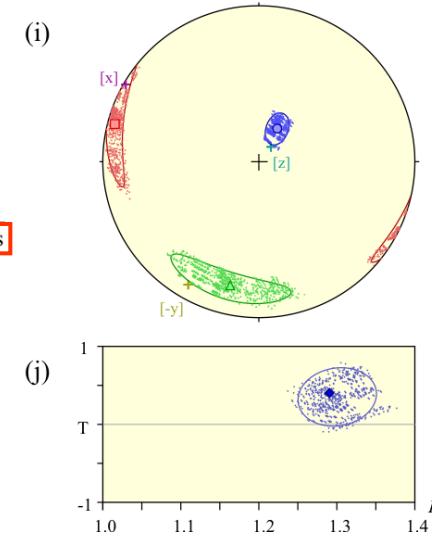
Orthopyroxene

P. Launeau, P.-Y. F. Robin
 (2005) "Determination of fabric and strain ellipsoids from measured sectional ellipses—implementation and applications". *Journal of Structural Geology*, 27, 2223-2233



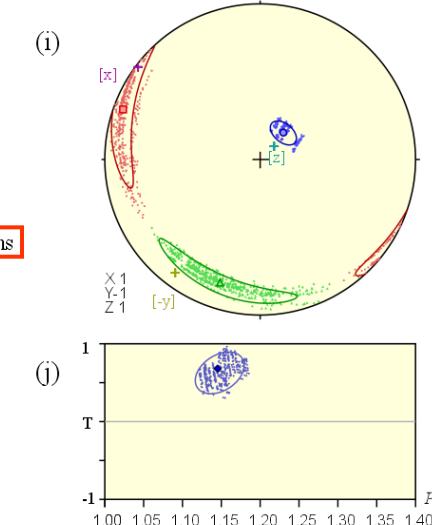
Application to gabbronorites of the Bushveld complex

(h)	A	B	C	\sqrt{F}	
L. norm	1.113	1.034	0.869	2.5%	
trend	284.7	192.9	29.0		
plunge	5.1	19.1	70.2		
A/C	1.282		Flinn	0.405	
A/B	1.077		P'	1.291	
B/C	1.192		T	0.396	
	A	B	C	729 combinations	
L. norm	1.120	1.030	0.867	\bar{X}	
	0.028	0.033	0.021	2σ	
trend	285.9	194.2	29.1	\bar{X}	
plunge	4.7	19.0	70.4	\bar{X}	
	27.4	27.4	8.8	$2\sigma_1$	
	5.6	8.8	5.7	$2\sigma_2$	
	\bar{X}	2σ	\bar{X}	2σ	
A/C	1.292	0.048	\sqrt{F}	2.6%	3.5%
A/B	1.087	0.057	P'	1.301	0.050
B/C	1.189	0.060	T	0.352	0.375



Inertia tensors

(h)	A	B	C	\sqrt{F}	
L. norm	1.050	1.029	0.926	1.4%	
trend	290.0	197.9	40.8		
plunge	6.8	17.2	71.5		
A/C	1.134		Flinn	0.179	
A/B	1.020		P'	1.145	
B/C	1.112		T	0.680	
	A	B	C	512 combinations	
L. norm	1.052	1.027	0.925	\bar{X}	
	0.011	0.014	0.015	2σ	
trend	289.6	197.6	41.3	\bar{X}	
plunge	6.9	16.8	71.5	\bar{X}	
	33.2	33.0	7.8	$2\sigma_1$	
	6.8	5.9	4.9	$2\sigma_2$	
	\bar{X}	2σ	\bar{X}	2σ	
A/C	1.138	0.027	\sqrt{F}	1.7%	2.1%
A/B	1.024	0.018	P'	1.148	0.031
B/C	1.111	0.031	T	0.625	0.268



Intercepts

P. Launeau, P.-Y. F. Robin (2005) "Determination of fabric and strain ellipsoids from measured sectional ellipses—implementation and applications". *Journal of Structural Geology*, 27, 2223-2233

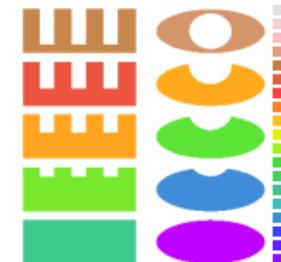
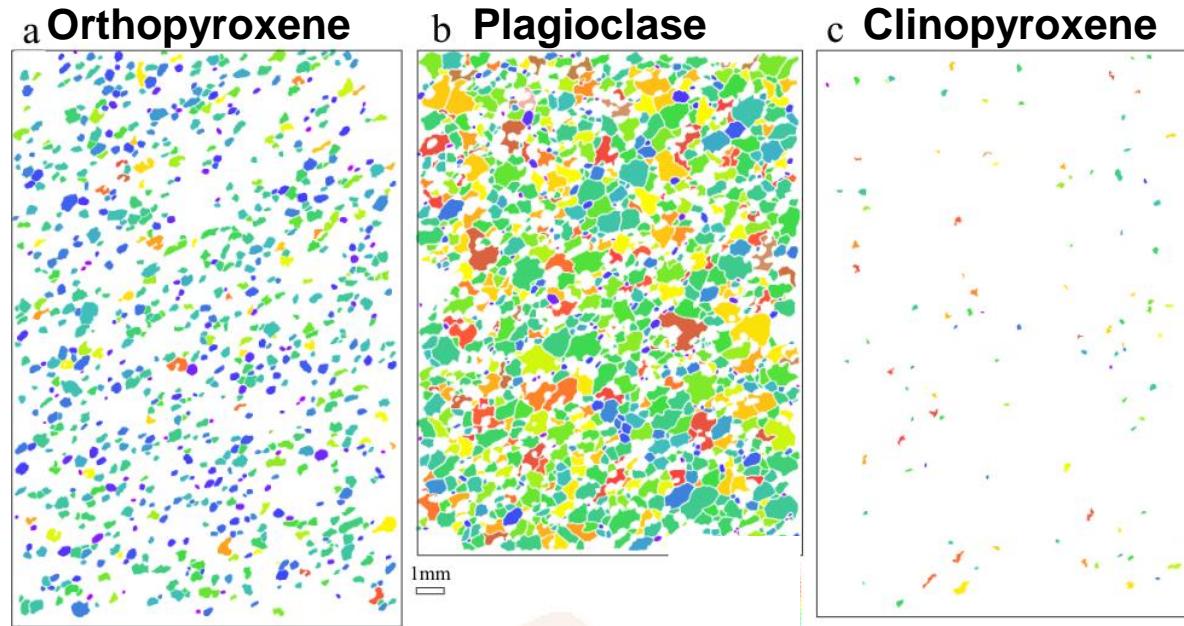
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<http://www.sciences.univ-nantes.fr/lpgnantes/SPO>

Application to gabbronorites of the Bushveld complex

P. Launeau (2004) "Mise en évidence des écoulements magmatiques par analyse d'images 2-D des distributions 3-D d'Orientations Préférentielles de Formes". *Bull. Soc. Géol. Fr.*, 175, 331-350

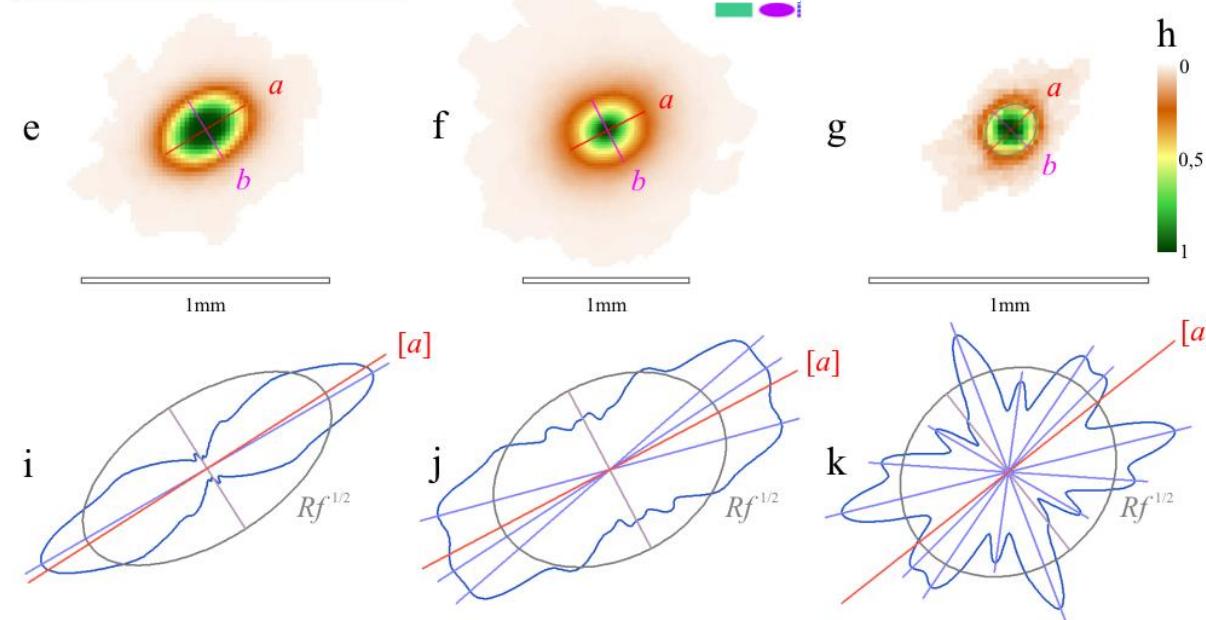
SPO



$$e_{ellipse} = \frac{\pi \cdot a \cdot b}{A} - 1$$

**pixel
density
of
crystal
stack**

PO

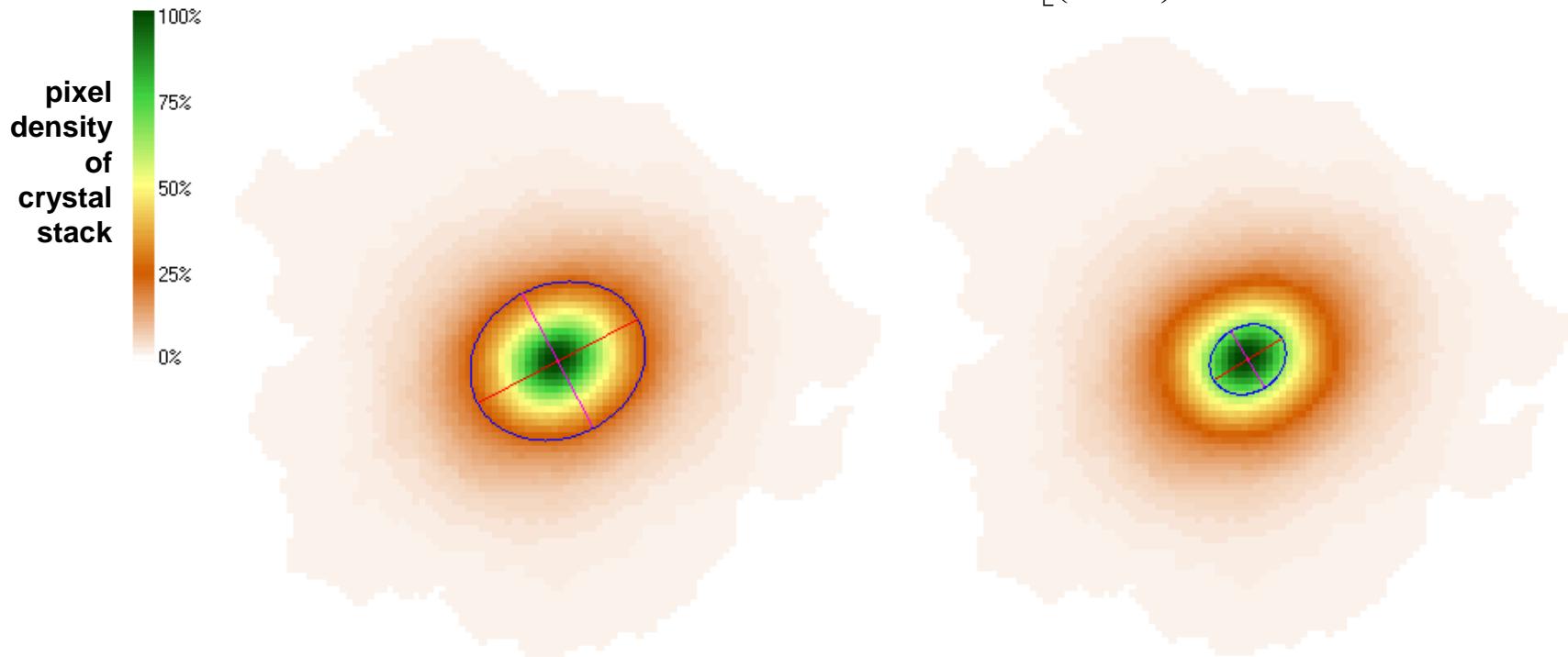


Application to gabbronorites of the Bushveld complex

Plagioclase

$$\mathbf{M}_T = \frac{1}{4 \cdot A} \begin{bmatrix} a^2 \cdot \cos^2 \varphi + b^2 \cdot \sin^2 \varphi & (a^2 - b^2) \cdot \cos \varphi \cdot \sin \varphi \\ (a^2 - b^2) \cdot \cos \varphi \cdot \sin \varphi & a^2 \cdot \sin^2 \varphi + b^2 \cdot \cos^2 \varphi \end{bmatrix}$$

$$\mathbf{M}_E = A \cdot \begin{bmatrix} \frac{1}{a^2} \cdot \cos^2 \varphi + \frac{1}{b^2} \cdot \sin^2 \varphi & \left(\frac{1}{a^2} - \frac{1}{b^2}\right) \cdot \cos \varphi \cdot \sin \varphi \\ \left(\frac{1}{a^2} - \frac{1}{b^2}\right) \cdot \cos \varphi \cdot \sin \varphi & \frac{1}{a^2} \cdot \sin^2 \varphi + \frac{1}{b^2} \cdot \cos^2 \varphi \end{bmatrix}$$



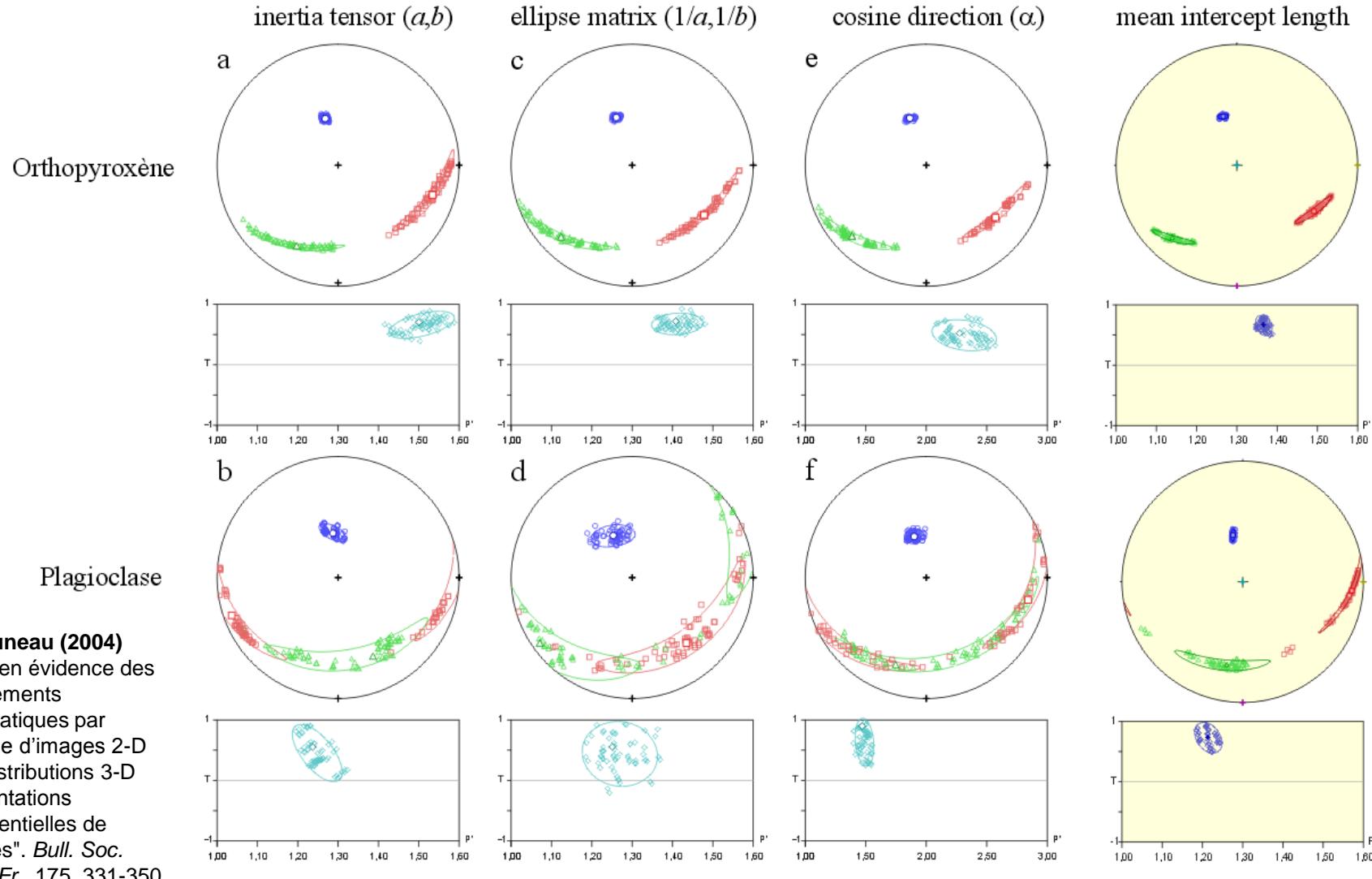
0.1 cm

P. Launeau (2004) "Mise en évidence des écoulements magmatiques par analyse d'images 2-D des distributions 3-D d'Orientations Préférentielles de Formes". *Bull. Soc. Géol. Fr.*, 175, 331-350

Quantitative Image Analysis of Minerals and Rocks

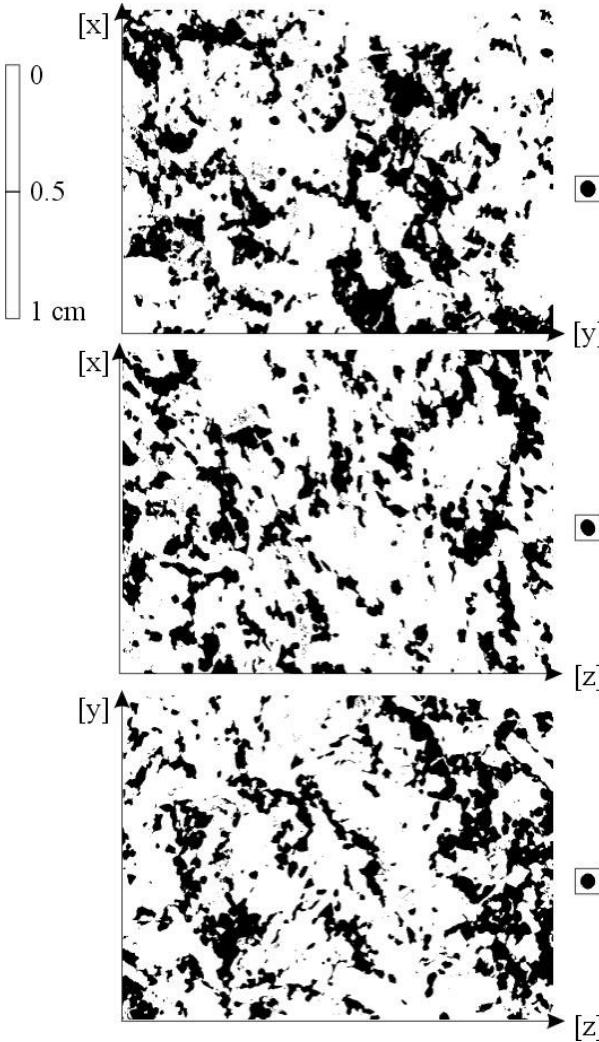
<http://www.sciences.univ-nantes.fr/lpgnantes/SPO>

Application to gabbronorites of the Bushveld complex



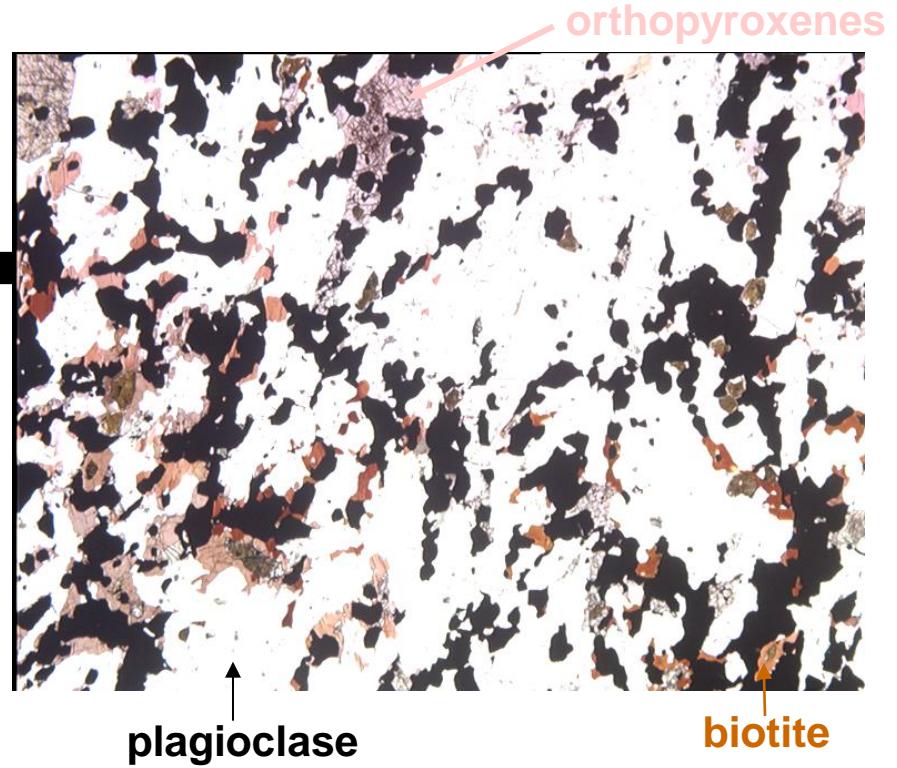
P. Launeau (2004)
 "Mise en évidence des
 écoulements
 magmatiques par
 analyse d'images 2-D
 des distributions 3-D
 d'orientations
 préférentielles de
 formes". *Bull. Soc.
 Géol. Fr.*, 175, 331-350

Application to ilmenite-rich norite

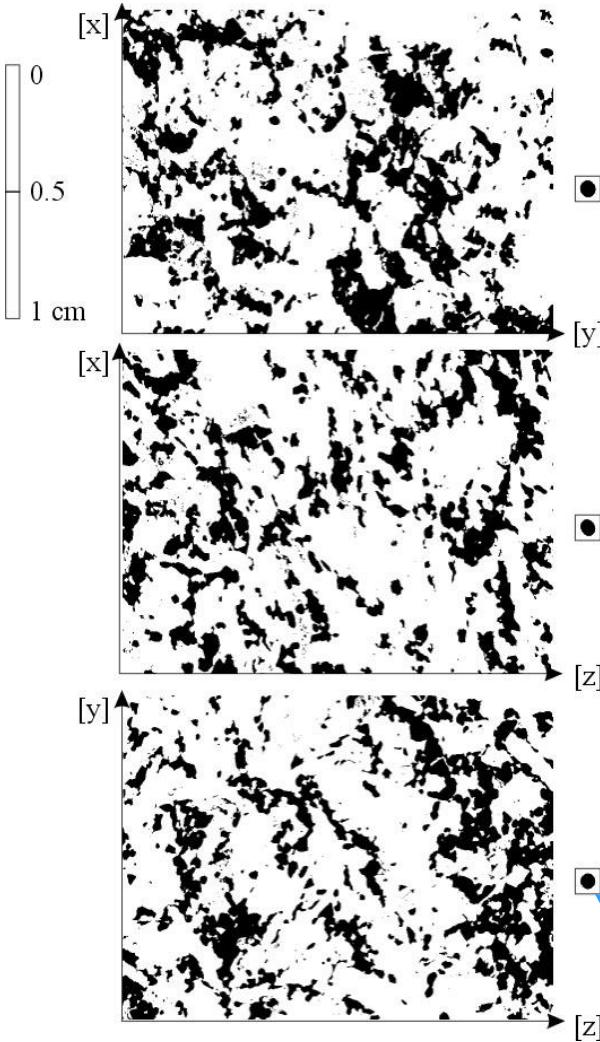


SPO with intercepts of one phase made of ilmenite and magnetite (Tellnes, Norway) H. Diot et al. 2003

Grey level threshold on dark minerals

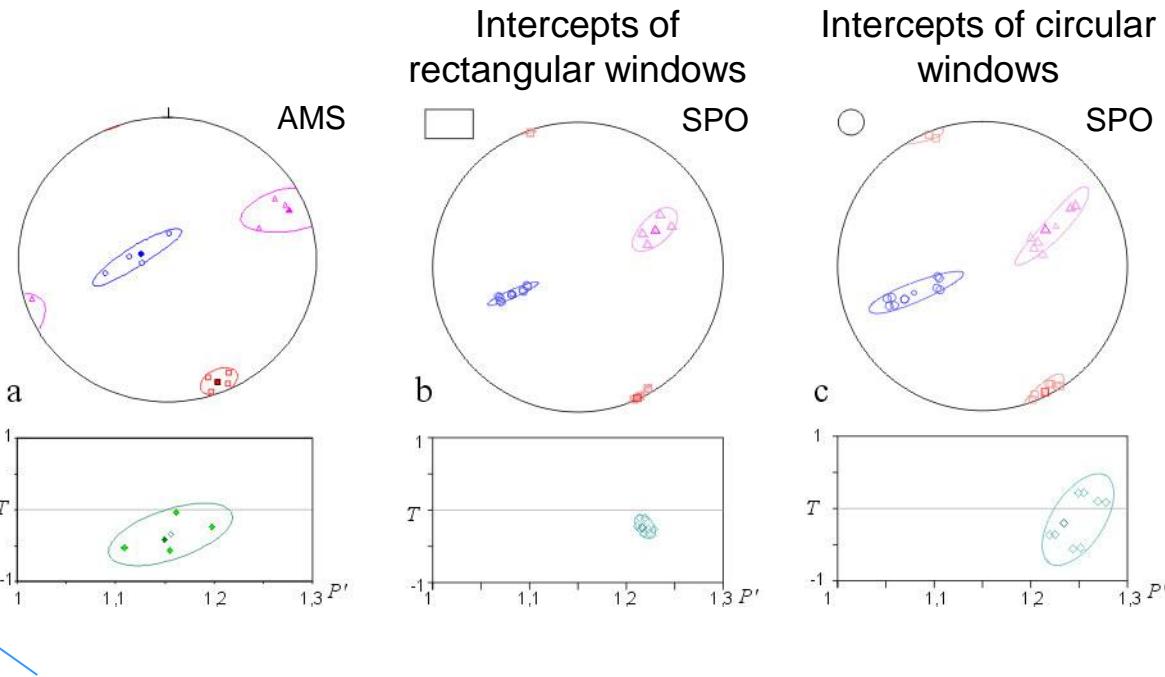


Application to ilmenite-rich norite



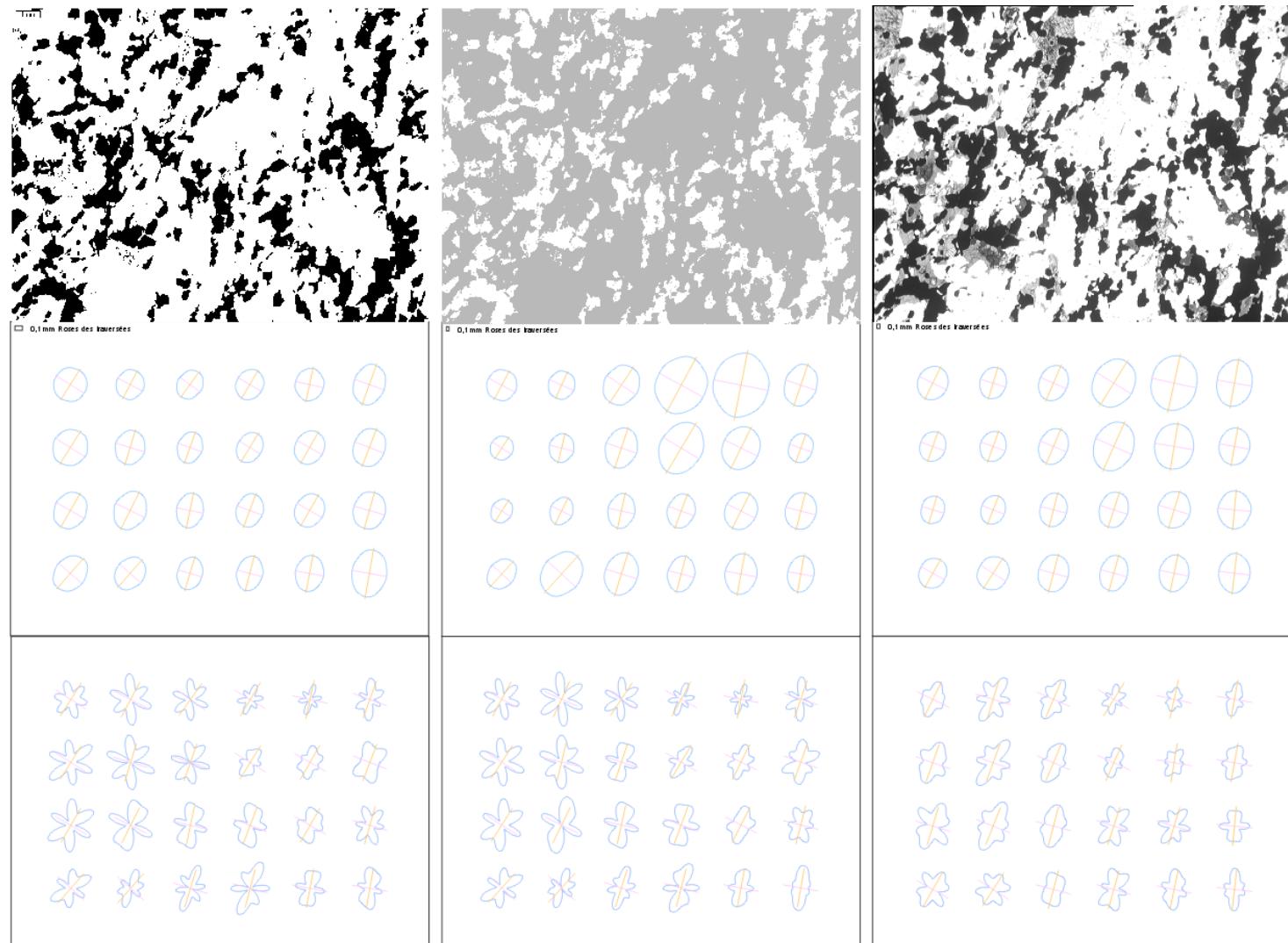
SPO with intercepts of one phase made of ilmenite
(Tellnes, Norway) H. Diot et al. 2003

Grey level threshold on dark minerals

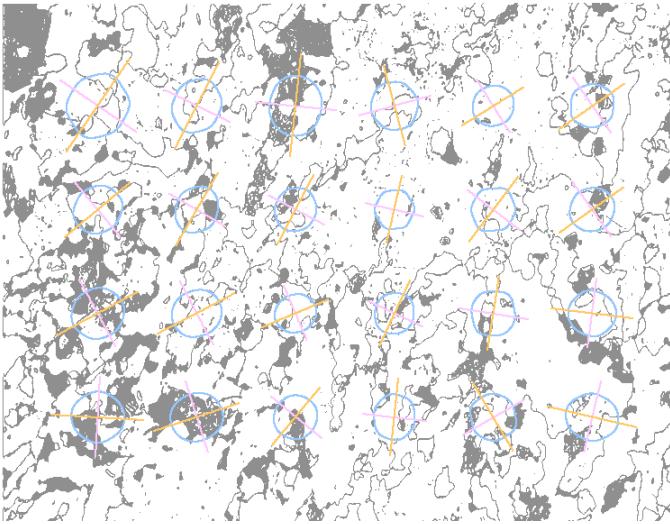
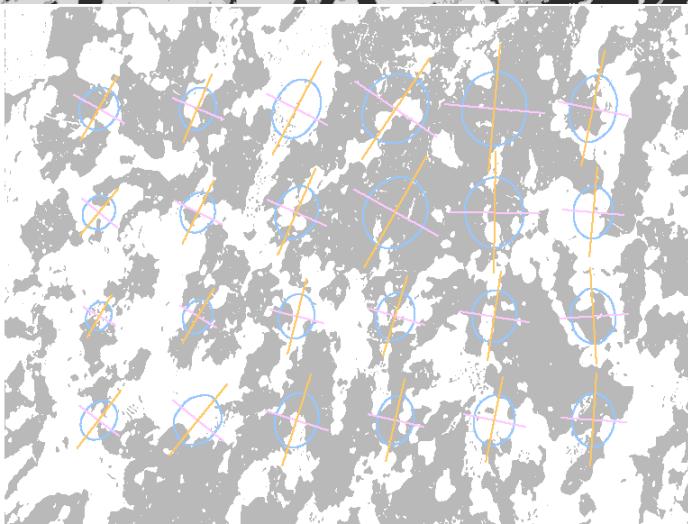
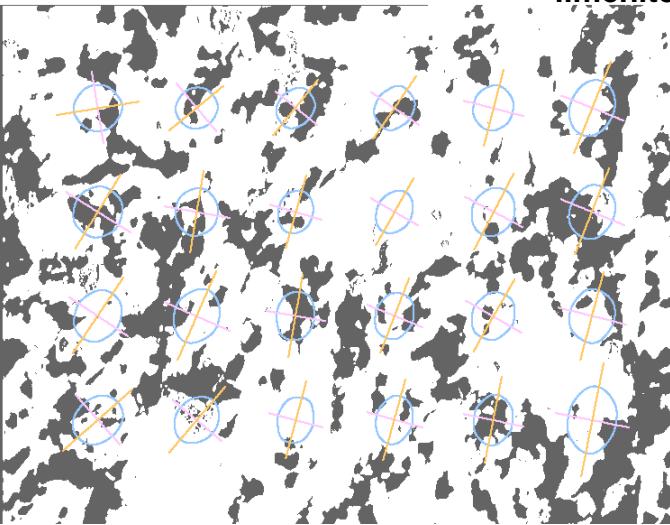
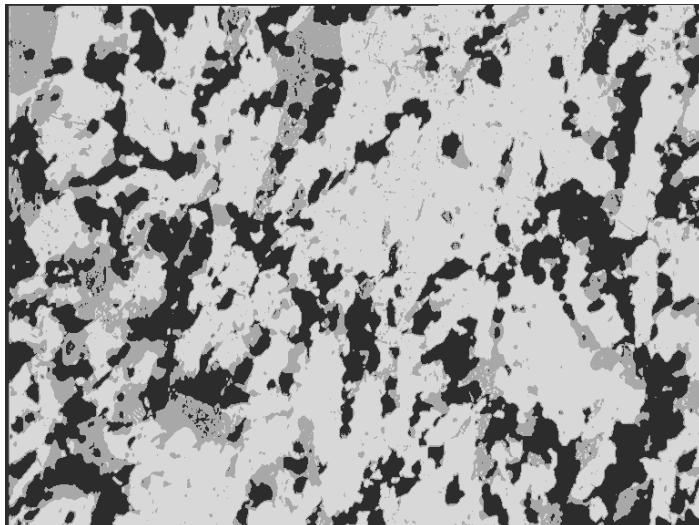


roses of mean intercept length

Application to ilmenite-rich norite

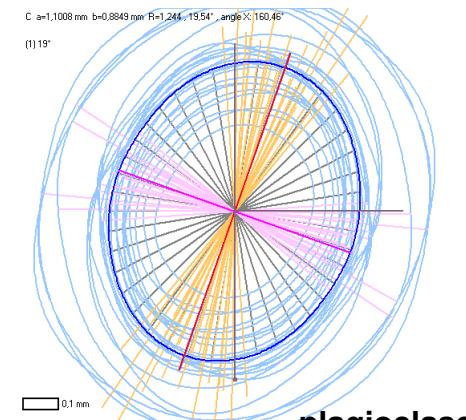
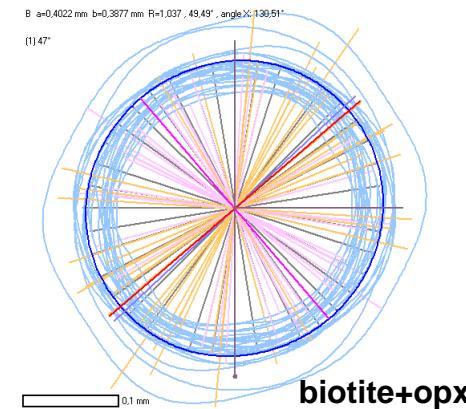
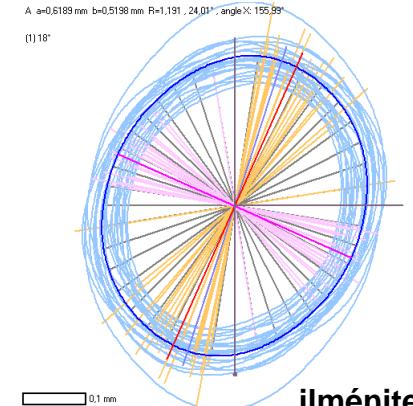


Application to ilmenite-rich norite

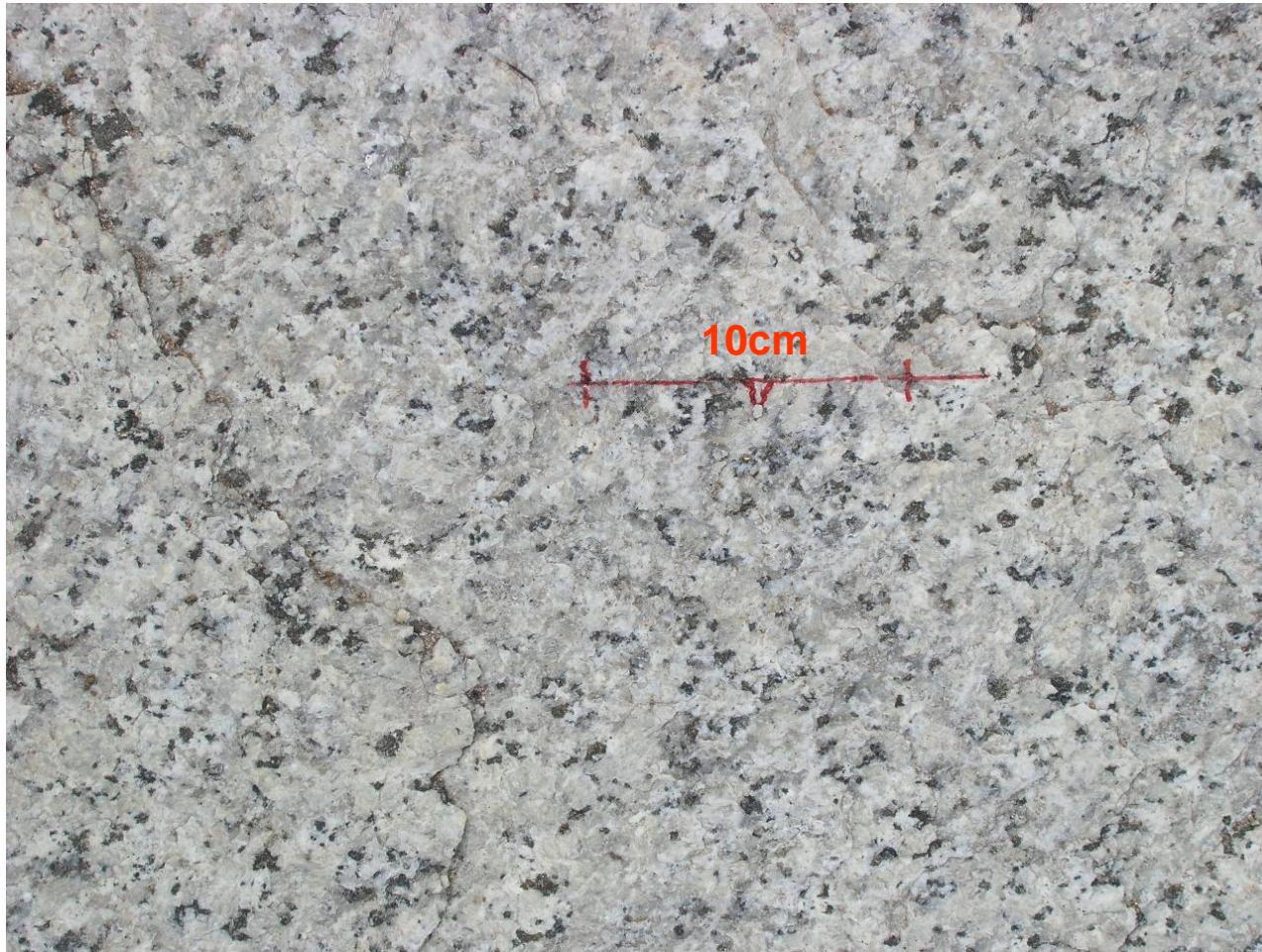


plagioclase

H. Diot, O. Bolle, J.-M. Lambert, P. Launeau and J.-C. Duchesne (2003) The Tellnes ilmenite deposit (Rogaland, South Norway): magnetic and petrofabric evidence for emplacement of a Ti-enriched noritic crystal mush in a fracture zone, *Journal of Structural Geology* 25, 481–501

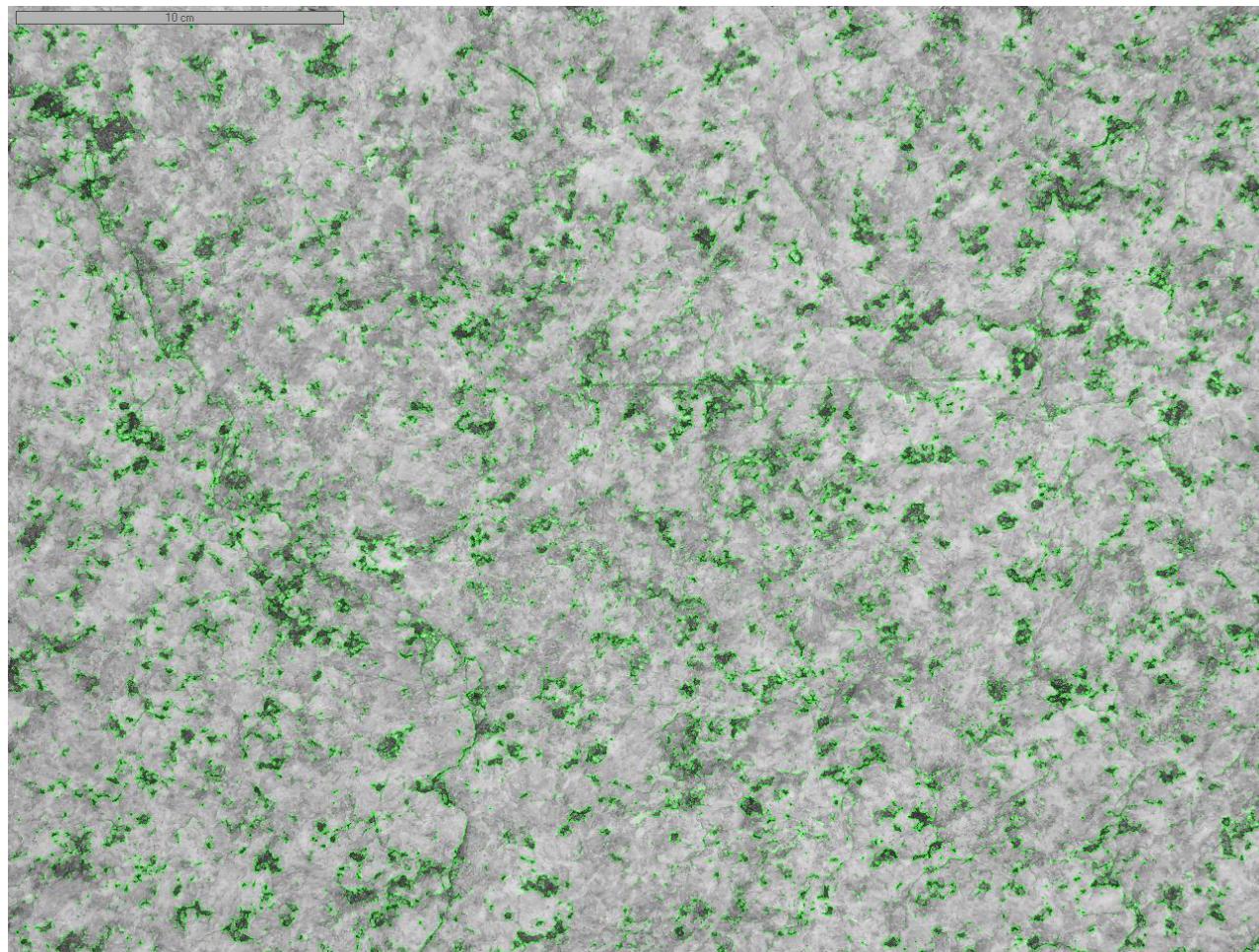


Application to the Pocinhos granite (Brazil)

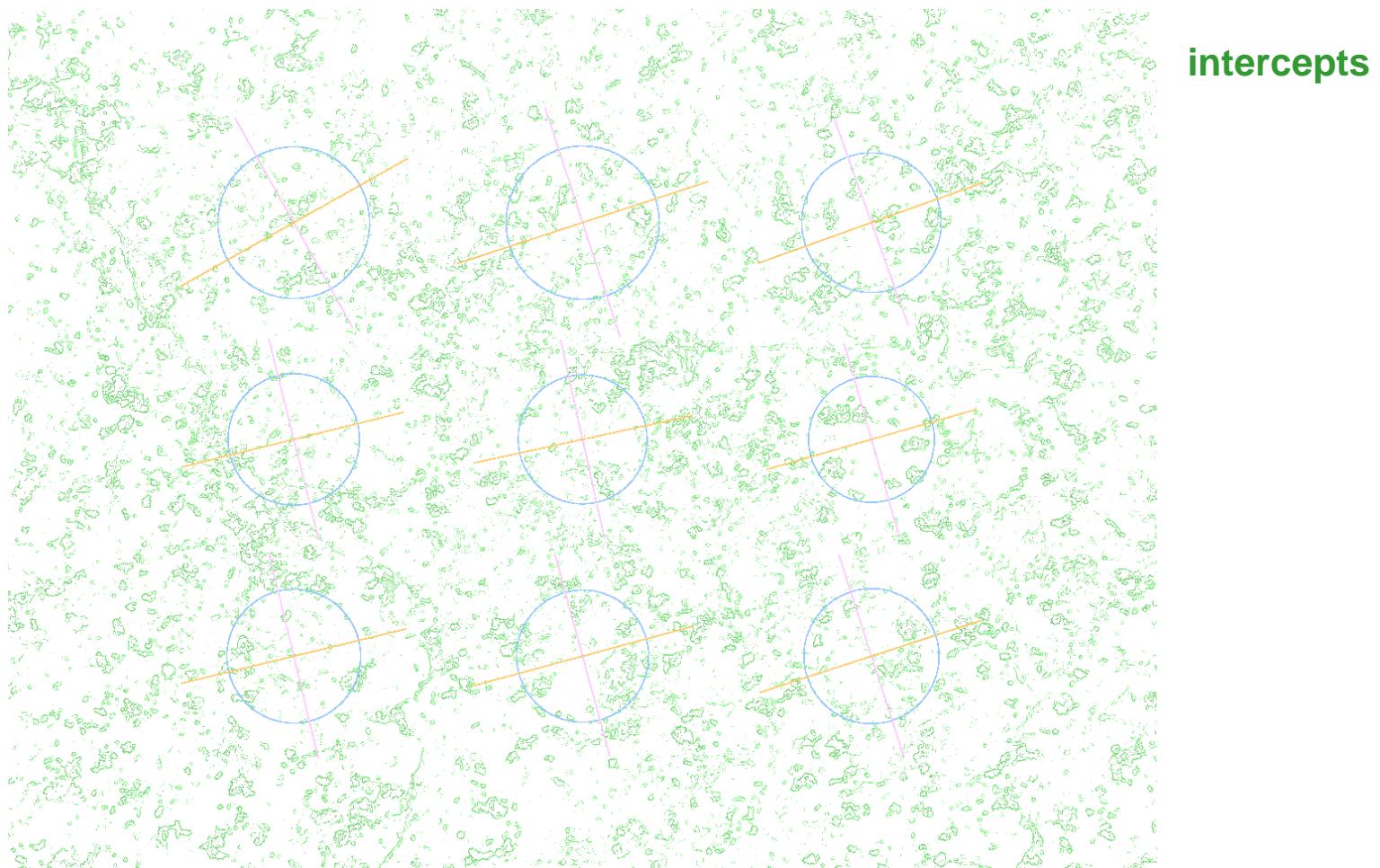


RGB

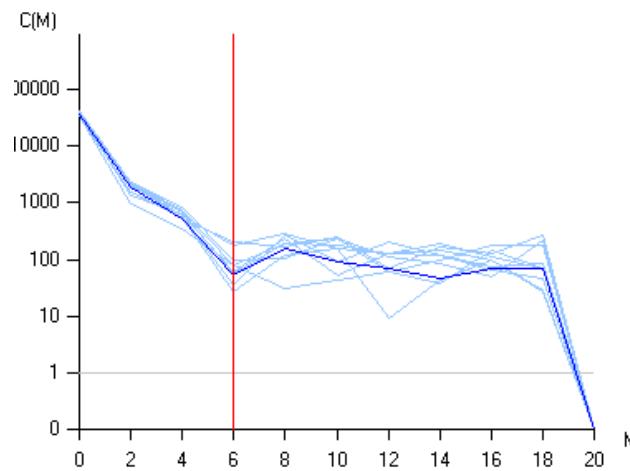
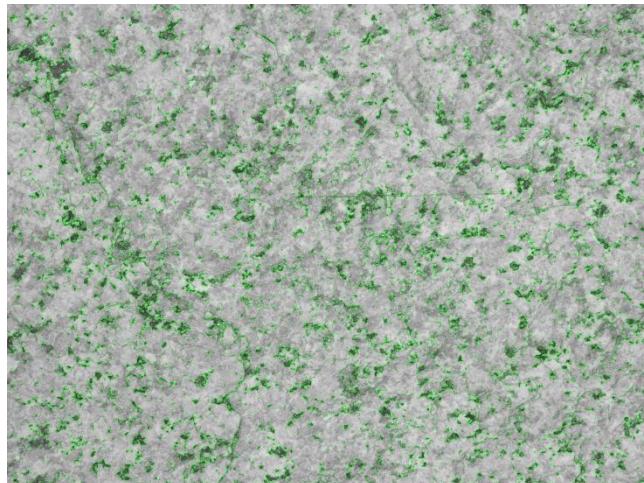
Application to the Pocinhos granite (Brazil)



Application to the Pocinhos granite (Brazil)

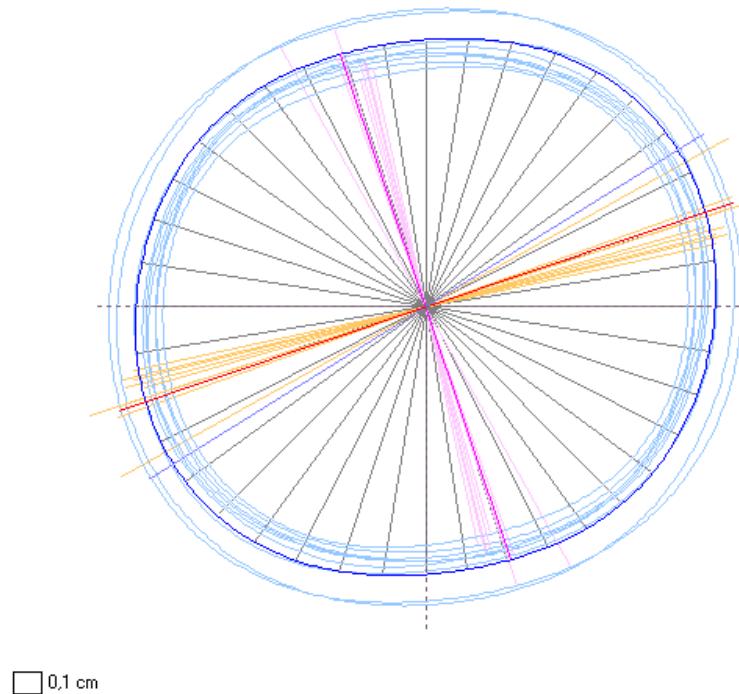


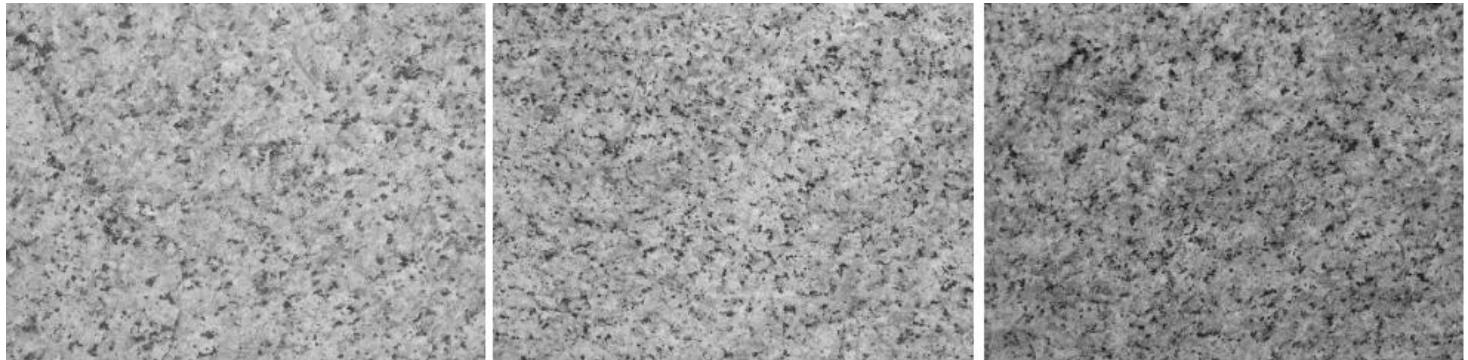
Application to the Pocinhos granite (Brazil)



A a=2,7108 cm b=2,4486 cm R=1,107 , 71,20° , angle X: 161,20°

(1) 58°

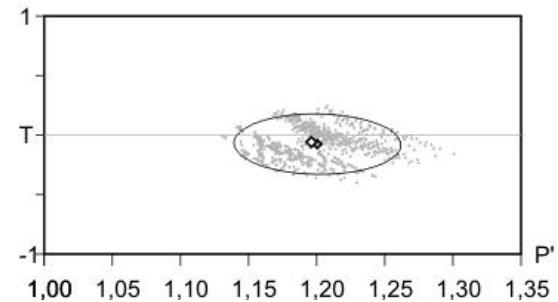
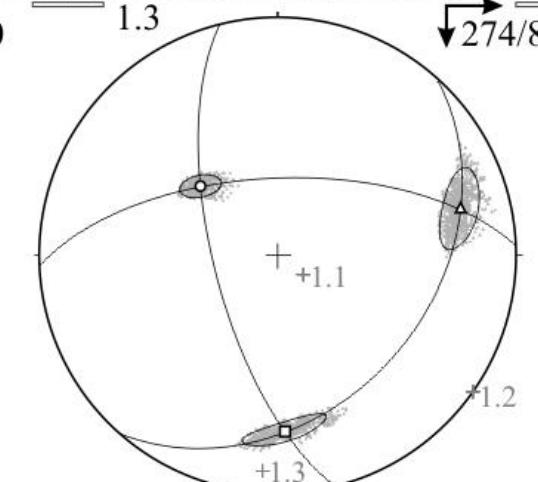




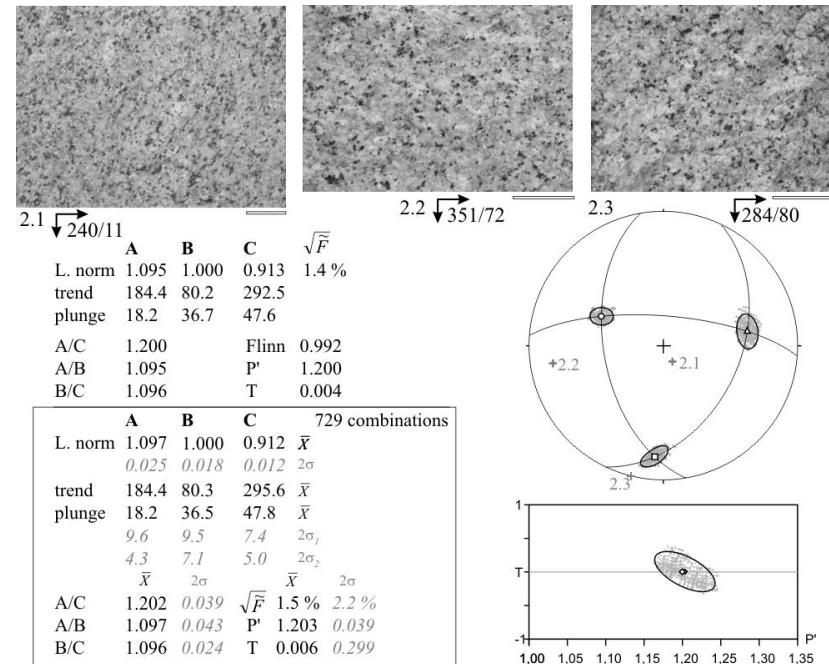
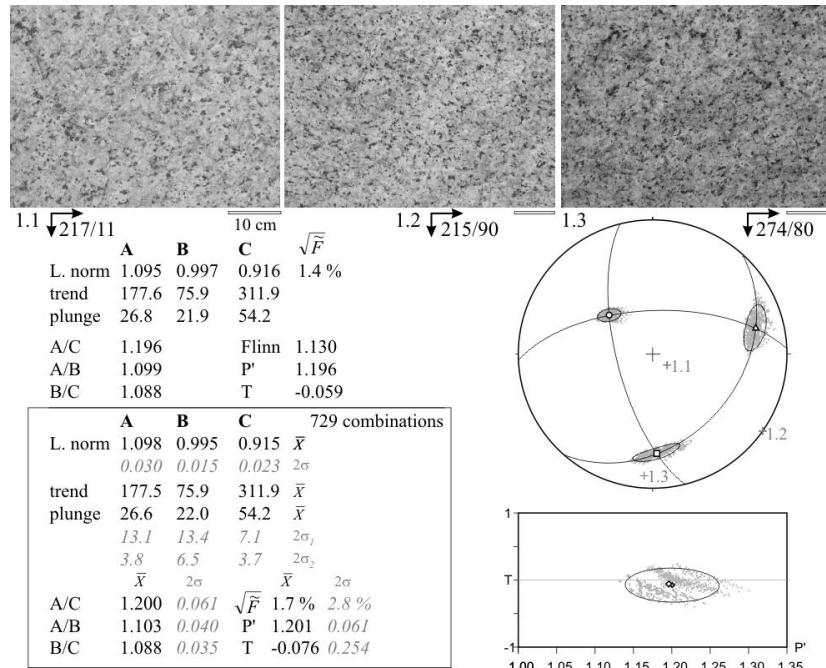
1.1 $\downarrow 217/11$ 10 cm 1.2 $\downarrow 215/90$ 1.3 $\downarrow 274/80$

Application to the Pocinhos granite (Brazil)

	A	B	C	\sqrt{F}	729 combinations
L. norm	1.095	0.997	0.916	1.4 %	
trend	177.6	75.9	311.9		
plunge	26.8	21.9	54.2		
A/C	1.196		Flinn	1.130	
A/B	1.099		P'	1.196	
B/C	1.088		T	-0.059	
L. norm	1.098	0.995	0.915	\bar{X}	
	0.030	0.015	0.023	2σ	
trend	177.5	75.9	311.9	\bar{X}	
plunge	26.6	22.0	54.2	\bar{X}	
	13.1	13.4	7.1	$2\sigma_1$	
	3.8	6.5	3.7	$2\sigma_2$	
	\bar{X}	2σ	\bar{X}	2σ	
A/C	1.200	0.061	\sqrt{F}	1.7 %	2.8 %
A/B	1.103	0.040	P'	1.201	0.061
B/C	1.088	0.035	T	-0.076	0.254

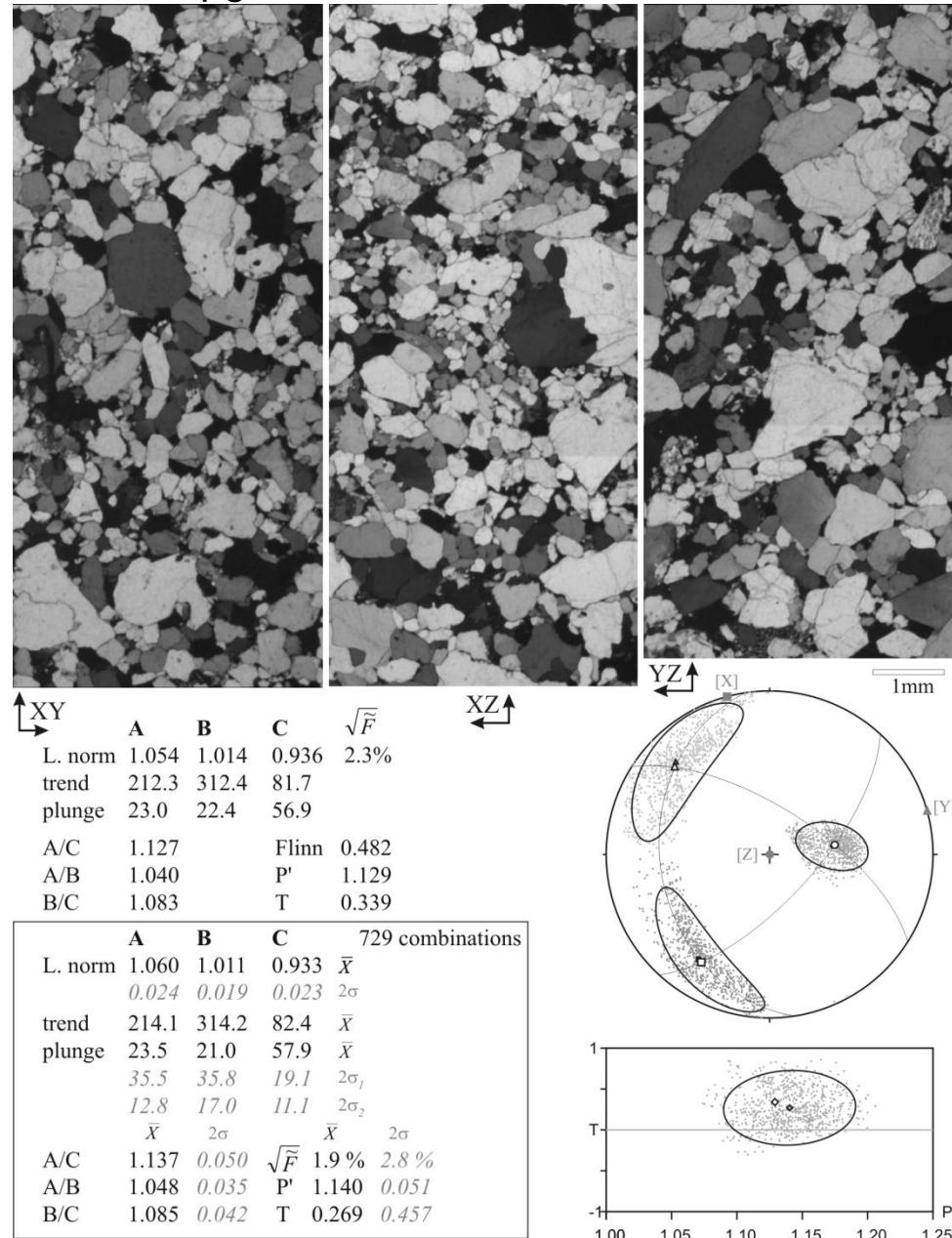


Application to the Pocinhos granite (Brazil)



Application to sandstones

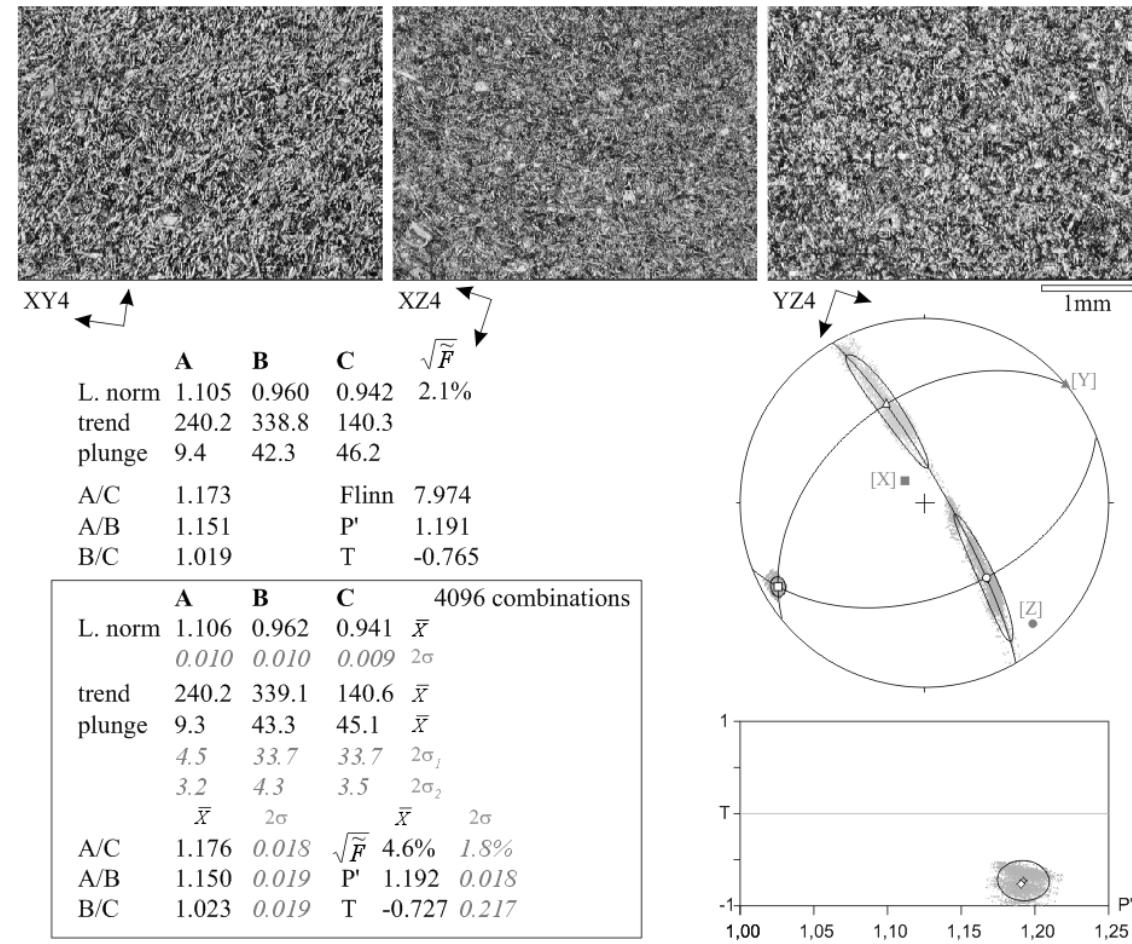
Application to maximum polarized light,
using Fueten and Goodchild (2001)
methodology; thin sections of
sandstones from the Devonian beds of
the Furnas Formation (Brazil)



Launeau P., Archanjo C. J., Picard D., Arbaret L., Robin P.Y. (2010). Two- and three-dimensional shape fabric analysis by the intercept method in grey levels. Tectonophysics, Volume 492, Issues 1-4, 20 September 2010, Pages 230-239

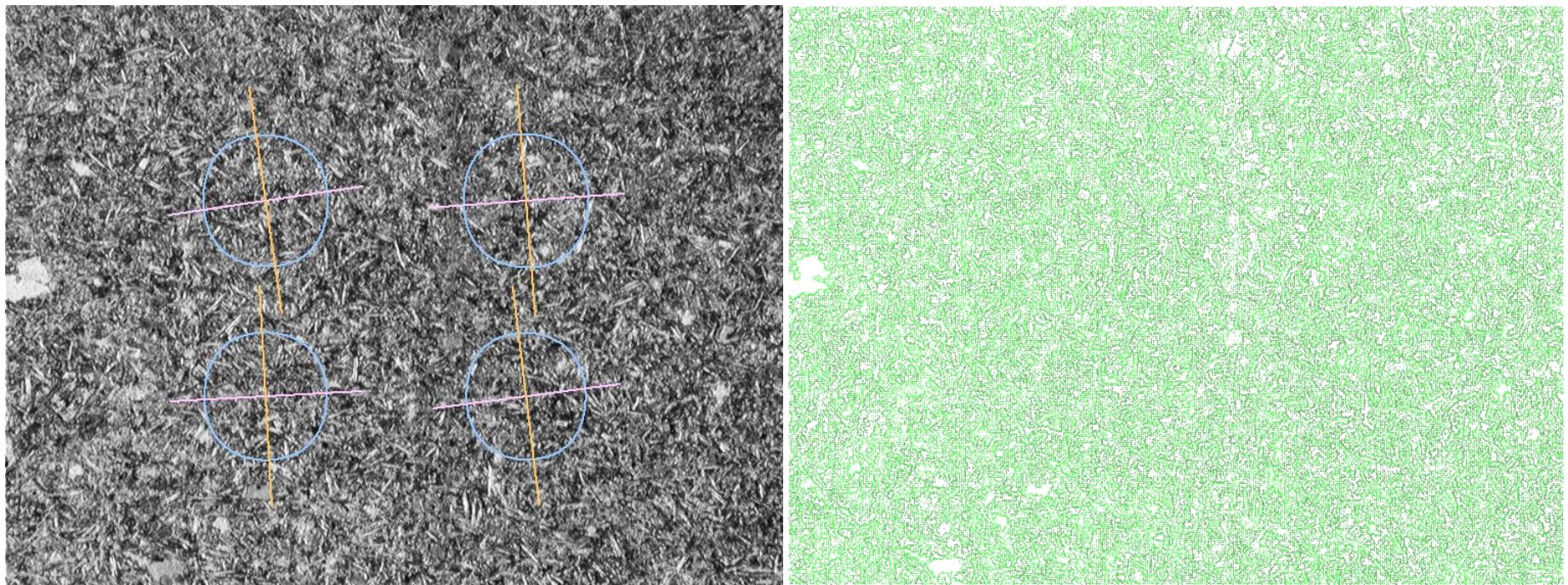
Application to a chilled margin of a diabase dike

Application to maximum polarized light, using Fueten and Goodchild (2001) methodology; The specimen comes from about 3 cm from the margin of a Mesozoic diabase dike (Rio Ceará-Mirim swarm, NE Brazil)



Application to a chilled margin of a diabase dike

Application to maximum polarized light, using Fueten and Goodchild (2001) methodology; The specimen comes from about 3 cm from the margin of a Mesozoic diabase dike (Rio Ceará-Mirim swarm, NE Brazil)



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